

# Characterizing Carbon Sink Responses to Decarbonization across Model Structures

Greta Shum<sup>1</sup>, James Yoon<sup>1</sup>, Olivia Truax<sup>1,2</sup>, Claire Zarakas<sup>1</sup>, Dargan Frierson<sup>1</sup>, Charlie Koven<sup>3</sup>, and Abby Swann<sup>1,4</sup>

<sup>1</sup> Atmospheric Sciences Department, University of Washington, Seattle, WA

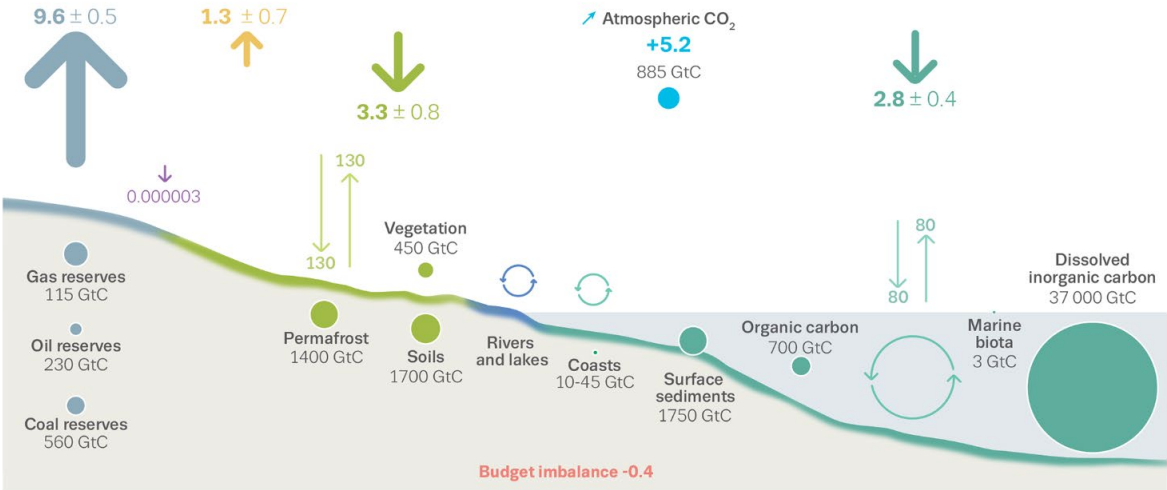
<sup>2</sup> School of Earth and Environment, University of Canterbury, Christchurch, NZ

<sup>3</sup> Lawrence Berkeley National Lab, Berkeley, CA

<sup>4</sup> Biology Department, University of Washington, Seattle, WA

# Biogeochemical feedbacks dictate the net land and ocean carbon sinks

## The global carbon cycle

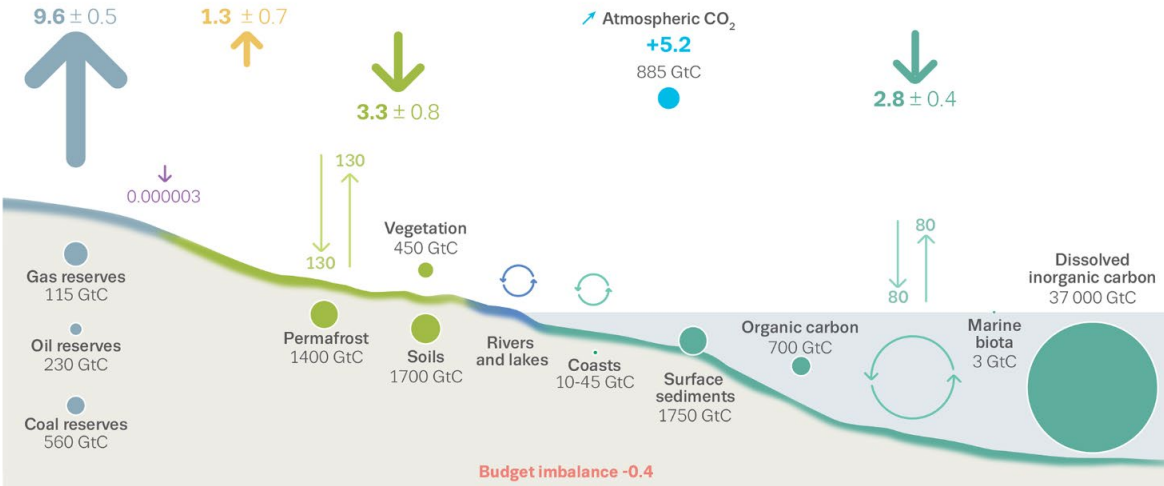


Anthropogenic fluxes 2013-2022 average GtC per year

- ↑ Fossil CO<sub>2</sub> E<sub>FOS</sub>
- ↑ Land-use change E<sub>LUC</sub>
- ↓ CDR not included in E<sub>LUC</sub>
- ↓ Land uptake S<sub>LAND</sub>
- ↓ Ocean uptake S<sub>OCEAN</sub>
- ↑ Carbon cycling GtC per year
- Stocks GtC
- + Atmospheric increase G<sub>ATM</sub>
- Budget Imbalance B<sub>IM</sub>

# Biogeochemical feedbacks dictate the net land and ocean carbon sinks

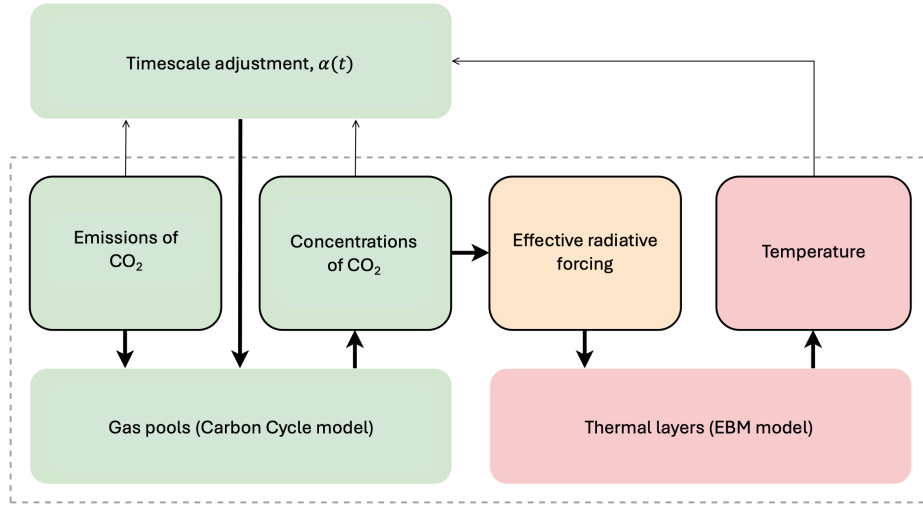
## The global carbon cycle



Anthropogenic fluxes 2013-2022 average GtC per year

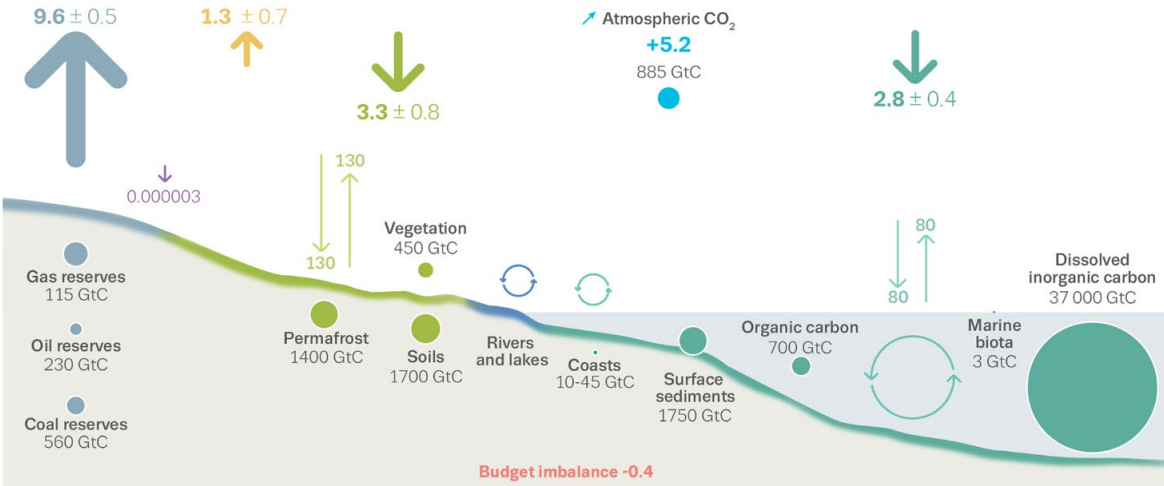
- ↑ Fossil CO<sub>2</sub> E<sub>FOS</sub>
- ↑ Land-use change E<sub>LUC</sub>
- ↓ CDR not included in E<sub>LUC</sub>
- ↓ Land uptake S<sub>LAND</sub>
- ↓ Ocean uptake S<sub>OCEAN</sub>
- ↑ Carbon cycling GtC per year
- Stocks GtC
- + Atmospheric increase G<sub>ATM</sub>
- Budget Imbalance B<sub>IM</sub>

## Structure of FaIR (Simple Climate Model)



# Biogeochemical feedbacks dictate the net land and ocean carbon sinks

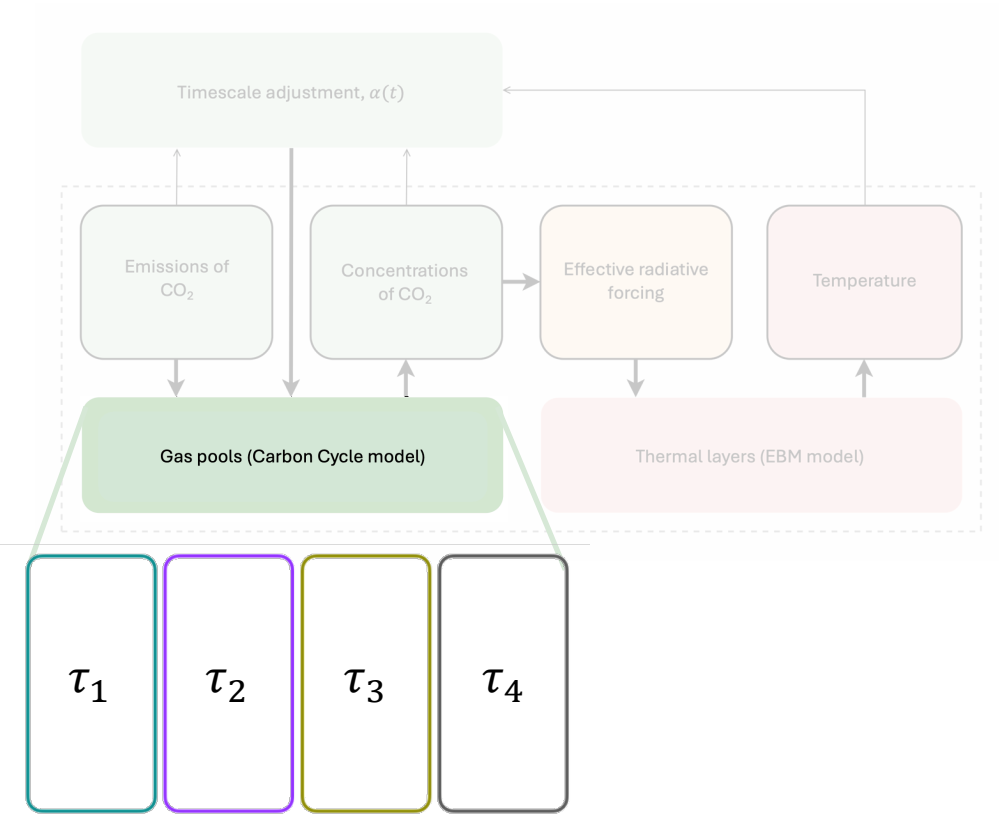
## The global carbon cycle



**Anthropogenic fluxes 2013-2022 average GtC per year**

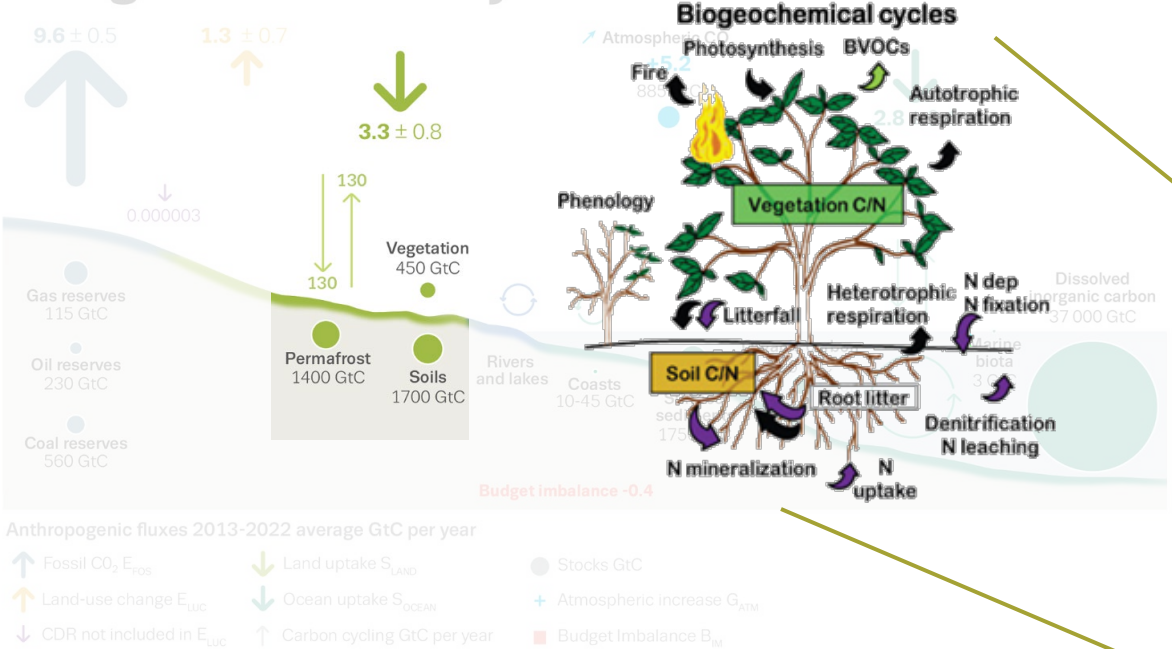
- ↑ Fossil CO<sub>2</sub> E<sub>FOS</sub>
- ↑ Land-use change E<sub>LUC</sub>
- ↓ CDR not included in E<sub>LUC</sub>
- ↓ Land uptake S<sub>LAND</sub>
- ↓ Ocean uptake S<sub>OCEAN</sub>
- ↑ Carbon cycling GtC per year
- Stocks GtC
- + Atmospheric increase G<sub>ATM</sub>
- Budget Imbalance B<sub>IM</sub>

## Structure of FaIR (Simple Climate Model)

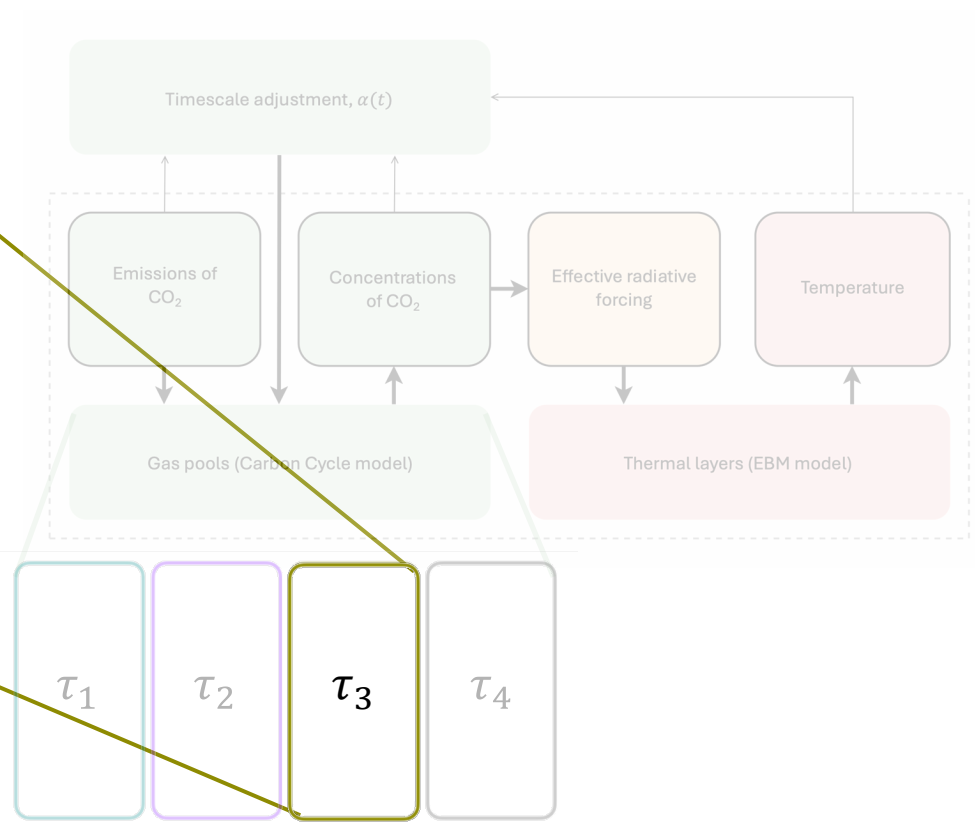


# Biogeochemical feedbacks dictate the net land and ocean carbon sinks

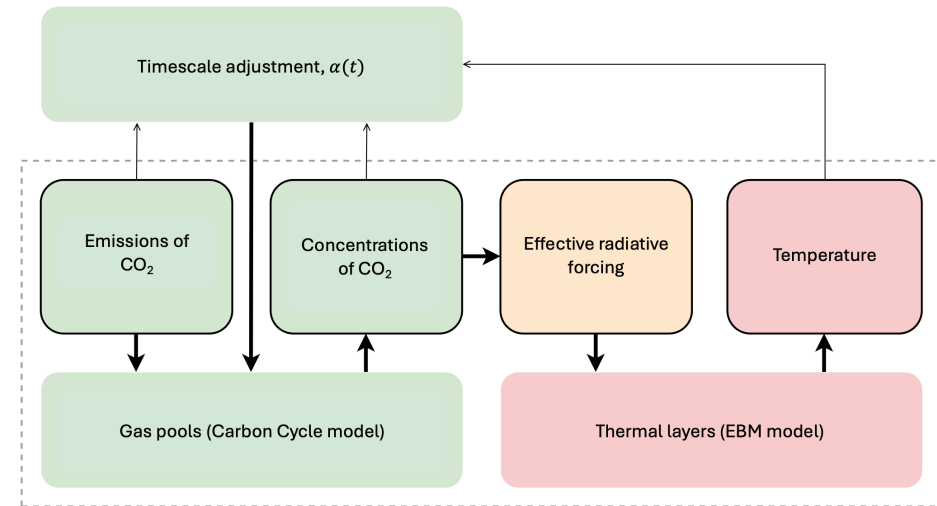
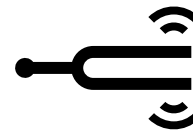
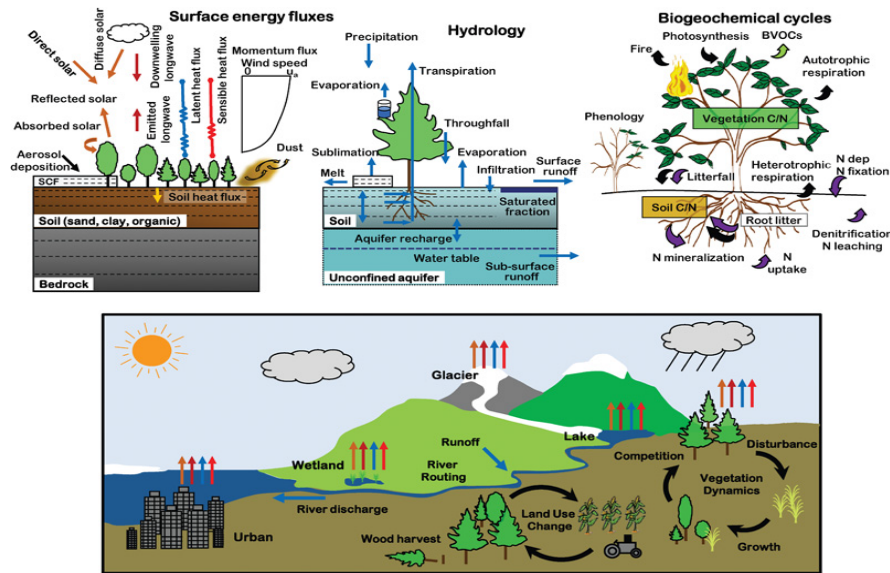
## The global carbon cycle



