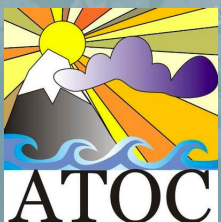


# Sea ice feedbacks cause more greenhouse cooling than greenhouse warming at high northern latitudes on multi-century timescales

**Jen Kay**

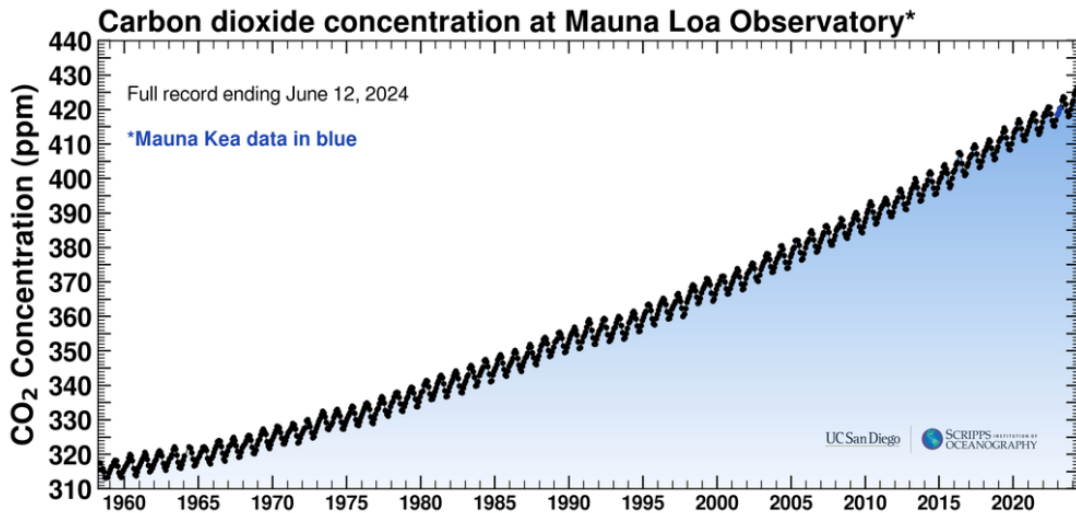
**Dept. Atmospheric and Oceanic Sciences (ATOC)  
Cooperative Institute for Research in Environmental Sciences (CIRES)  
University of Colorado Boulder (CU)**



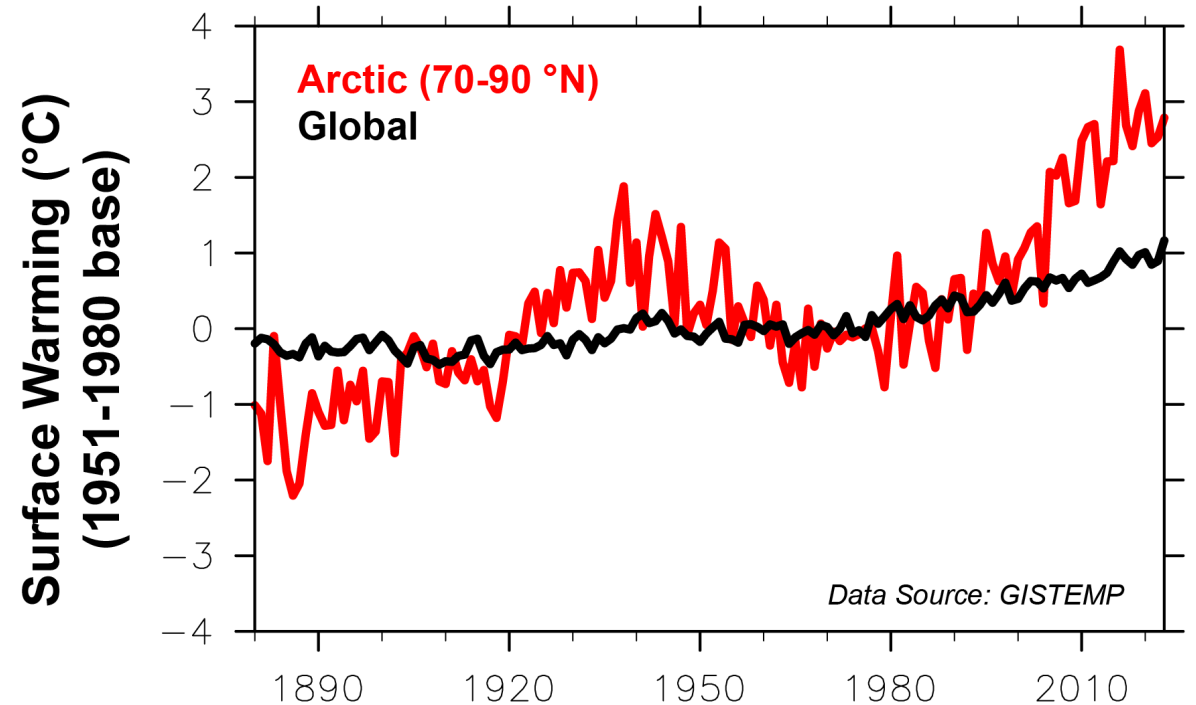
*Image: My foot in melting Arctic sea ice*