## Structure of FaIR (Simple Climate Model)



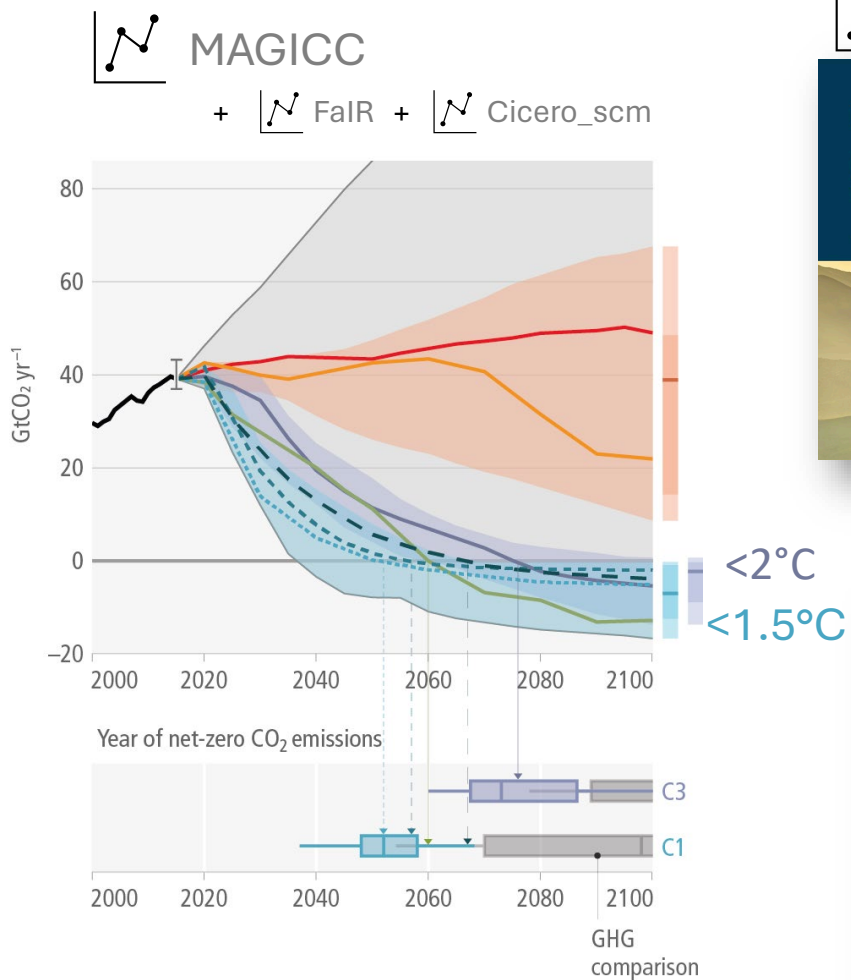
# SCMs are tuned to observations and ESMs, only ESMs in the future



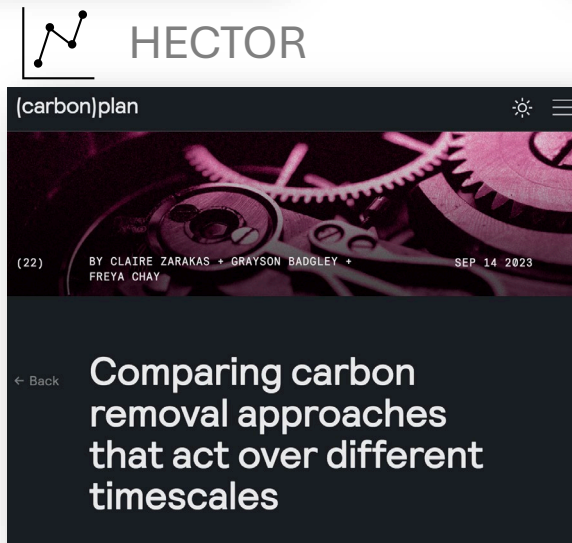
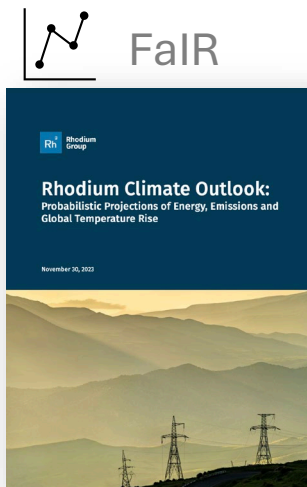
ESM process representation

SCMS parameters are tuned to reproduce the ESM output (using CMIP scenarios for future)

# SCMs increasingly used to assess and interpret climate mitigation



IPCC AR6 WGIII Figure 3.6



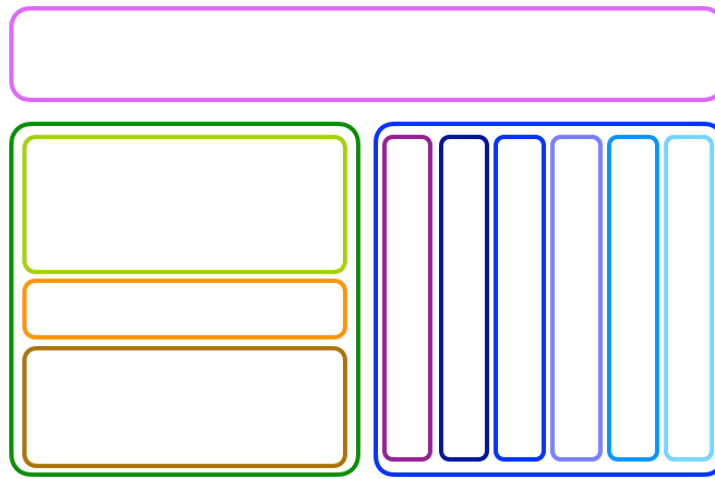


# Structure of the carbon cycle inside these SCMs varies a lot

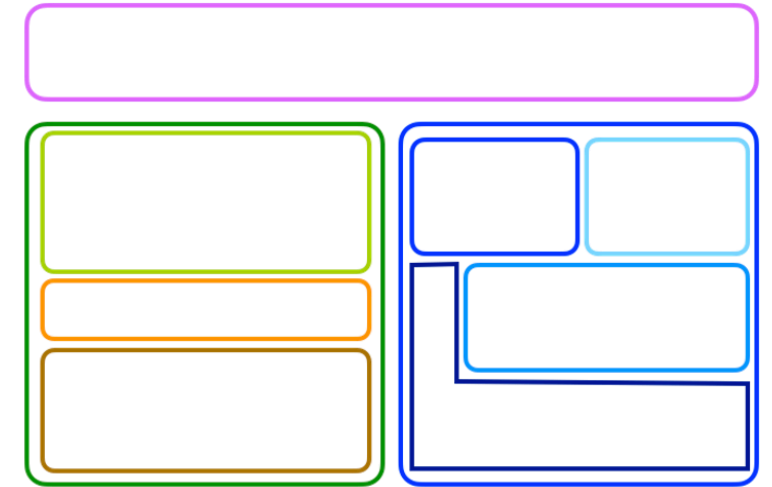
**FaIR**



**MAGICC**



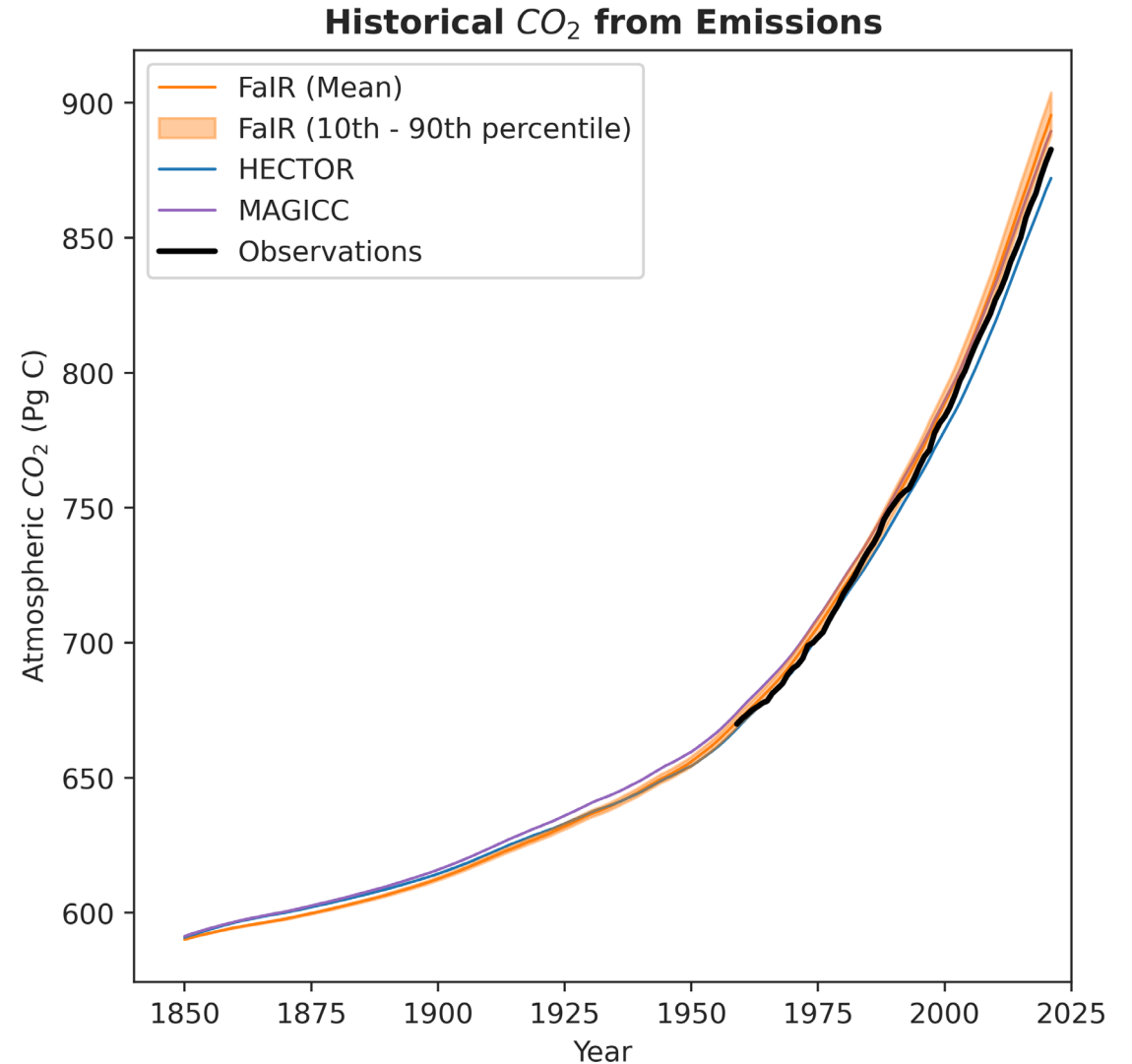
**HECTOR**





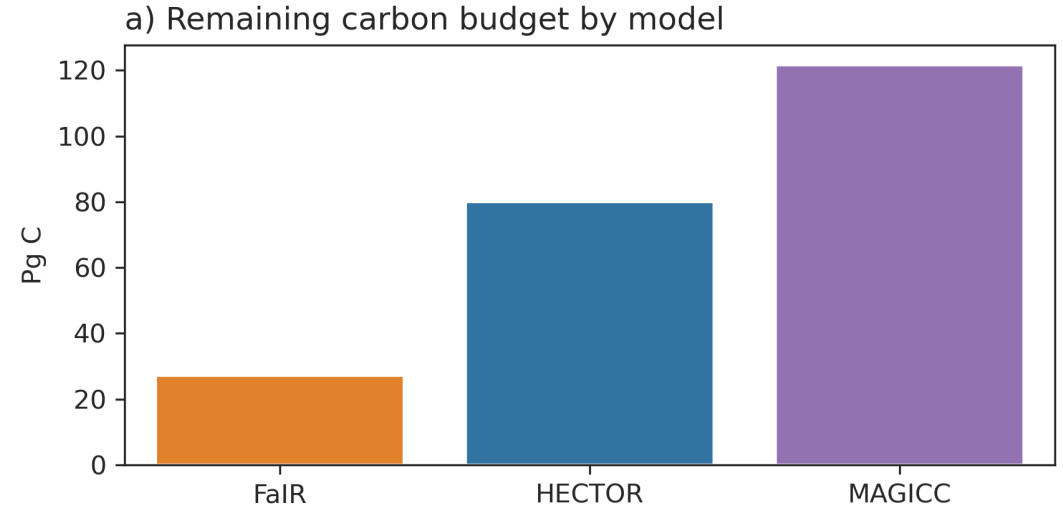
# Models reproduce historical emissions to concentration

- Out of the box, models can reproduce observed record of CO<sub>2</sub> concentrations since pre-industrial
  - FF and LUC emissions data from Global Carbon Budget 2021 (Friedlingstein et al., 2022)
- Demonstrates their ability to capture carbon cycle in an emissions regime.



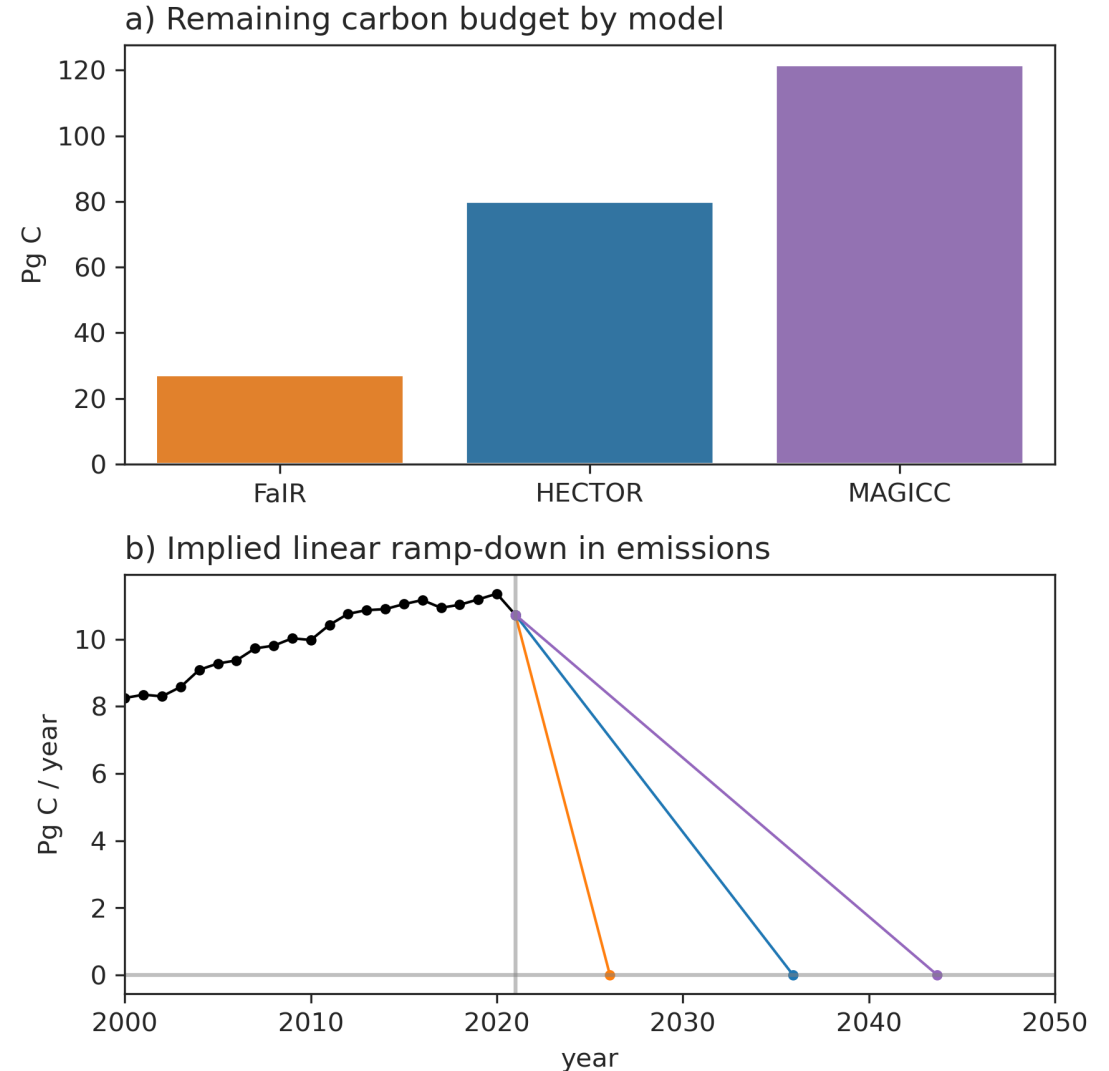
# SCMs vary by ~100 Pg C in remaining carbon budget

- Using each model's TCRE and ZEC, we can compute remaining carbon budgets for a temperature change limit of 1.5°C.

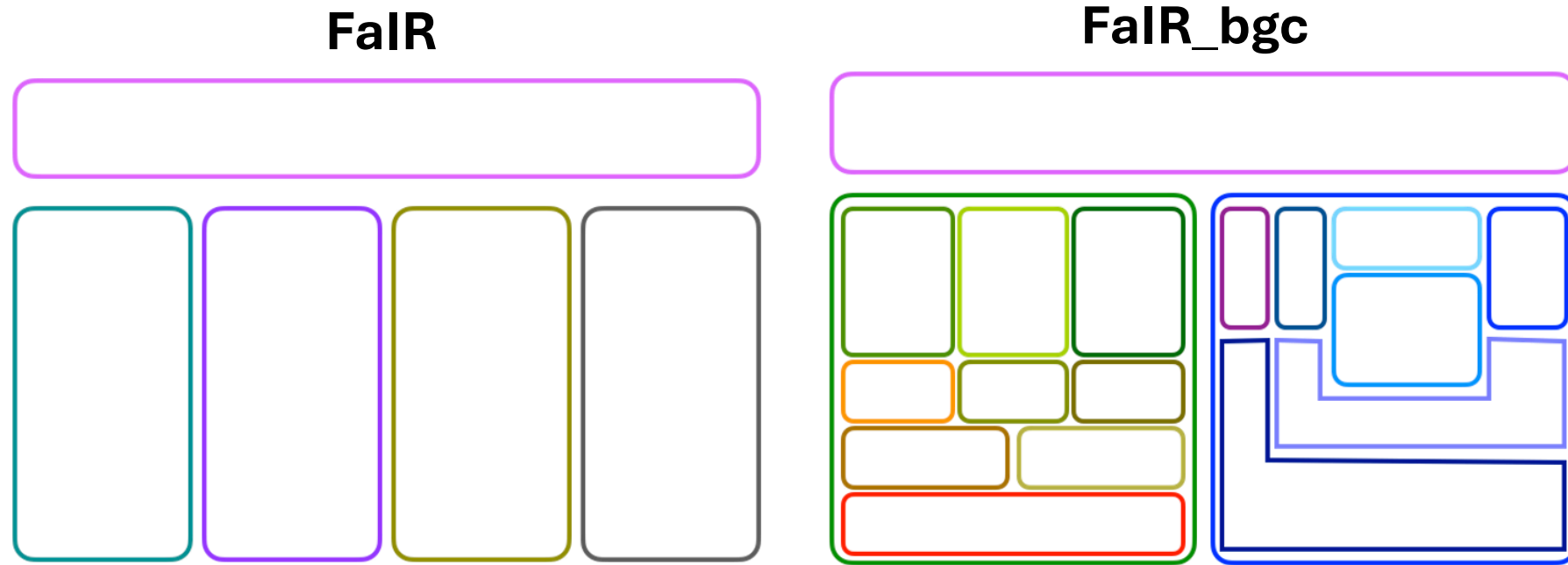


# SCMs show spread in remaining carbon budget and net-zero date

- Using each model's TCRE and ZEC, we can compute remaining carbon budgets for a temperature change limit of 1.5°C.
- From these budgets, we can project simplified linear ramp-down trajectories to estimate time-to-net-zero emissions.
- RCBs vary by ~100 Pg C and 20 years in net-zero date.



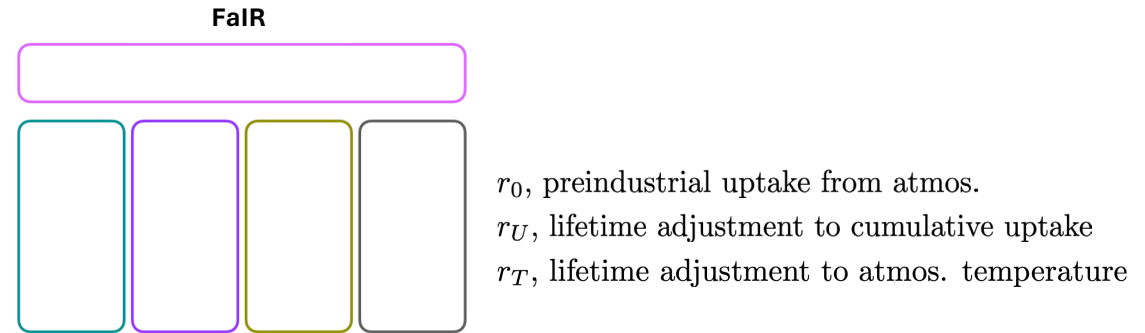
# Replaced timescales in FaIR with a 9-box land and 7-box ocean model



# Sampled biogeochemical feedbacks by perturbing parameters

## FaIR:

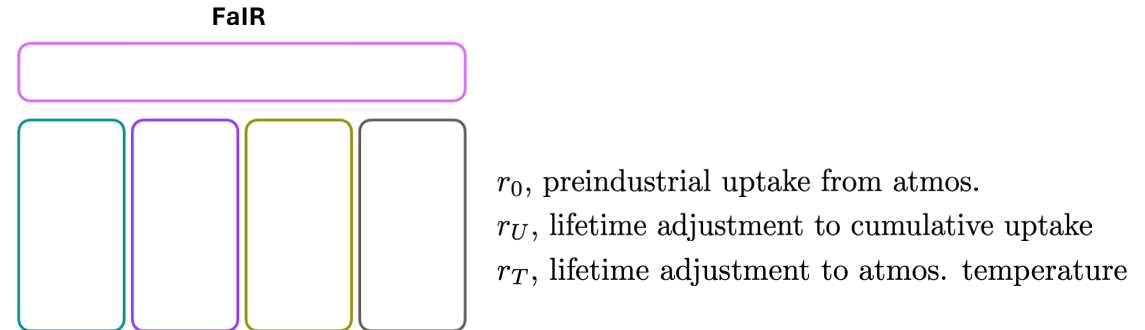
- **3 parameters that influence the carbon cycle**
- 11 parameters → energy balance model
- Concentrations of CH<sub>4</sub> and N<sub>2</sub>O are held constant



# Sampled biogeochemical feedbacks by perturbing parameters

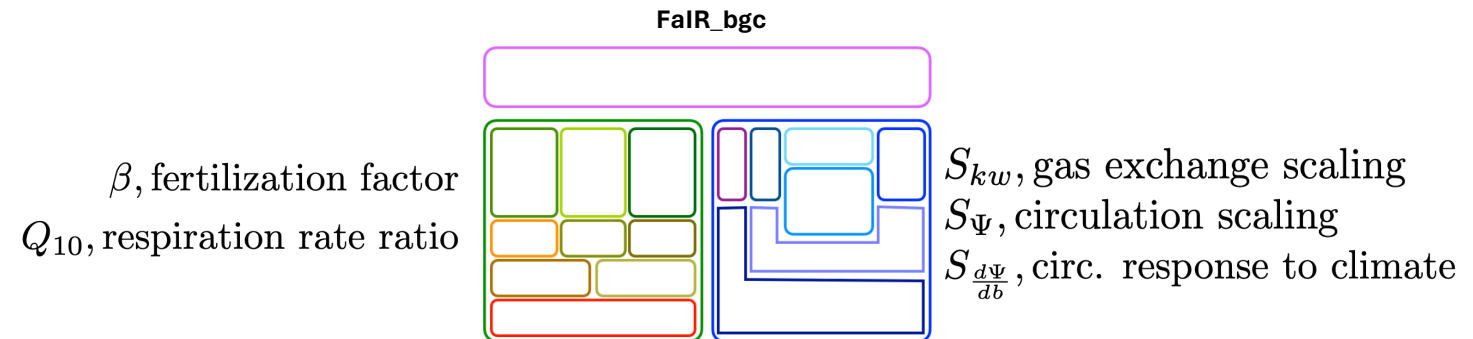
## FaIR:

- **3 parameters that influence the carbon cycle**
- 11 parameters → energy balance model
- Concentrations of CH<sub>4</sub> and N<sub>2</sub>O are held constant

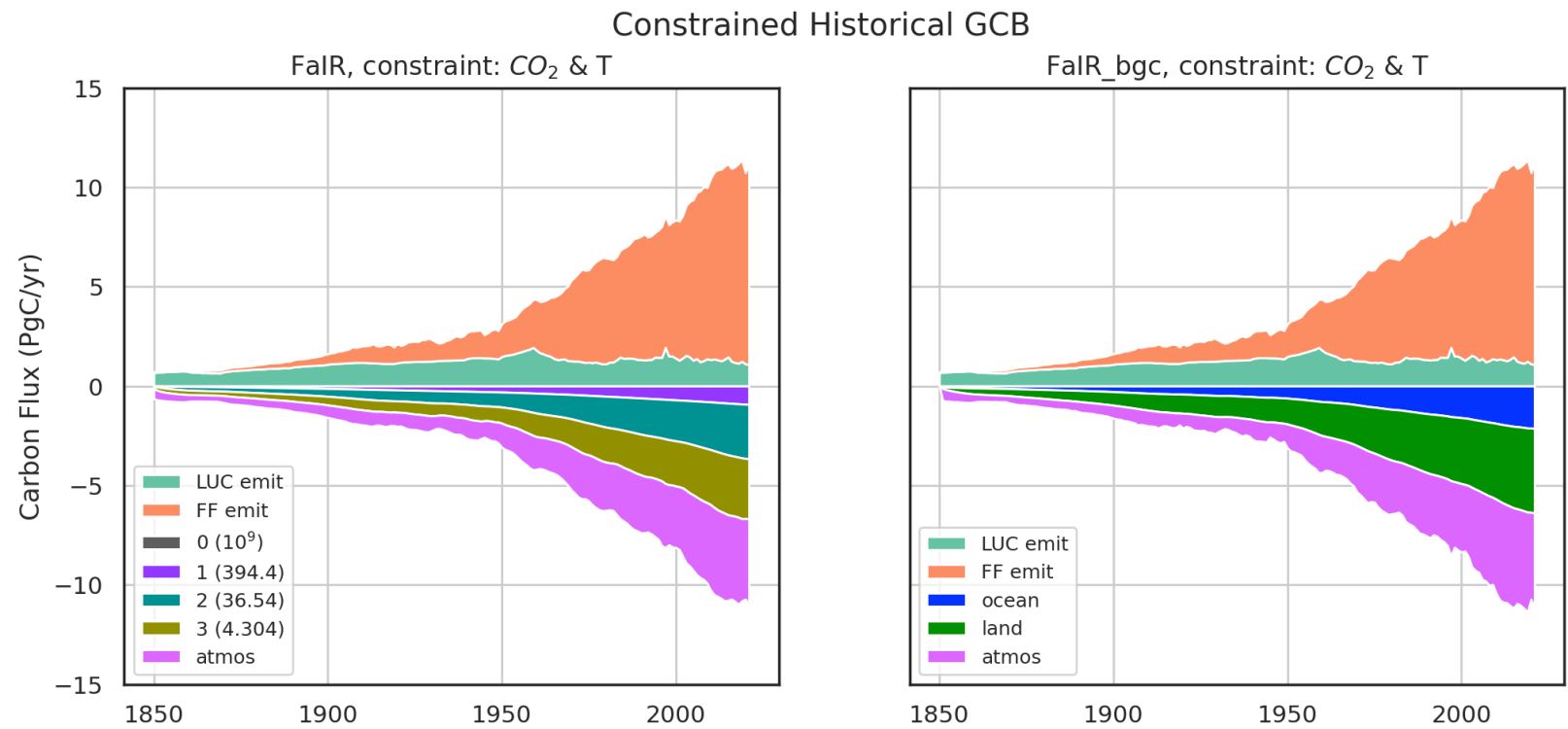


## FaIR\_bgc:

- **5 parameters that influence the carbon cycle**
- 11 parameters → energy balance model
- Concentrations of CH<sub>4</sub> and N<sub>2</sub>O are held constant

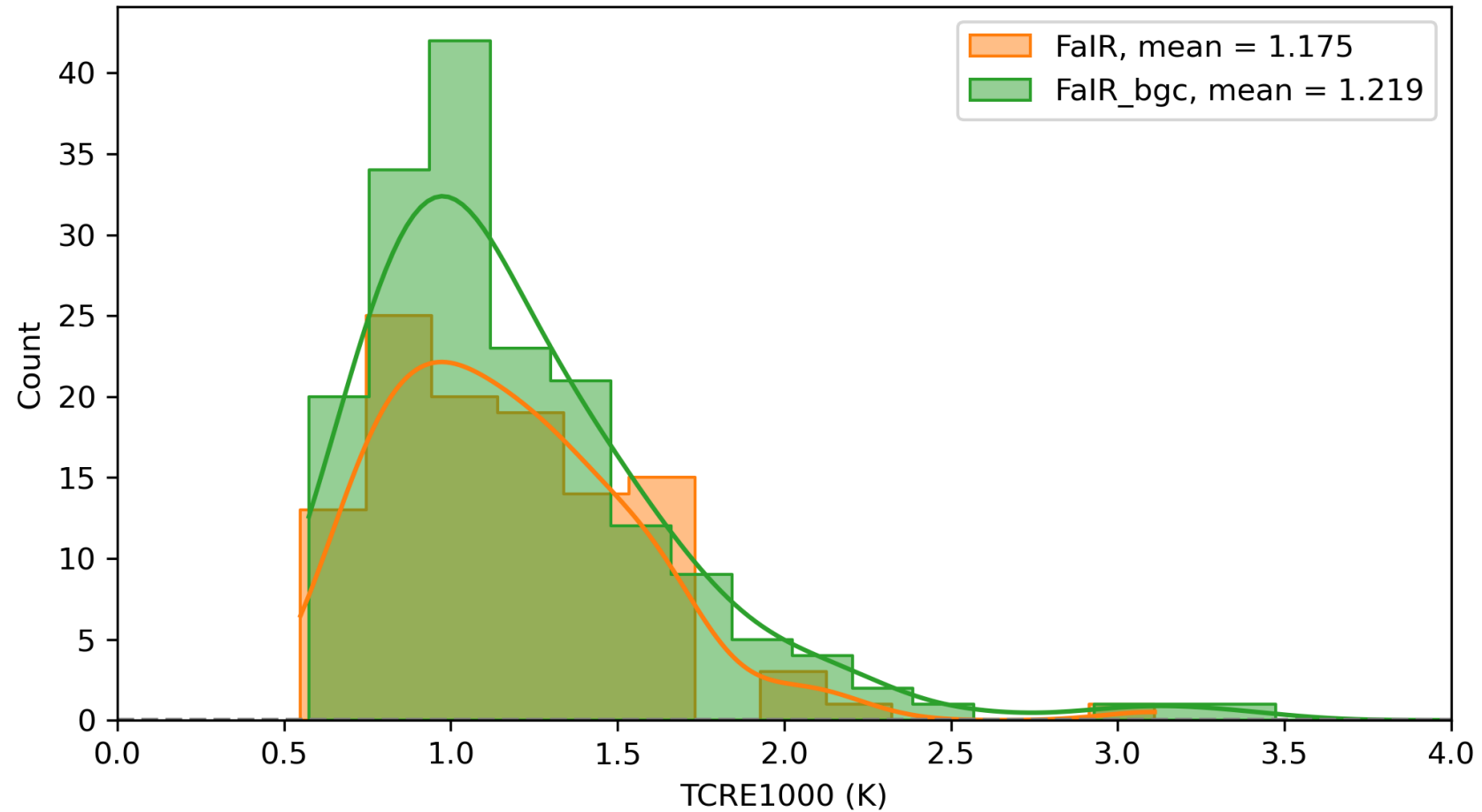


# Compare atmospheric CO<sub>2</sub> and carbon sinks to observations

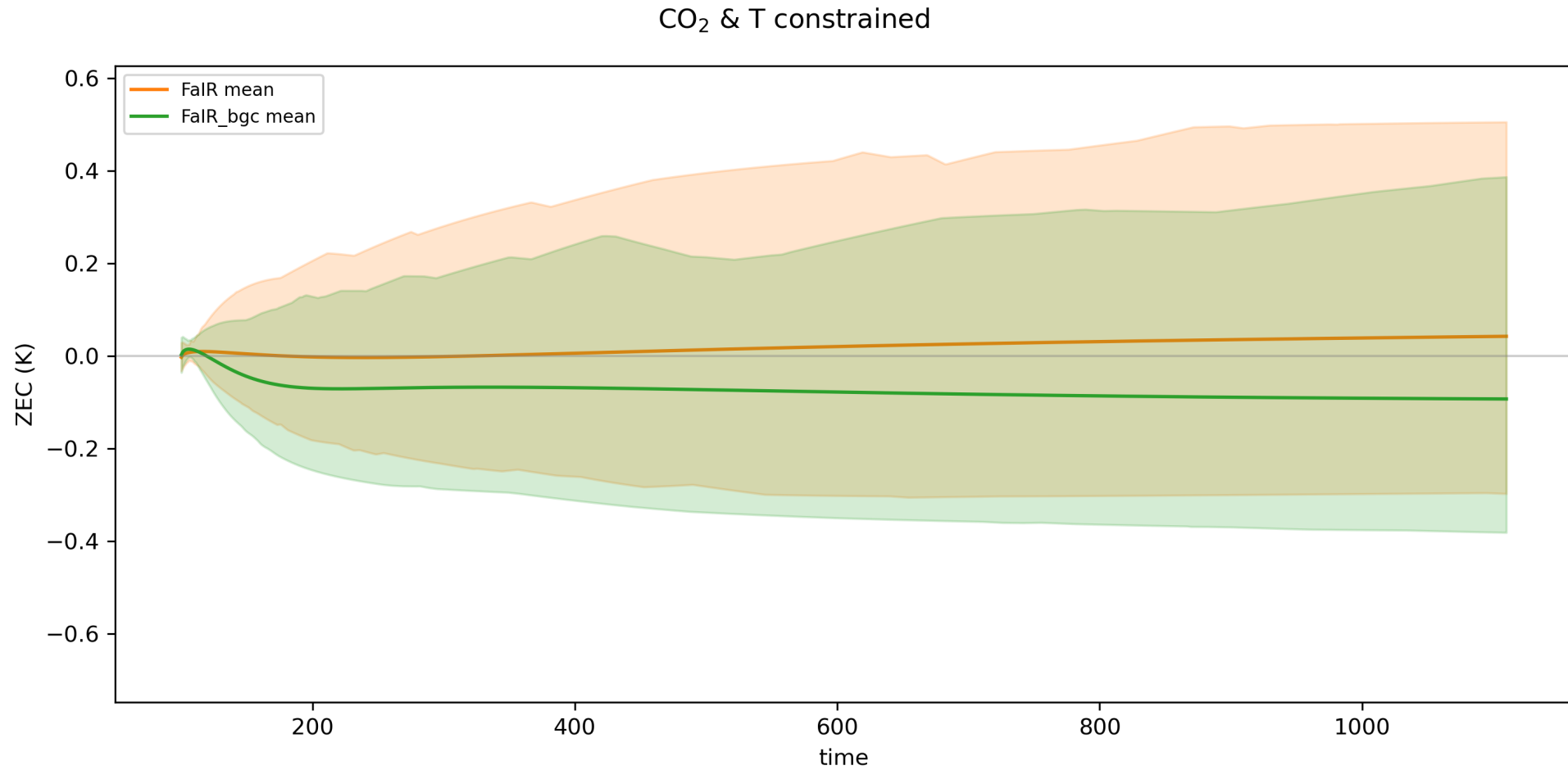




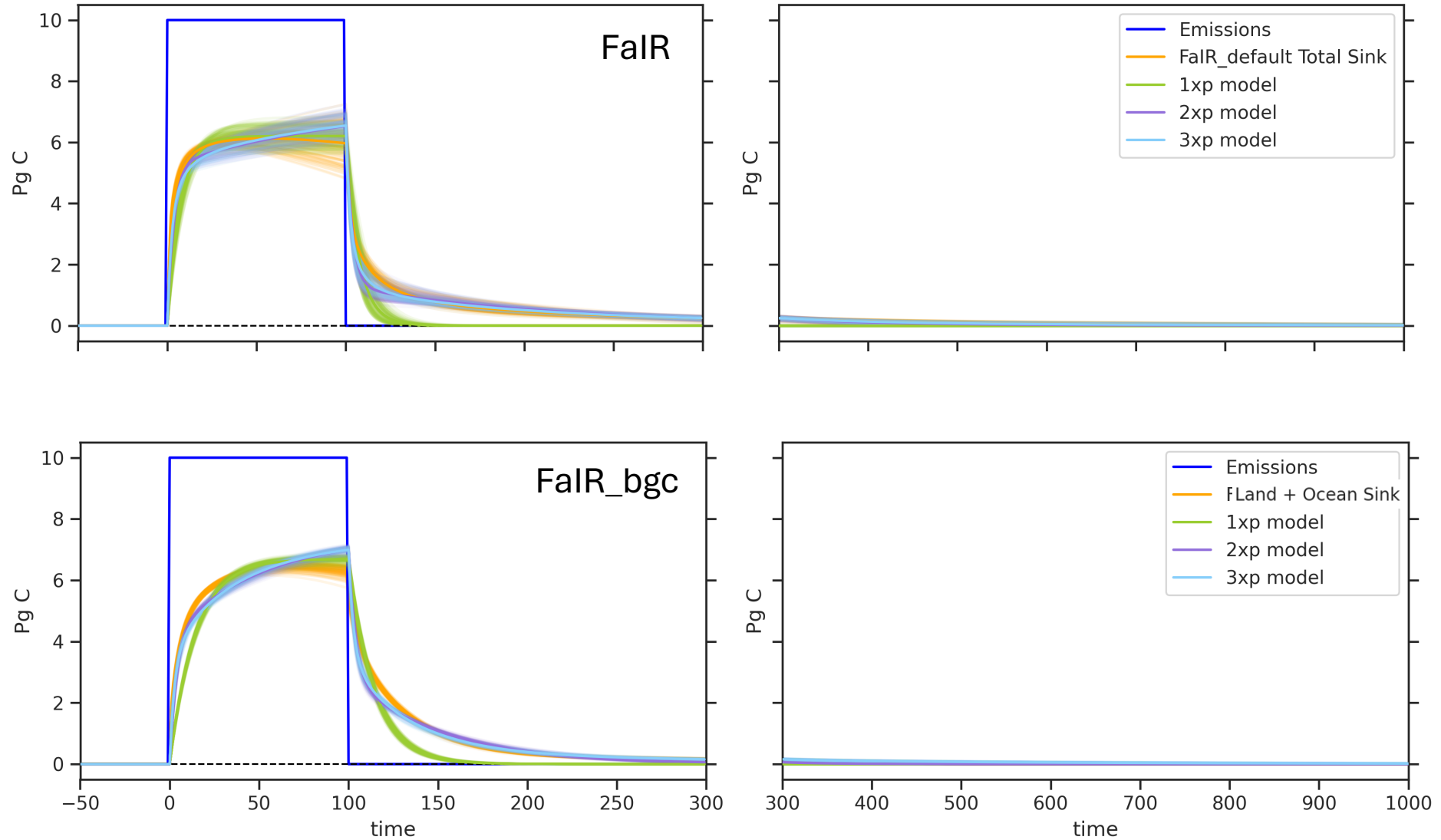
We see minimal changes to TCRE (calculated using the esm-flat10-zec emissions trajectory)



# Small shift in ZEC with different carbon cycle structures

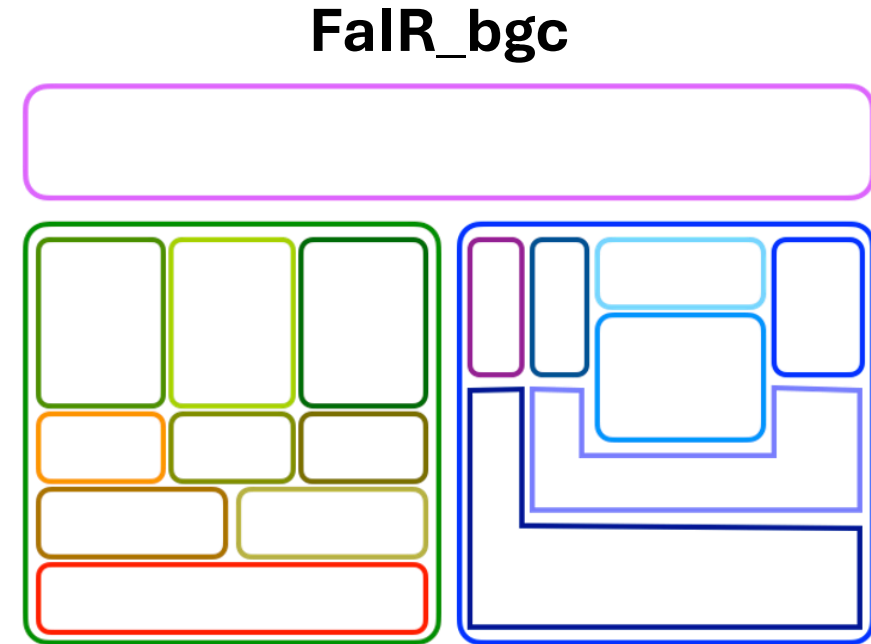
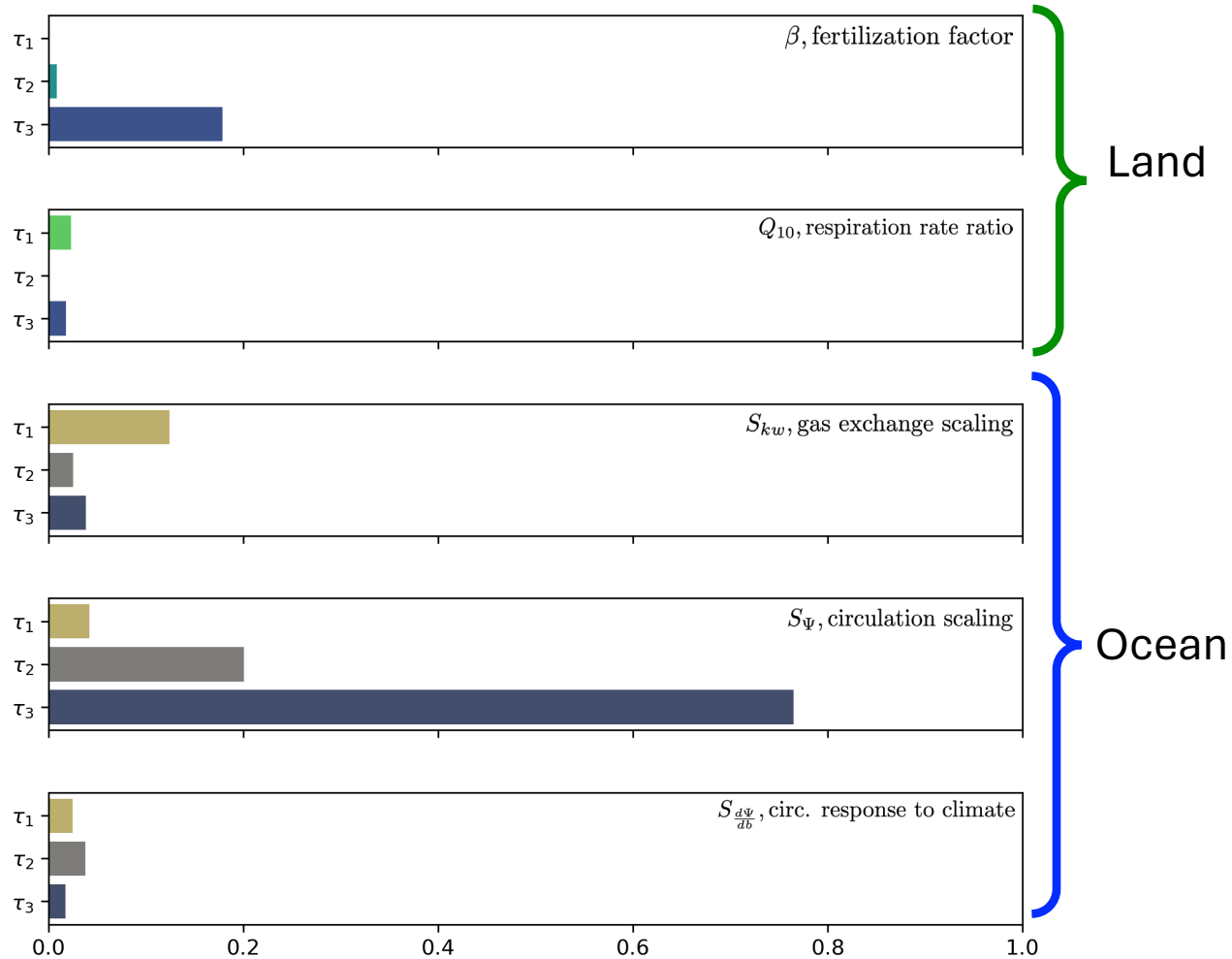


# Use 3xp to diagnose emergent timescales of carbon removal



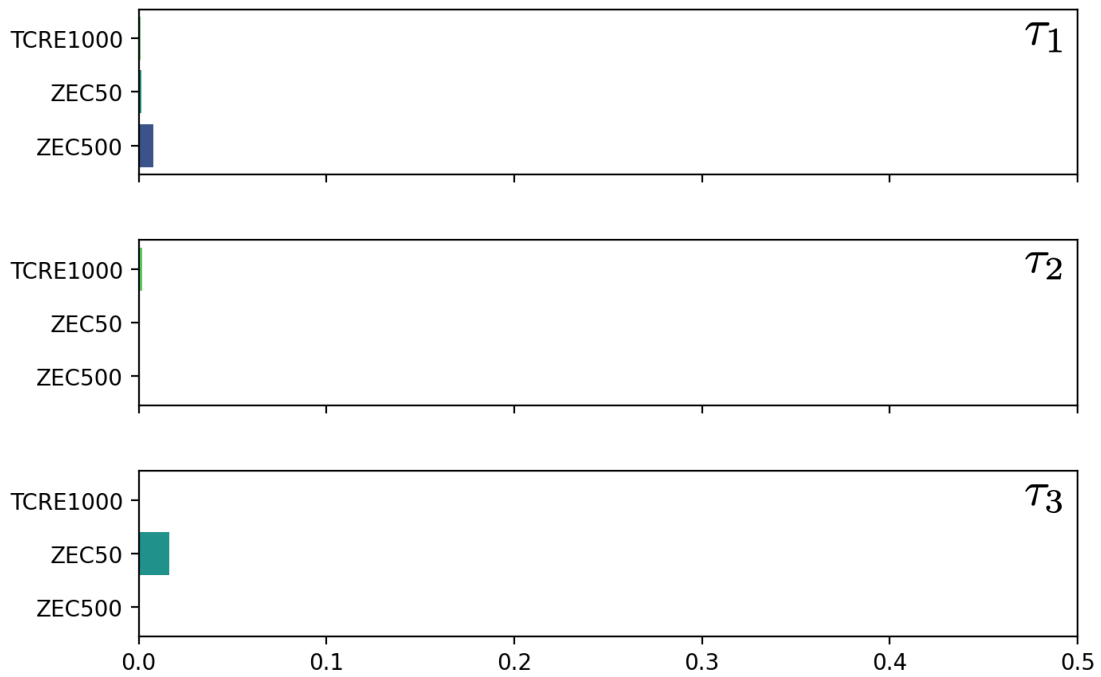
# FaIR\_bgc carbon cycle parameters connect timescales with processes

Variance explained ( $R^2$ ) by FaIR\_bgc parameter



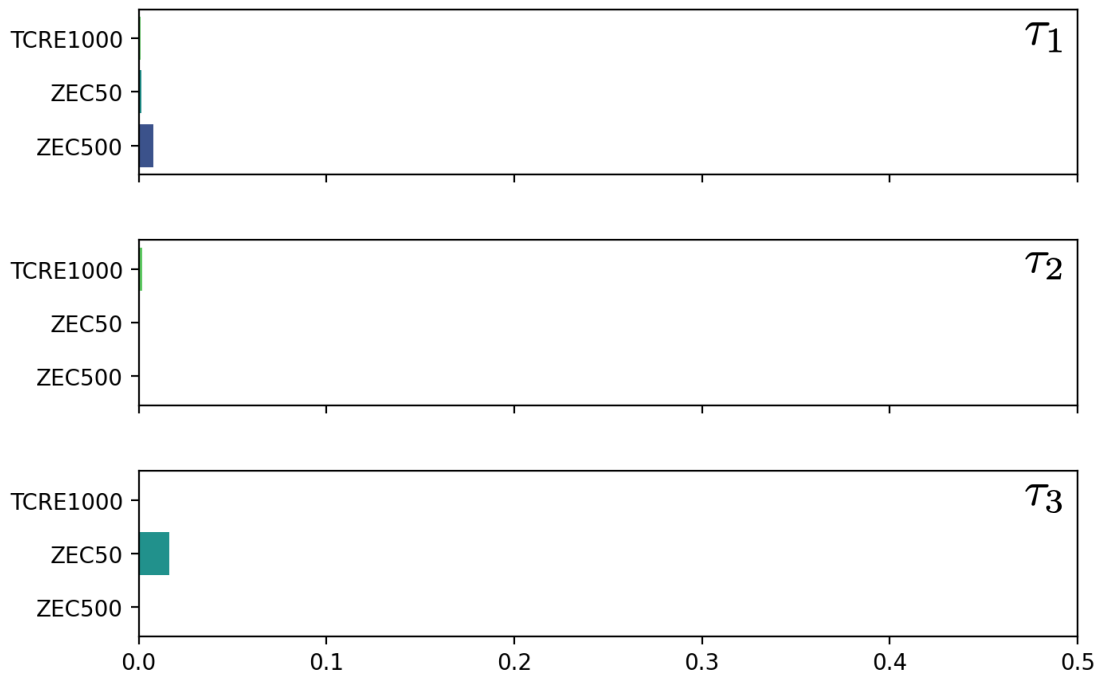
# FaIR\_bgc timescales cannot explain variance in temperature response

Variance explained ( $R^2$ ) by FaIR\_bgc timescales



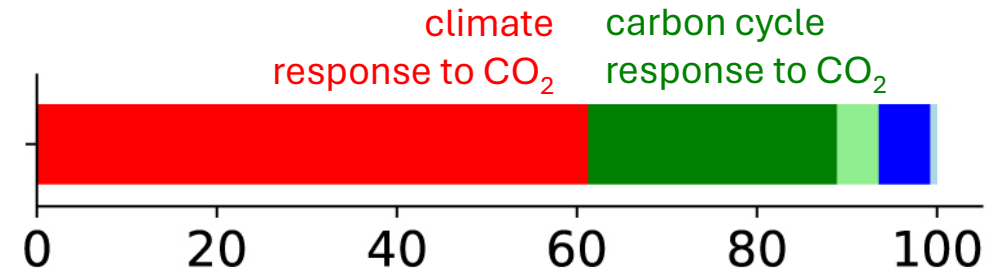
# FaIR\_bgc timescales cannot explain variance in temperature response

Variance explained ( $R^2$ ) by FaIR\_bgc timescales



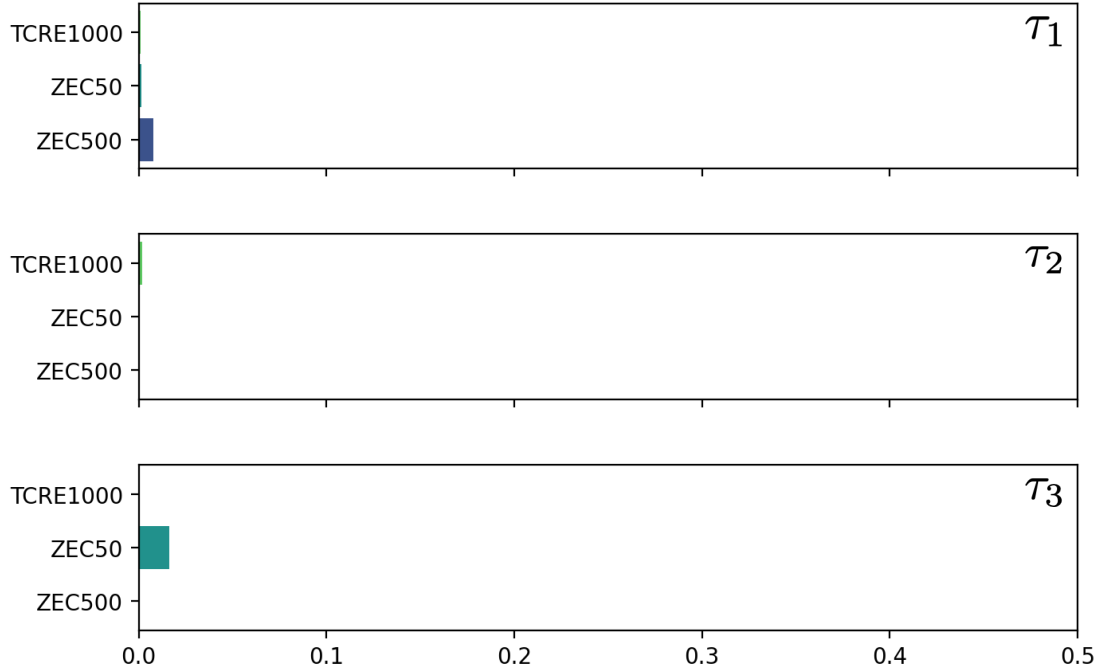
*For comparison:*

Contributions to TCRE in C4MIP ESMs:

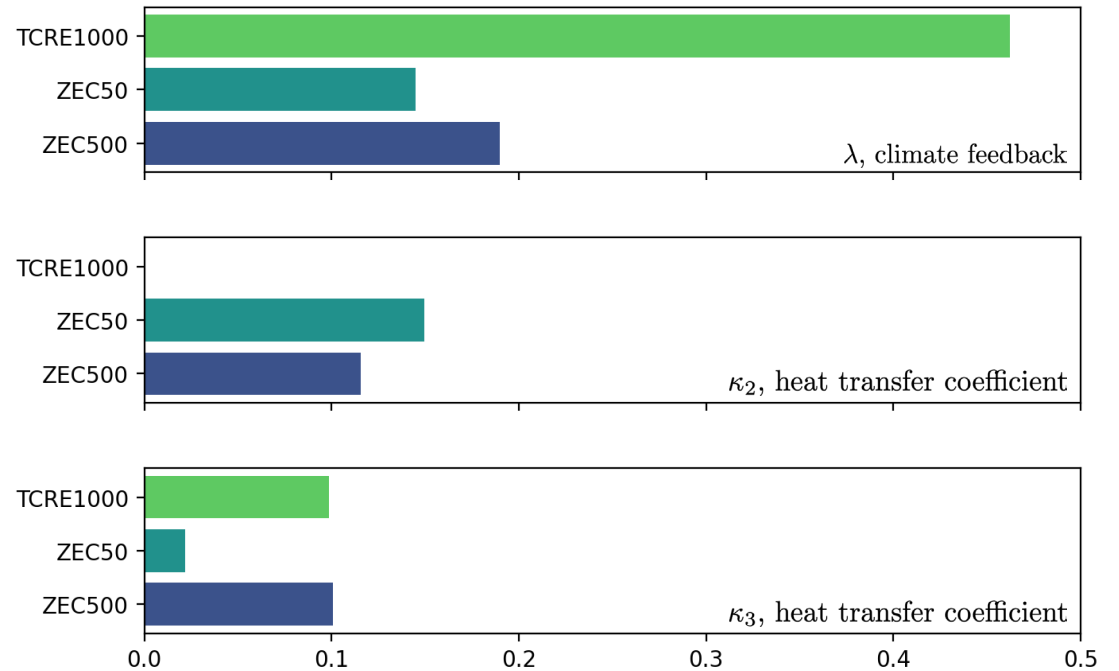


# FaIR\_bgc energy balance model parameters explain variance in temperature response

Variance explained ( $R^2$ ) by FaIR\_bgc timescales

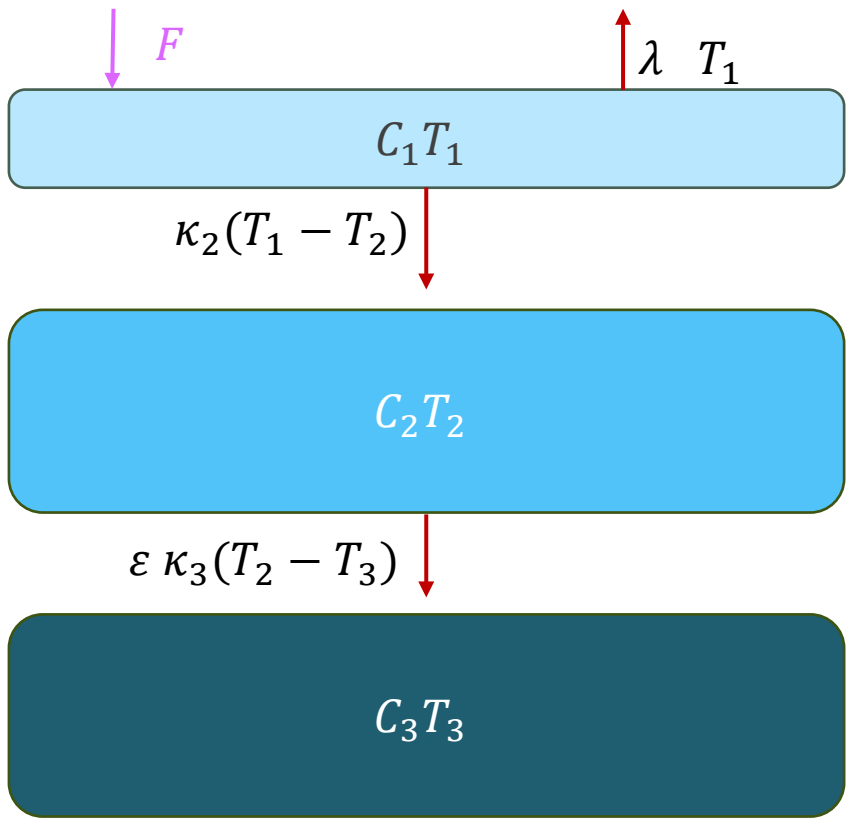


Variance explained ( $R^2$ ) by EBM parameters

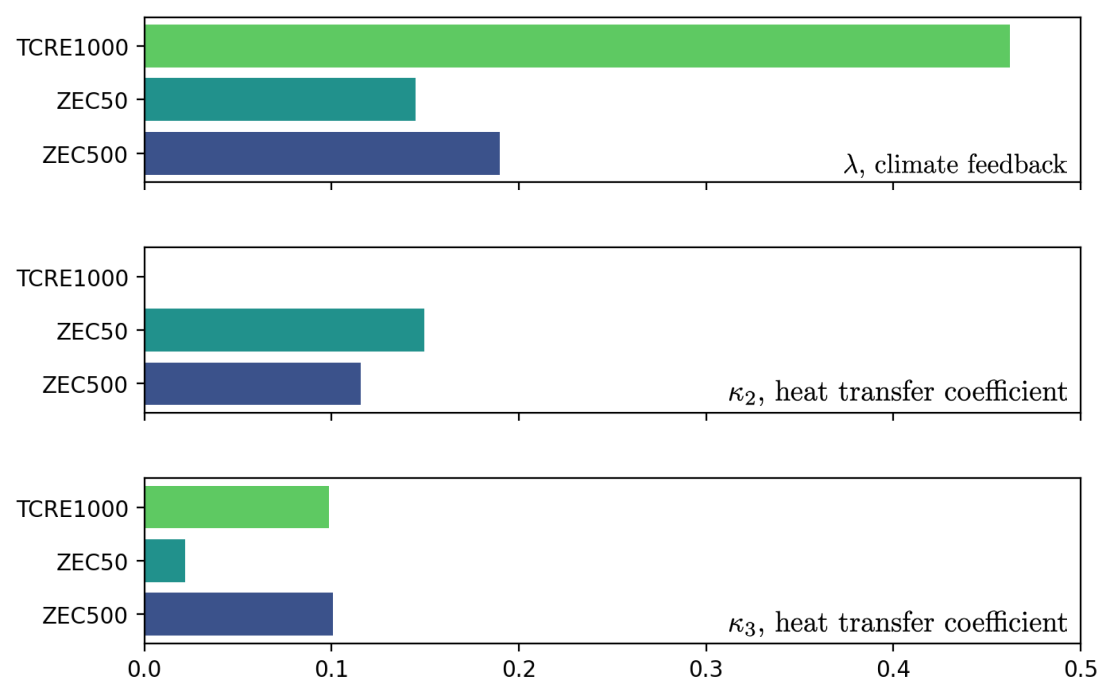




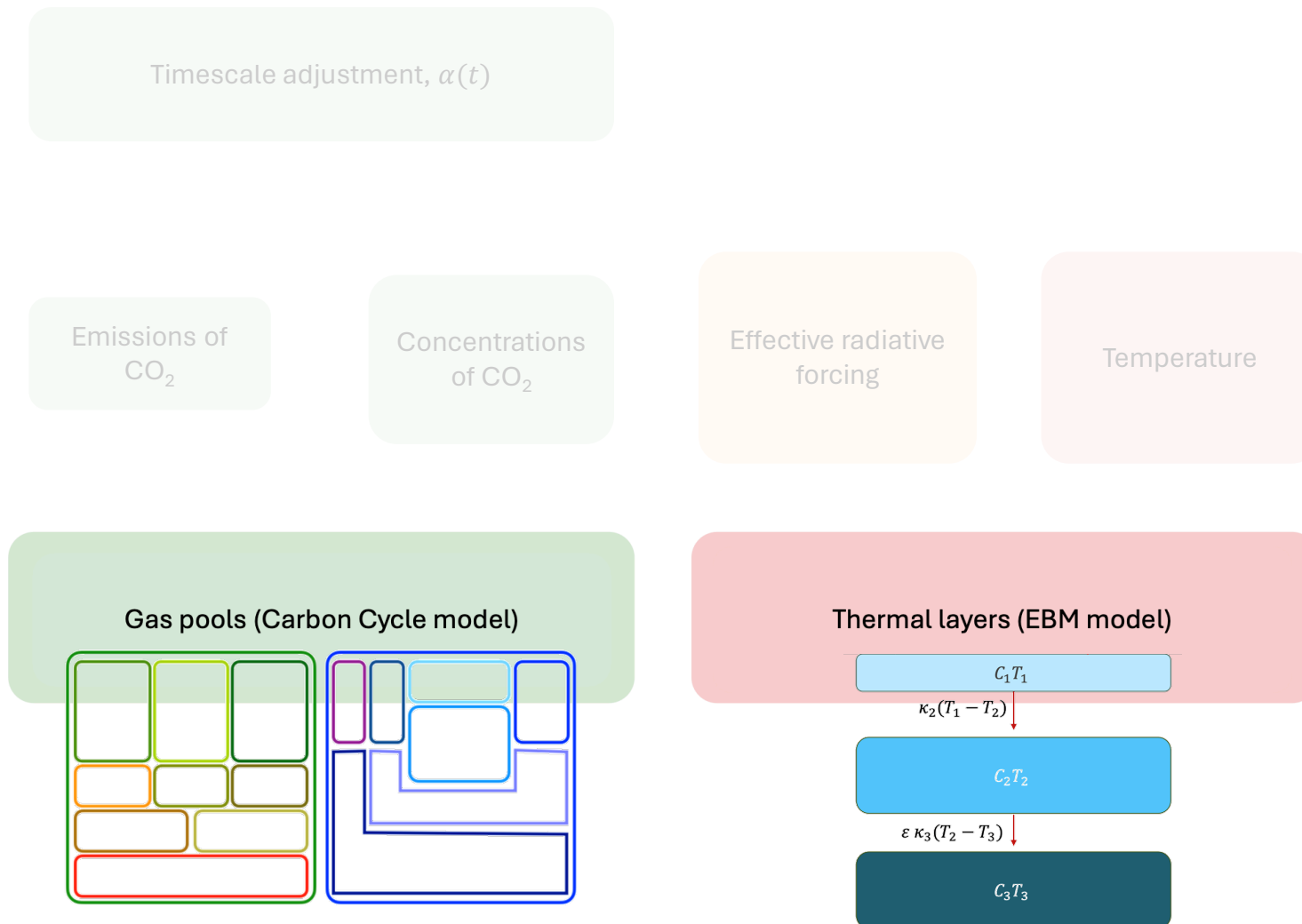
# EBM parameters the control energy fluxes and evolution of temperature



Variance explained ( $R^2$ ) by EBM parameters



# Why do carbon sink timescales influence ZEC so little?



# Why do carbon sink timescales influence ZEC so little?

Timescale adjustment,  $\alpha(t)$

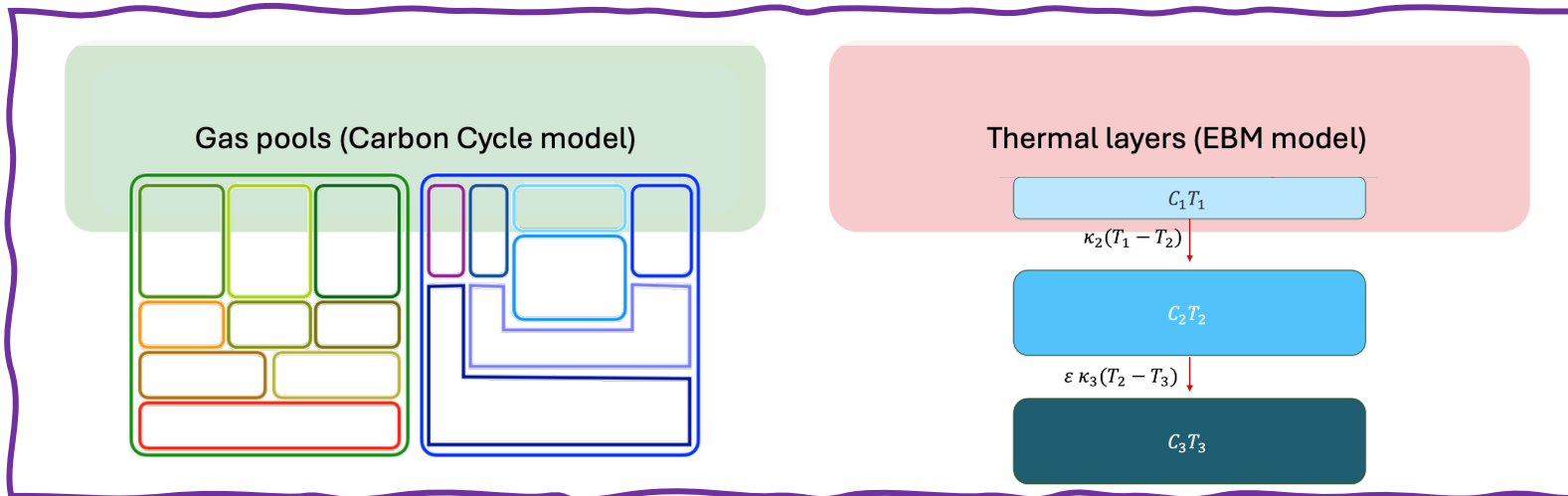
Emissions of  $\text{CO}_2$

Concentrations of  $\text{CO}_2$

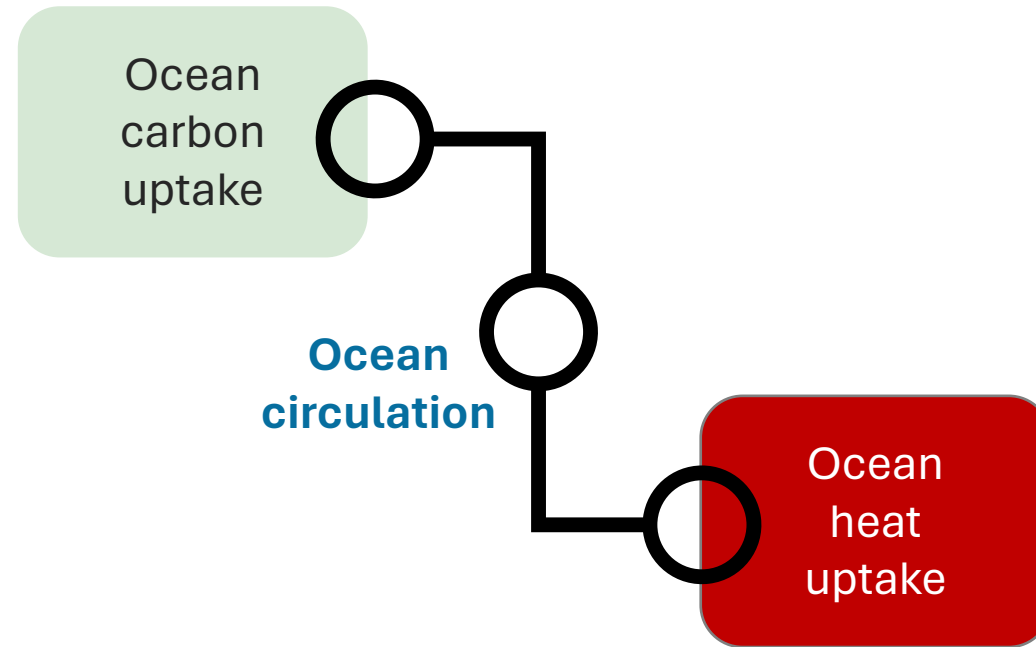
Effective radiative forcing

Temperature

*No structural link between processes that govern carbon and energy fluxes!*



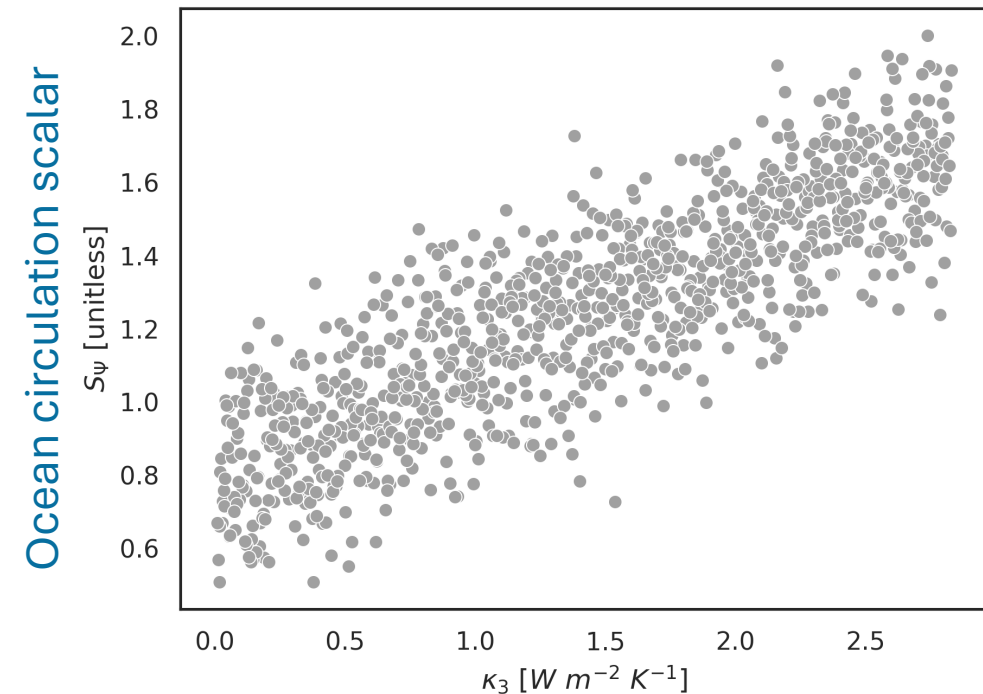
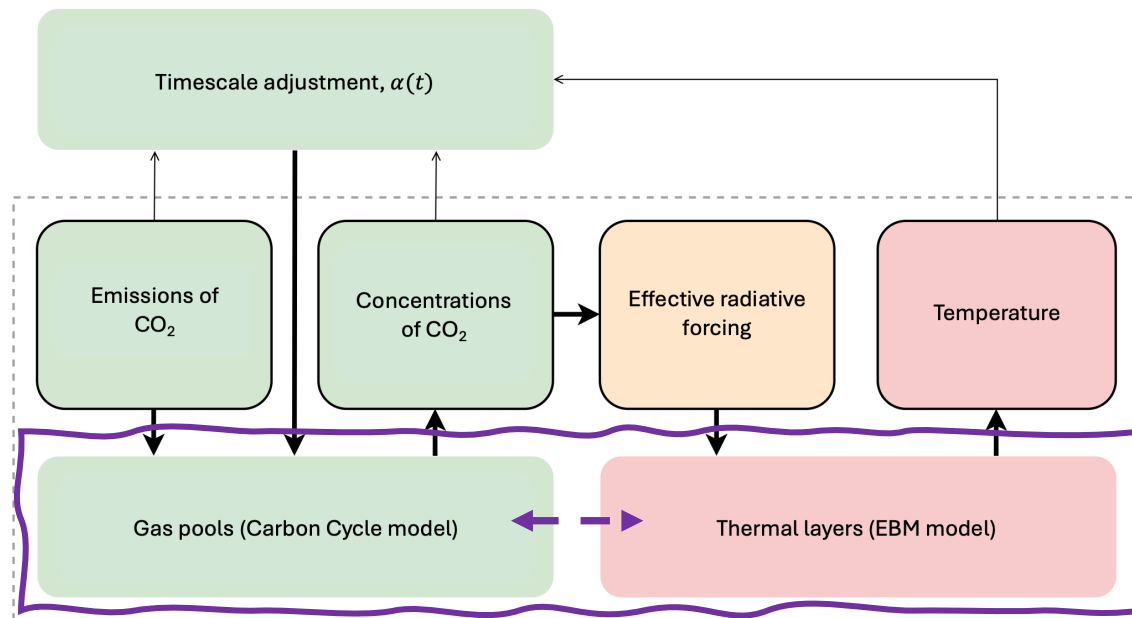
# Ocean carbon uptake and heat uptake are both impacted by circulation



# Hypothesis: if we correlate ocean circulation with ocean heat uptake, the influence of carbon uptake on ZEC may change

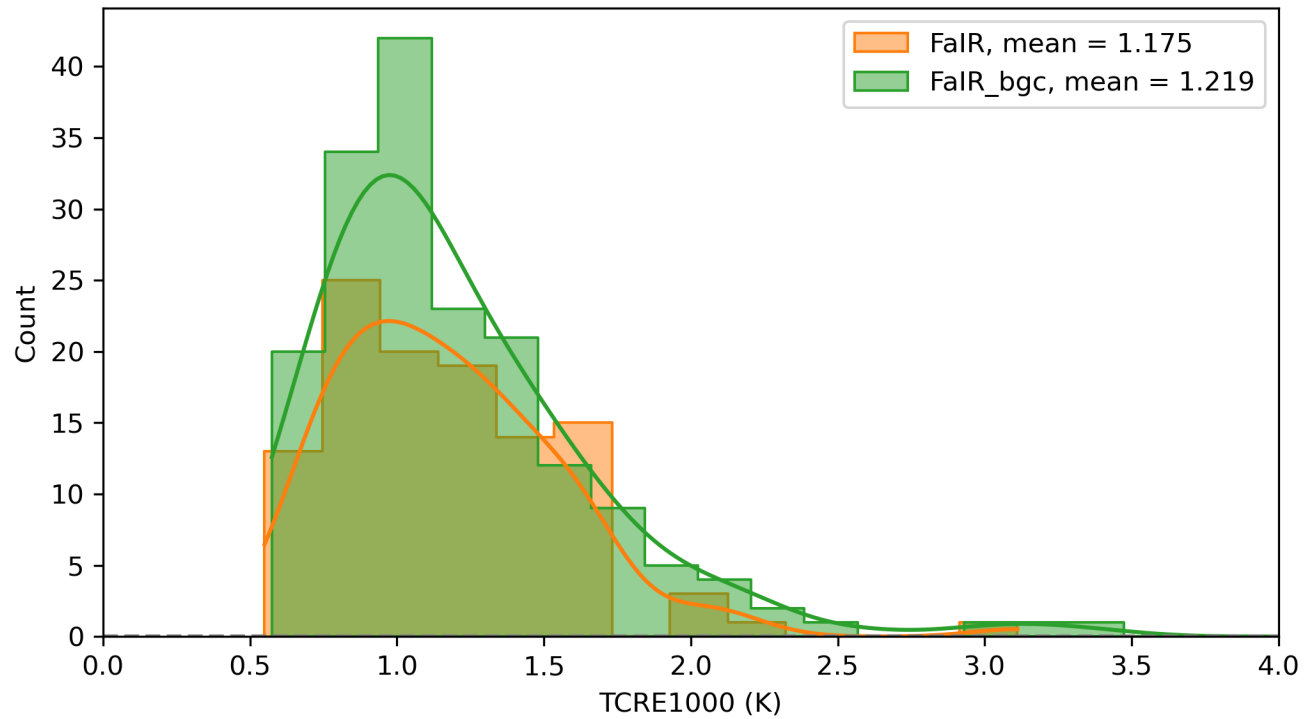
Does the influence of carbon cycle timescales increase when thermal boxes are connected through correlation to carbon cycle boxes?

- Created a new Latin Hypercube ensemble with **correlated parameter values** for  $S_\psi$  and heat transport,  $\kappa_3$ .

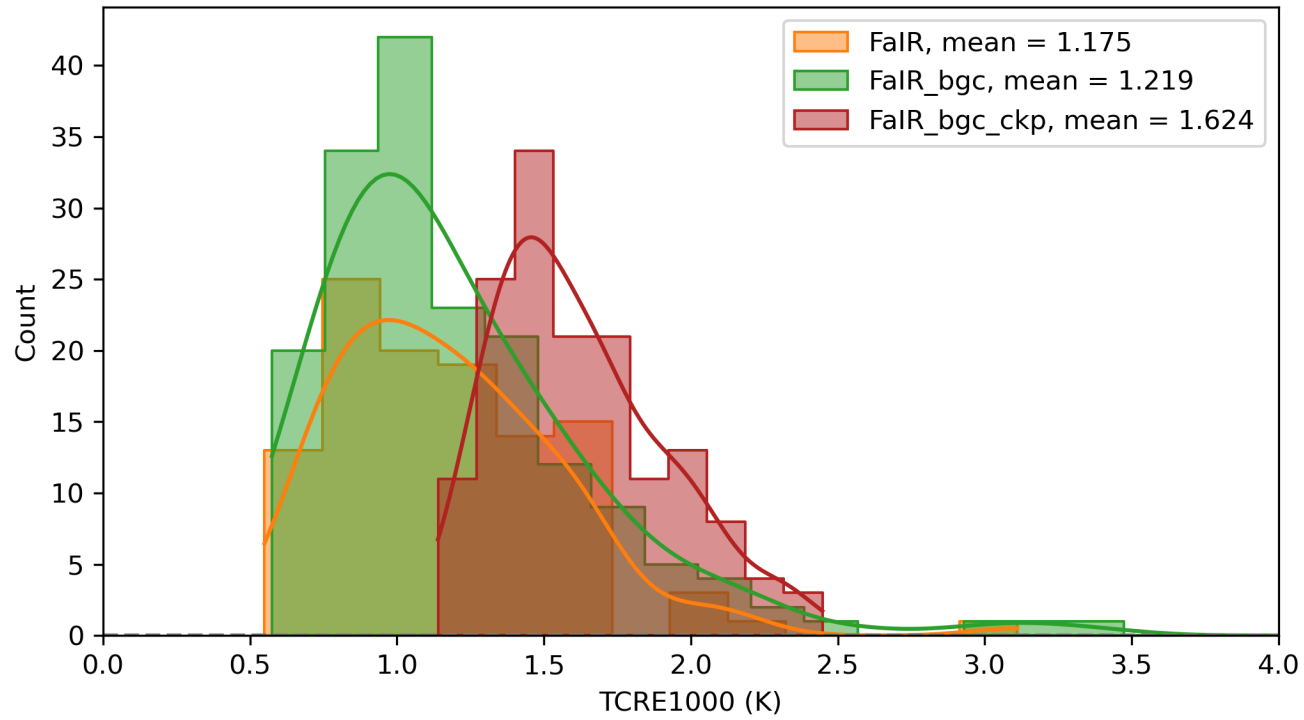


Heat transport to deep ocean

# Can adding correlation between circulation rate and thermal transfer shift TCRE?

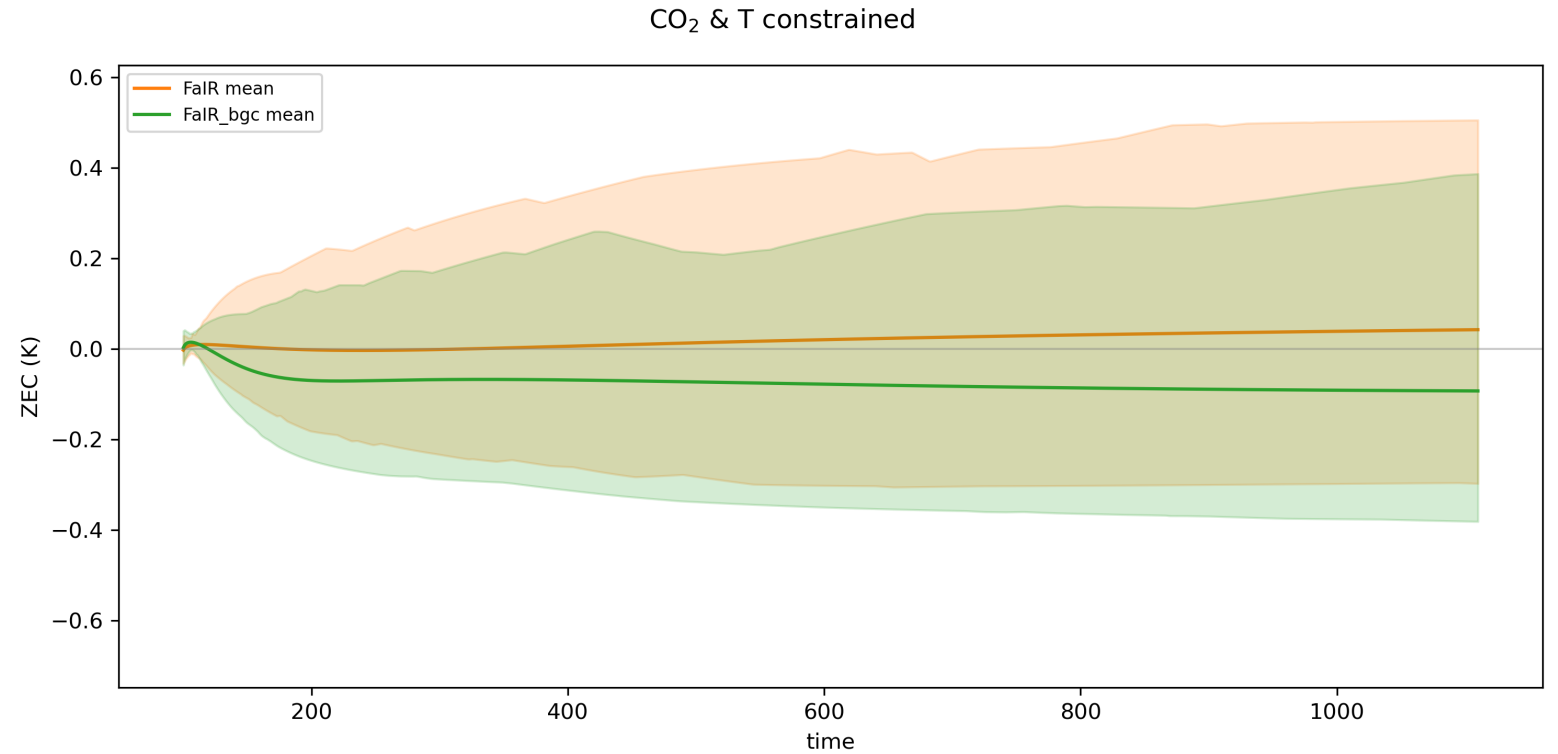
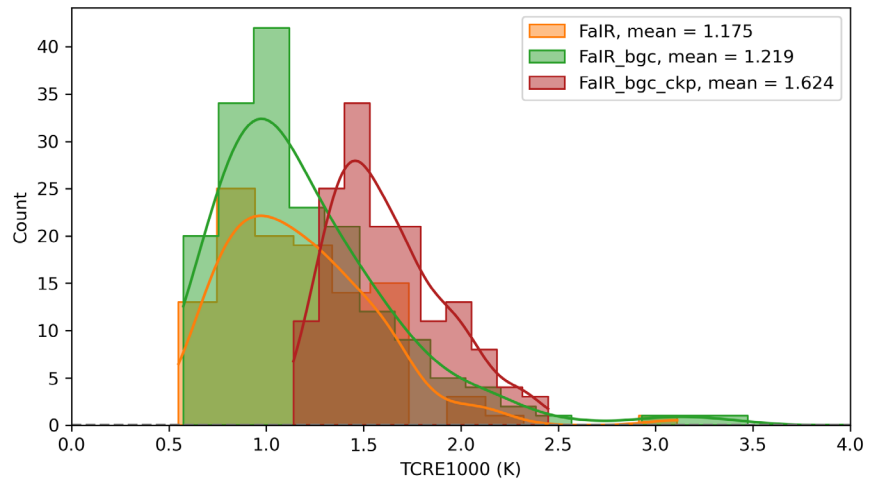


# TCRE gets stronger!

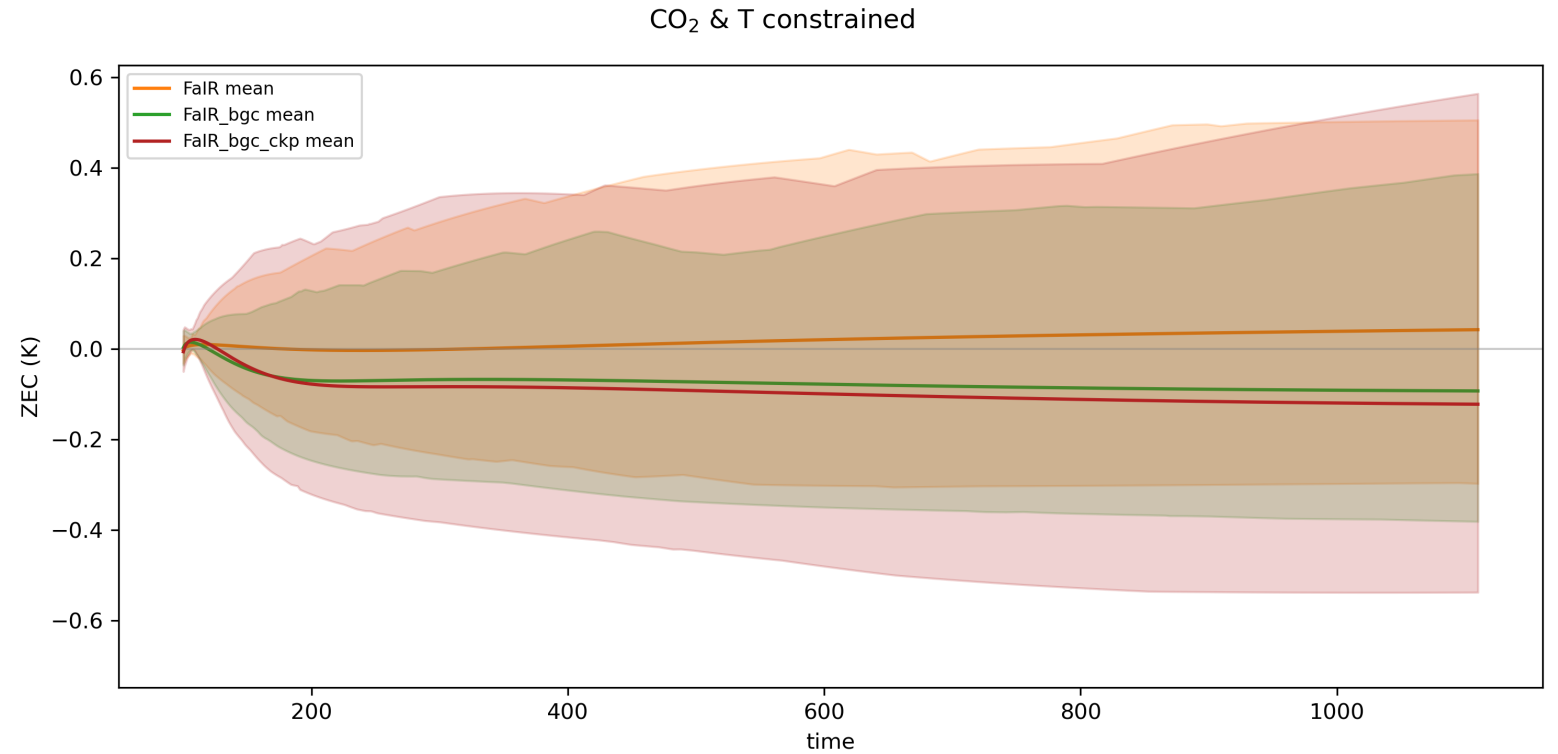
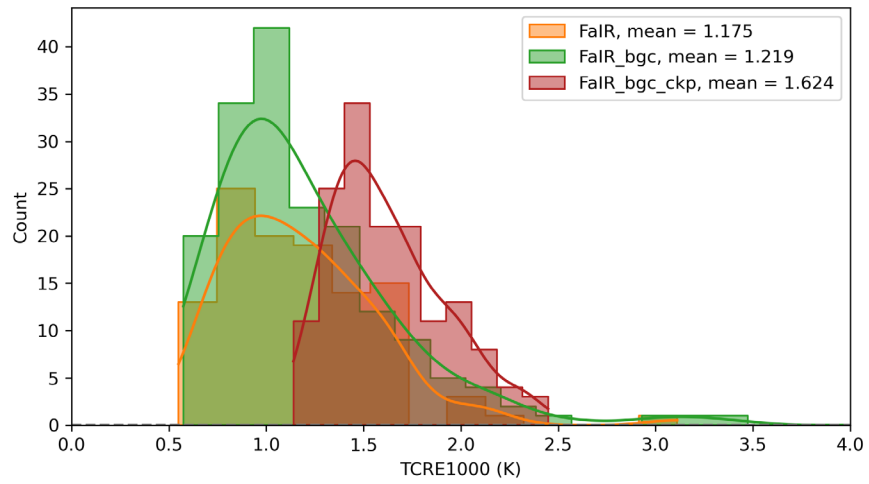




# Can adding correlation between circulation rate and thermal transfer shift ZEC?

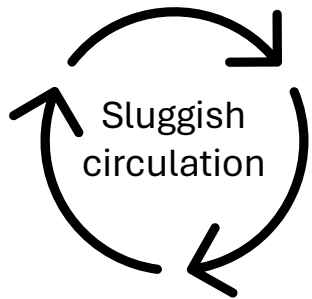


# ZEC gets weakly more negative



# Uncorrelated: circulation affects CO<sub>2</sub> and climate through atmos. CO<sub>2</sub>

Historical  
emission of  
CO<sub>2</sub> →

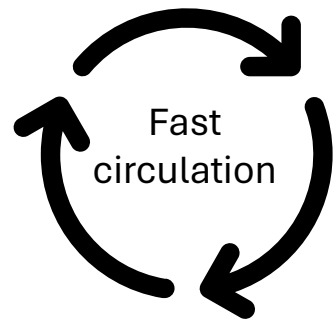


slow ocean  
sink



higher atmos.  
CO<sub>2</sub>

Stronger  
atmos.  
warming



rapid ocean  
sink



lower atmos.  
CO<sub>2</sub>

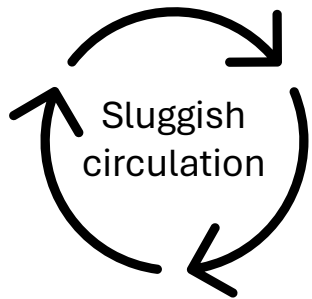
Weaker  
atmos.  
warming

In constraining to  
observed CO<sub>2</sub> and ΔT

We select other  
parameters that  
compensate for these  
circulation-driven  
outcomes.

# Uncorrelated: circulation affects CO<sub>2</sub> and climate through atmos. CO<sub>2</sub>

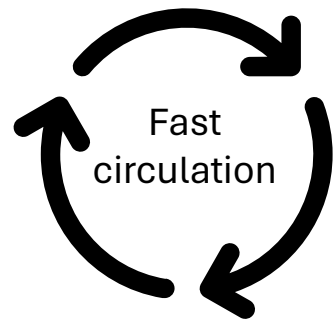
Historical emission of CO<sub>2</sub> →



slow ocean sink

higher atmos. CO<sub>2</sub>

Stronger atmos. warming



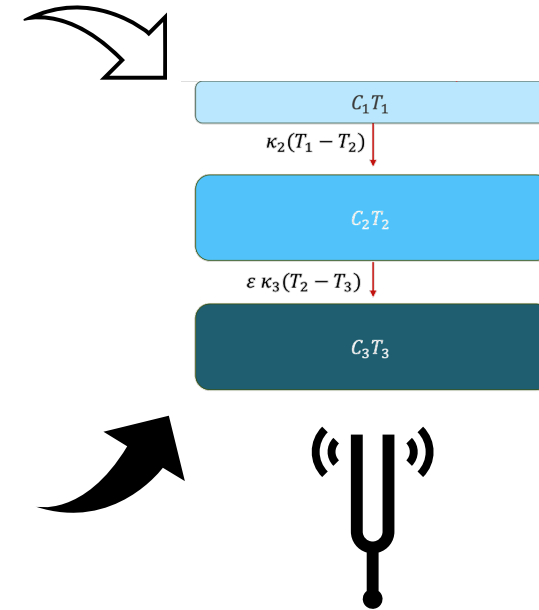
rapid ocean sink

lower atmos. CO<sub>2</sub>

Weaker atmos. warming

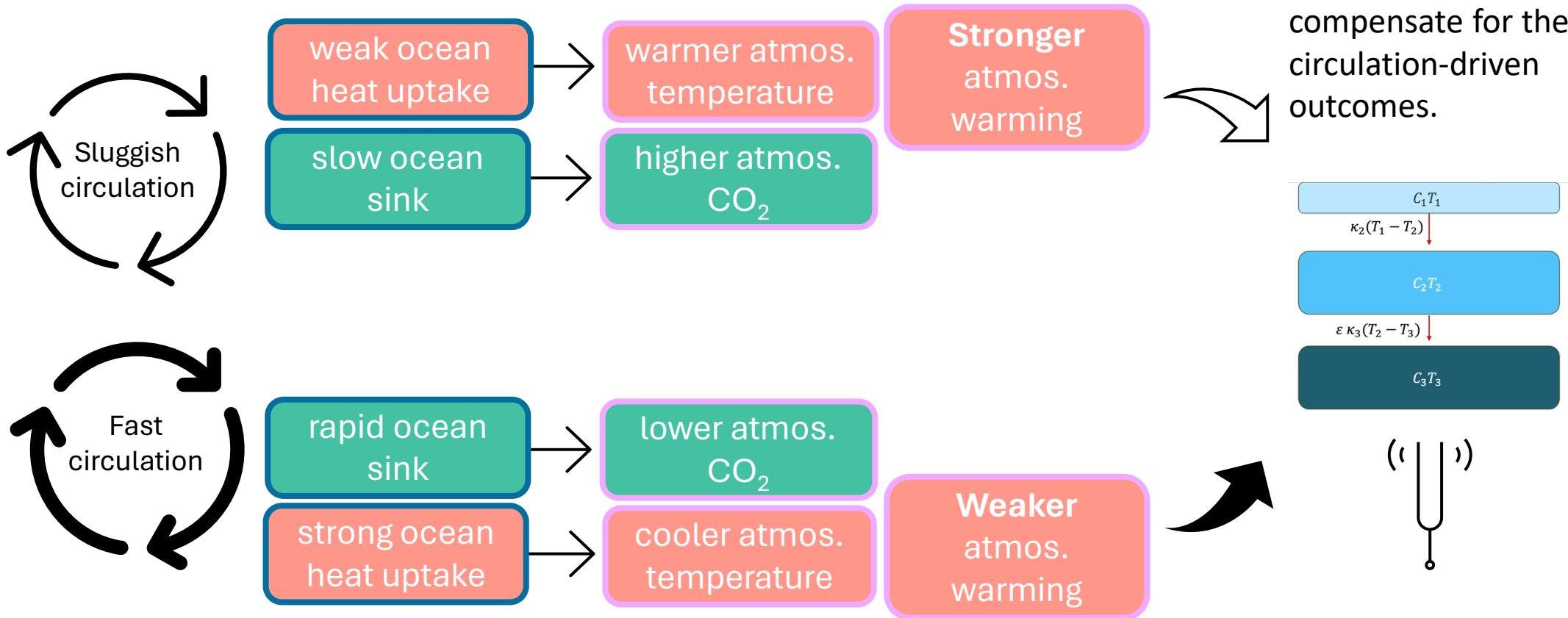
In constraining to observed CO<sub>2</sub> and ΔT

We select other parameters that compensate for these circulation-driven outcomes.



# Correlated: circulation affects climate through atmospheric CO<sub>2</sub> and through heat uptake

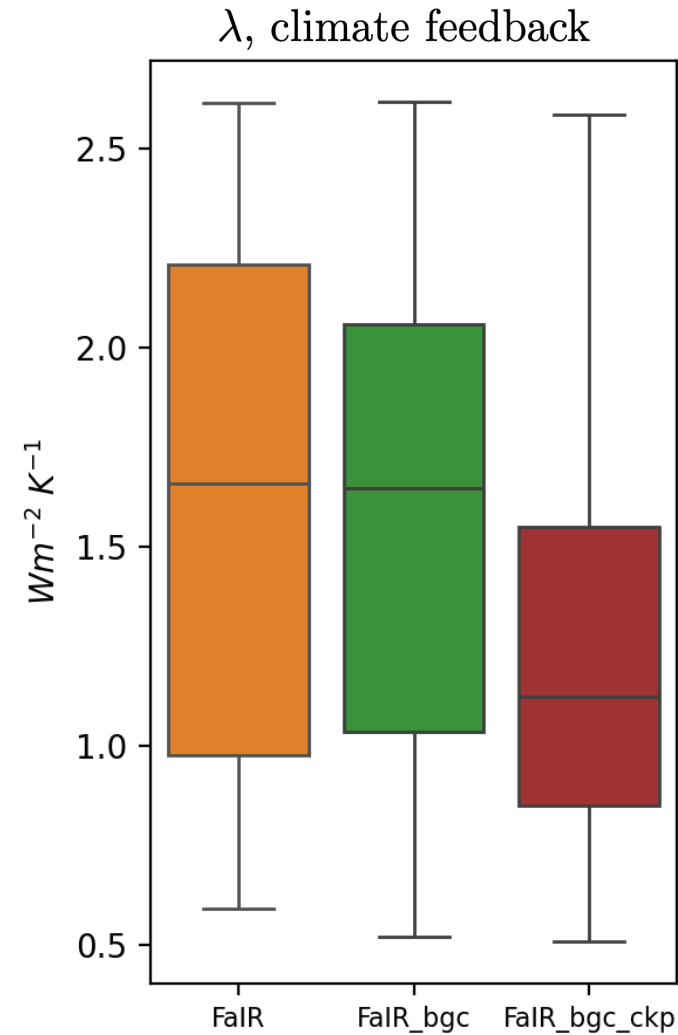
Historical emission of CO<sub>2</sub> →



In constraining to observed CO<sub>2</sub> and  $\Delta T$

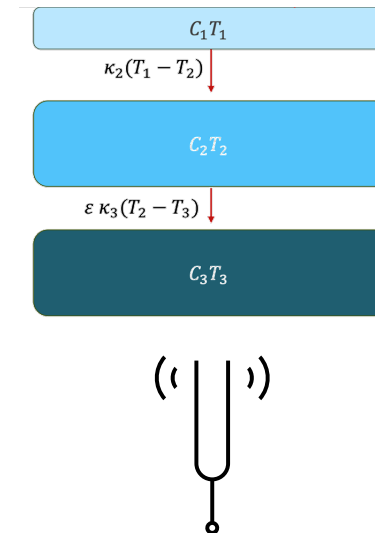
We select other parameters that compensate for these circulation-driven outcomes.

# Compensating shift in EBM parameters when constrained to CO<sub>2</sub> & T

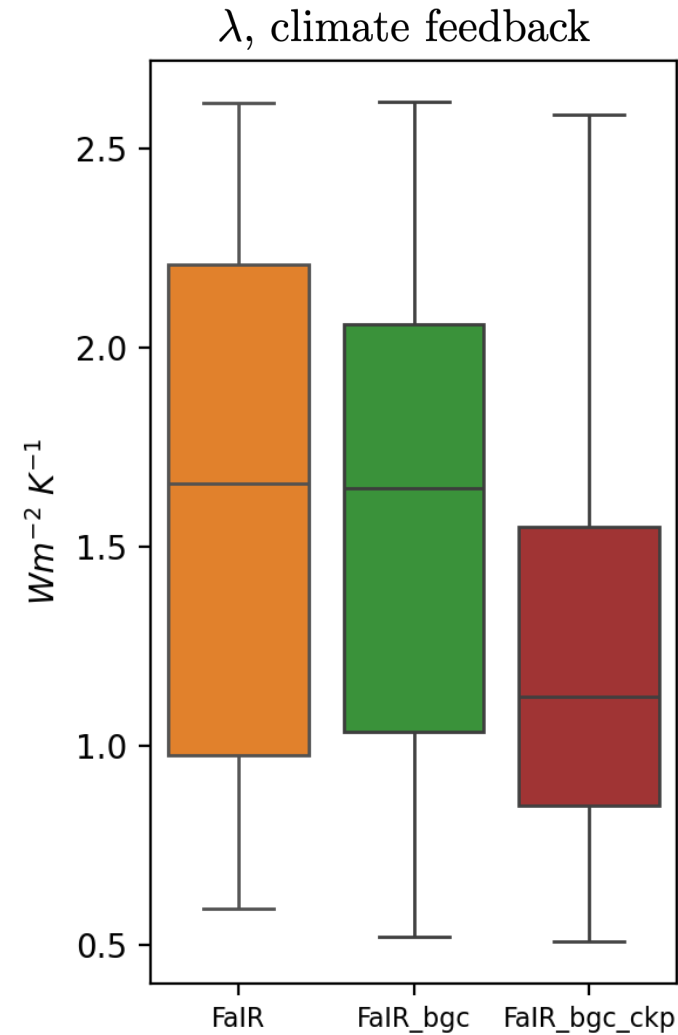
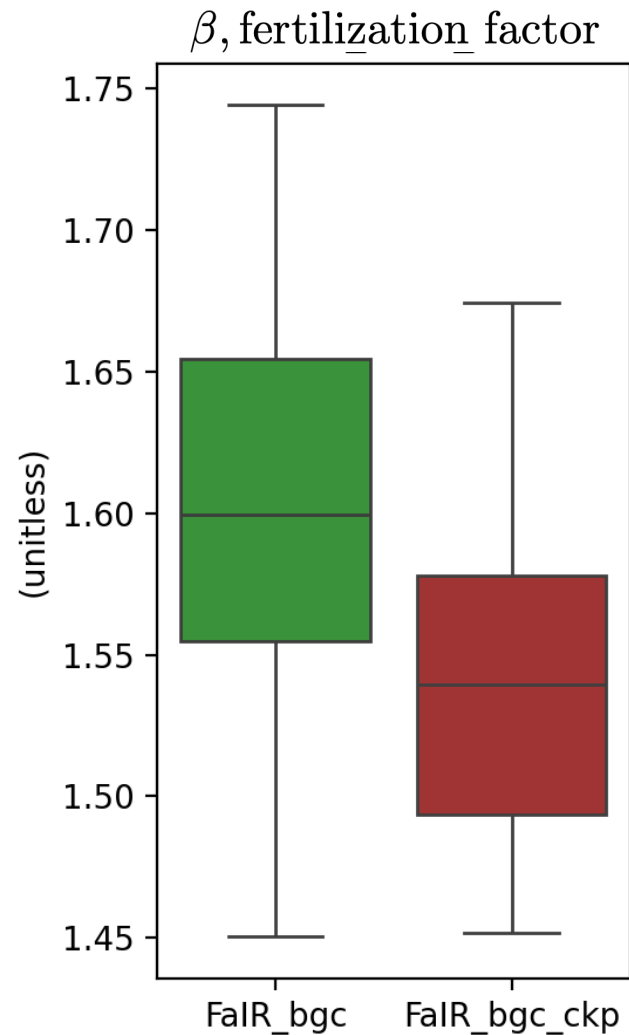


In constraining to observed CO<sub>2</sub> and  $\Delta T$

We select other parameters that compensate for these circulation-driven outcomes.

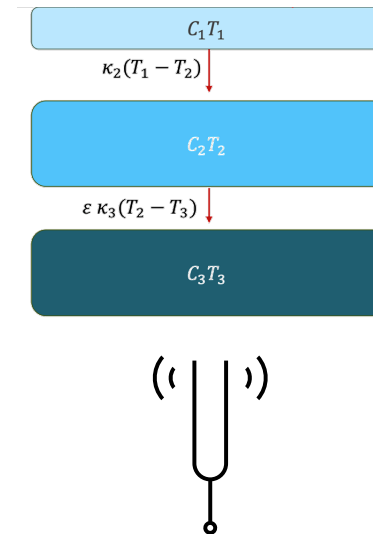


# Compensating shift in CC parameters when constrained to CO<sub>2</sub> & T

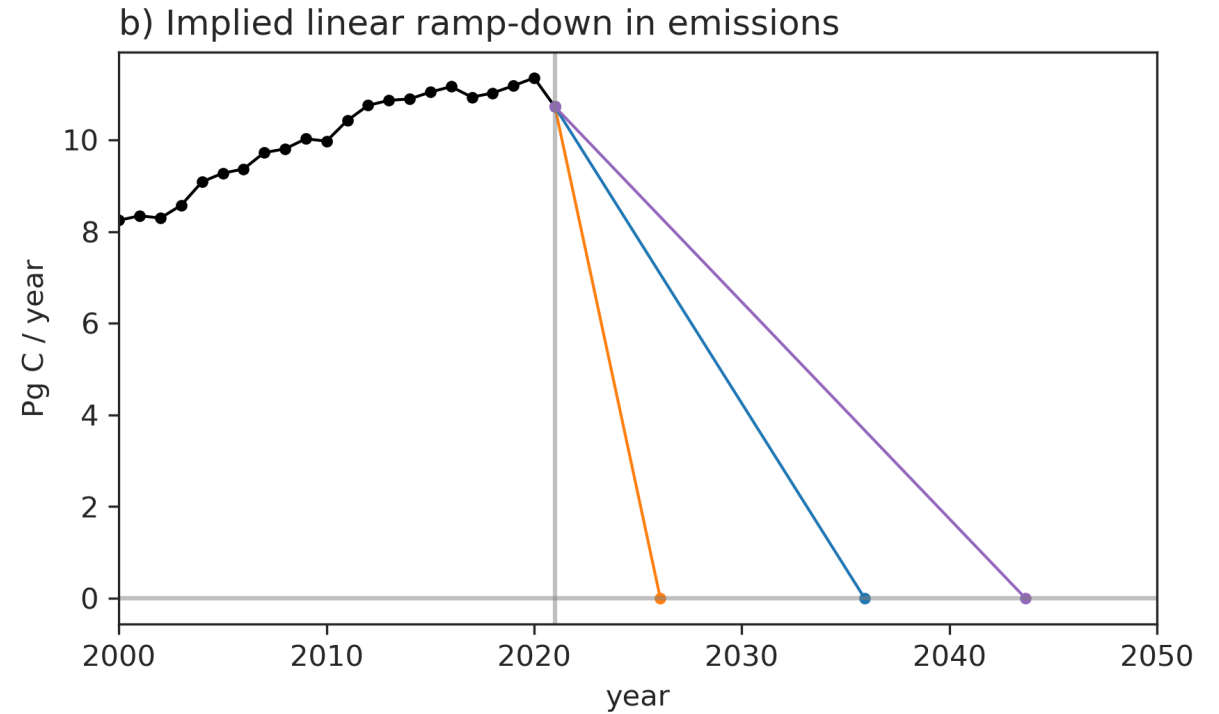
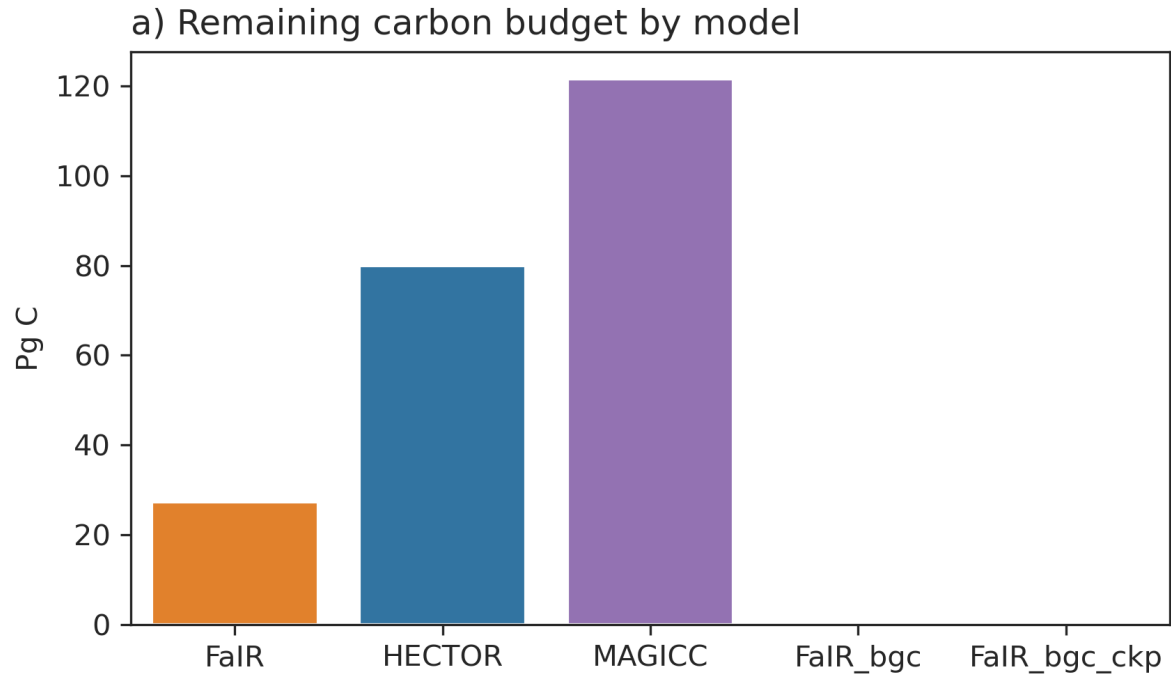


In constraining to observed CO<sub>2</sub> and  $\Delta T$

We select other parameters that compensate for these circulation-driven outcomes.

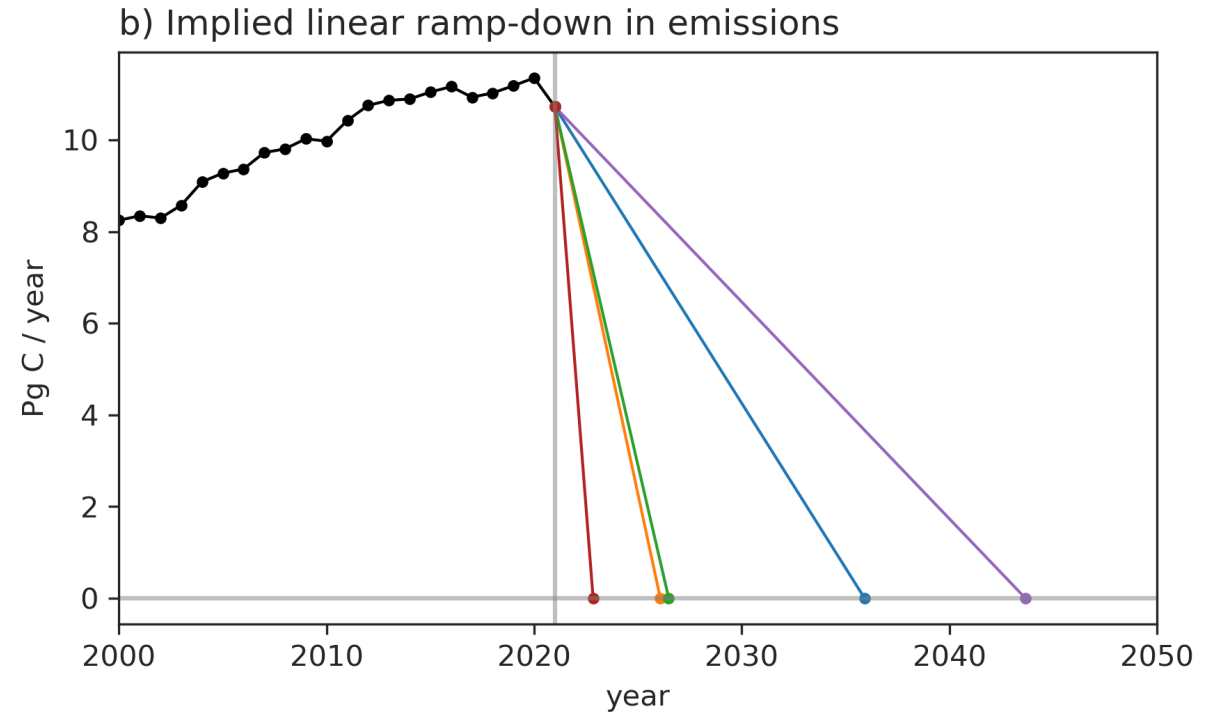
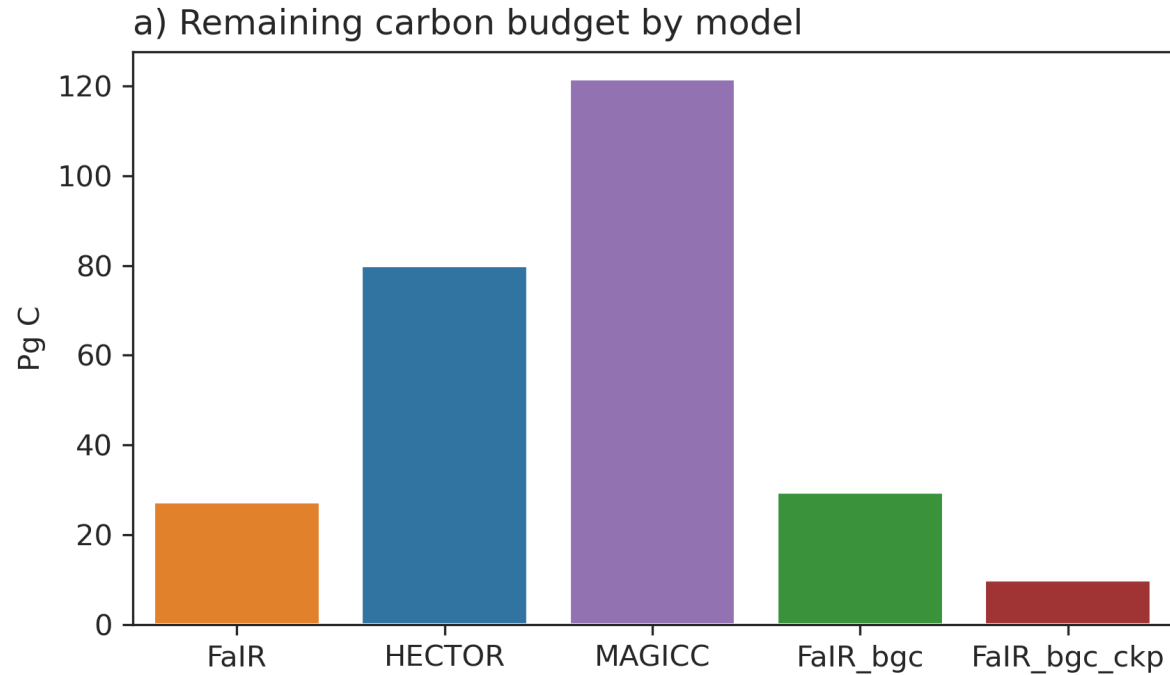


# Implications for remaining carbon budget?

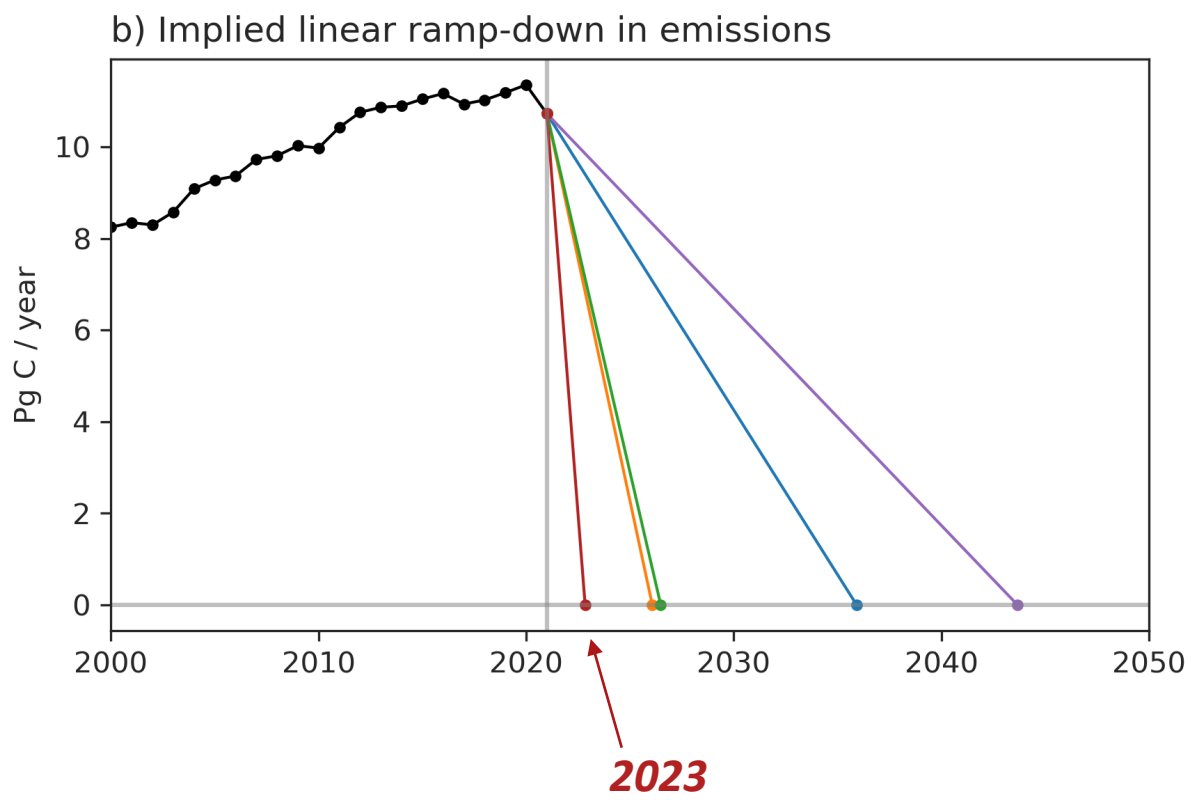
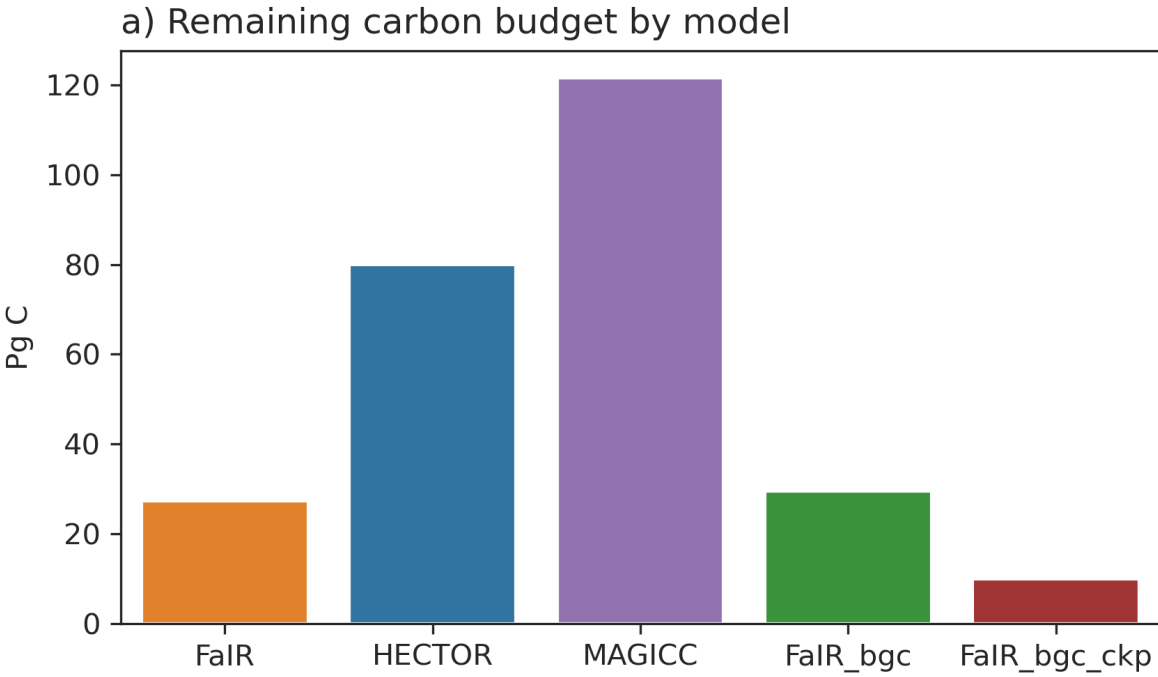




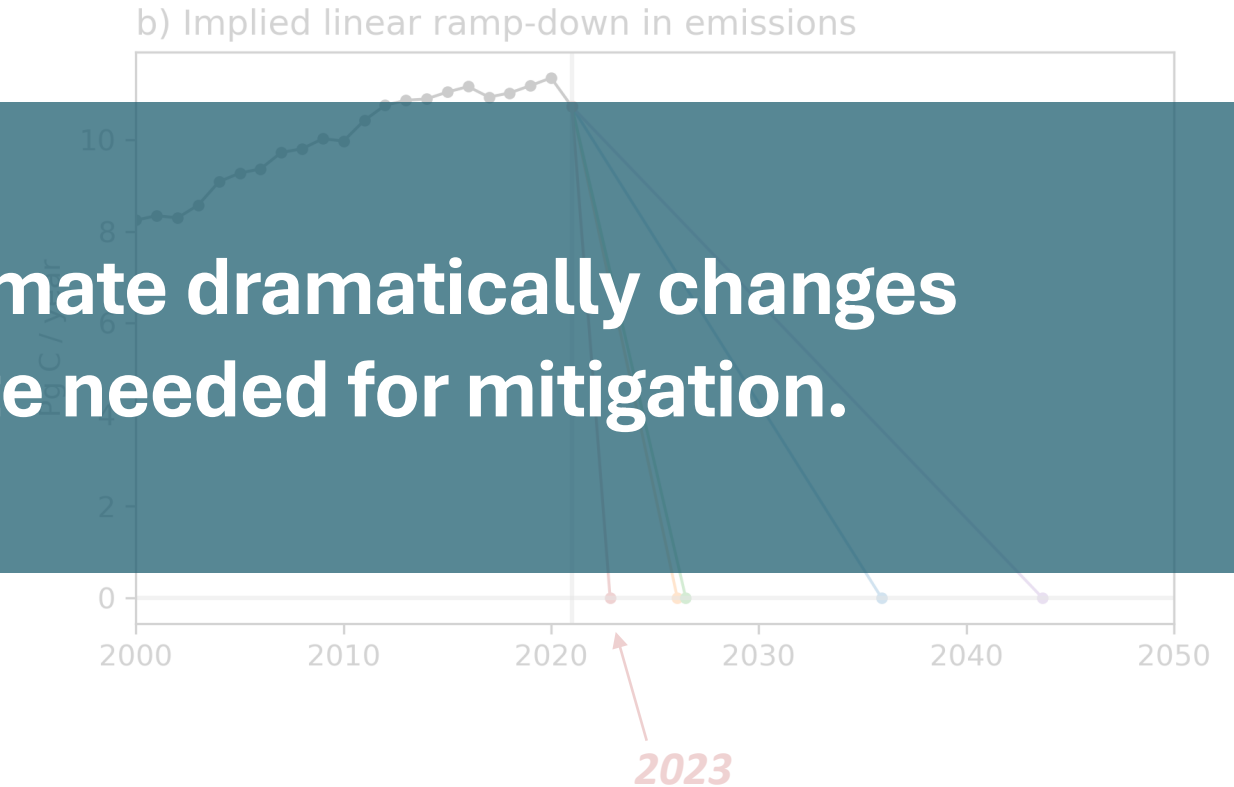
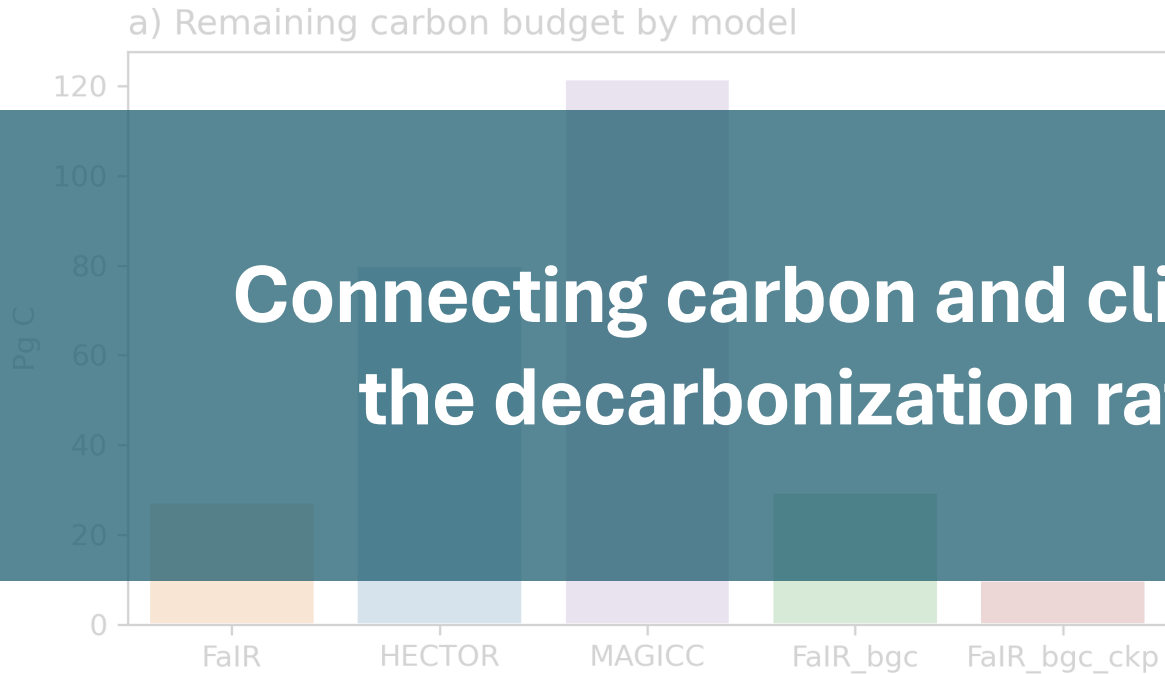
# Drop in RCB in correlated ensemble $\rightarrow$ different emissions necessary to meet mitigation goals.



# Drop in RCB in correlated ensemble → different emissions necessary to meet mitigation goals.



Drop in RCB in correlated ensemble → different emissions necessary to meet mitigation goals.



**Connecting carbon and climate dramatically changes the decarbonization rate needed for mitigation.**

Need for ESM-driven constraints on connections between processes that influence carbon and climate :

→ e.g. ESM carbon fluxes and heat fluxes as outputs from emissions-driven runs

→ Carbon cycle PPEs would allow us to sample across realizations of the carbon-climate system