# Observed greenhouse gas increases and surface warming almost everywhere, **especially in the Arctic**

\*Latest CO<sub>2</sub> reading: **427.44 ppm**

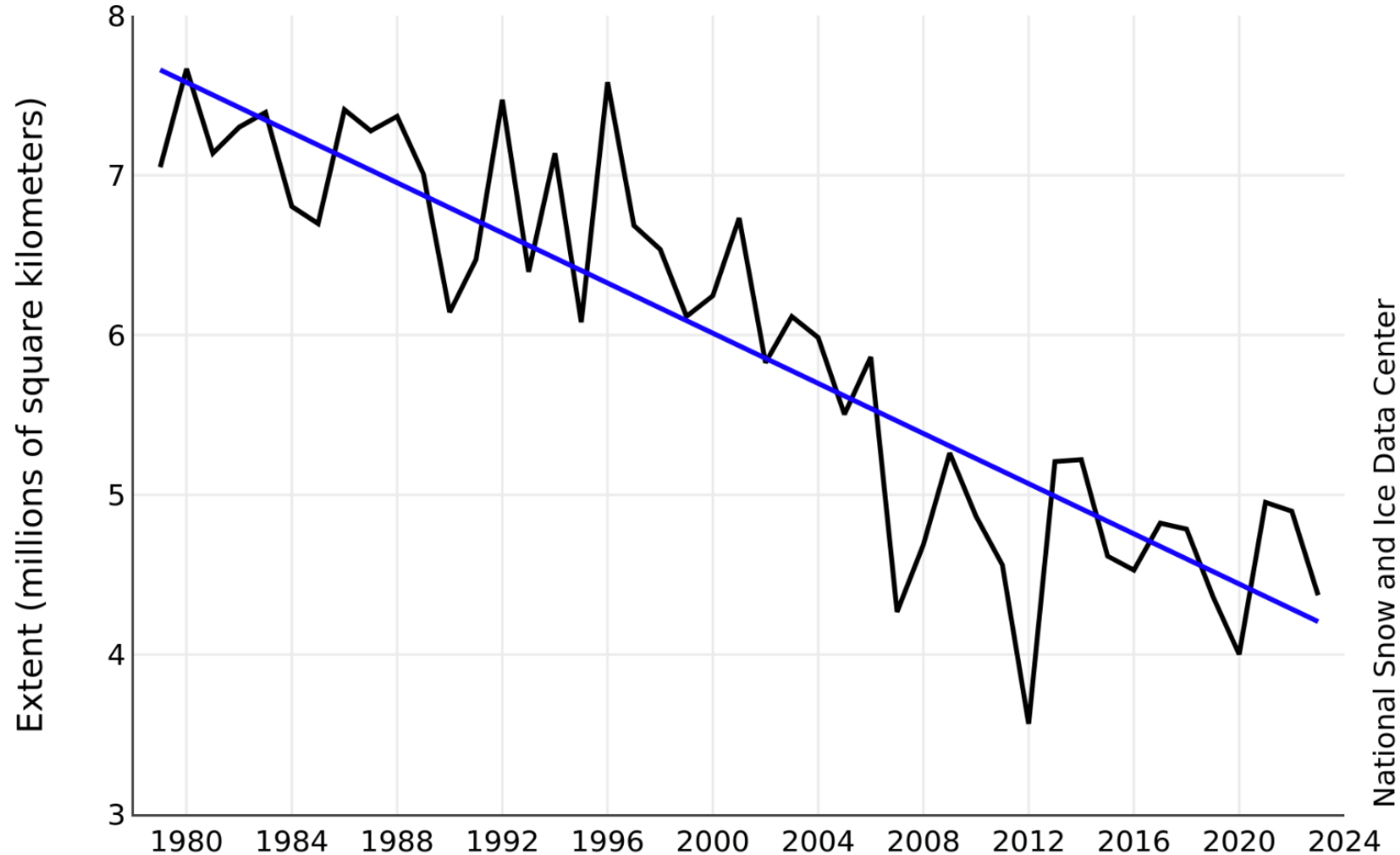


Data Source: <https://keelingcurve.ucsd.edu/>





Are you expecting a new September Arctic sea ice record minimum soon? **I am.**

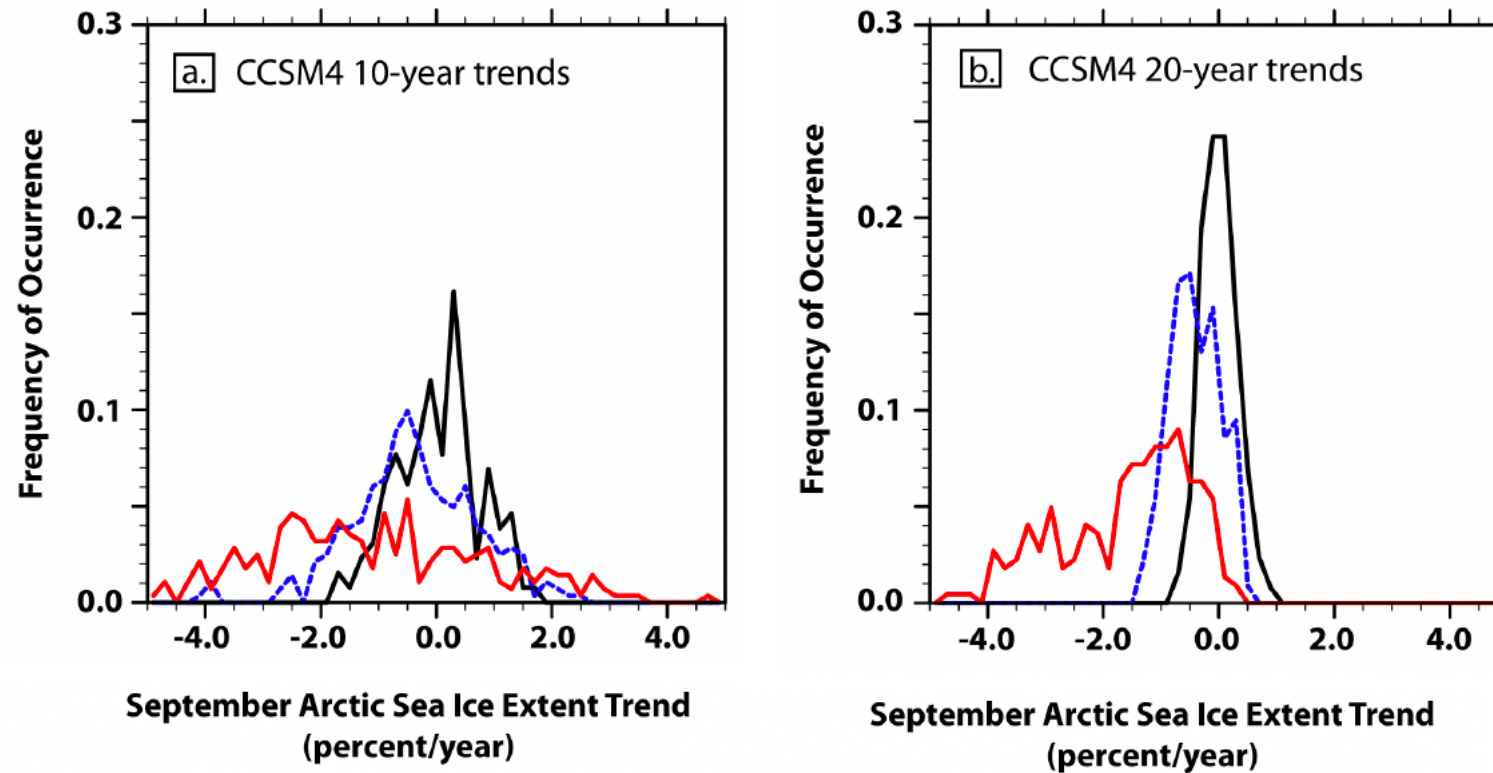


**FLASHBACK!**  
**FLASHBACK!**  
**FLASHBACK!**  
**FLASHBACK!**  
**FLASHBACK!**

## Inter-annual to multi-decadal Arctic sea ice extent trends in a warming world

Jennifer E. Kay,<sup>1</sup> Marika M. Holland,<sup>1</sup> and Alexandra Jahn<sup>1</sup>

Received 2 May 2011; revised 1 July 2011; accepted 8 July 2011; published 11 August 2011.



- 21stC transient (2006-2061, 21STC\_6)
- - - 20thC transient (1950-2005, 20THC\_6)
- 1850 control (1300 years, 1850CNT)

# DECK

amip

piControl and esm-piControl

1pctCO2

abrupt-4xCO2

piClim-control

piClim-anthro

piClim-4xCO2

historical and/or esm-historical

Additions to the DECK since CMIP6

# CMIP AR7 Fast Track

Climate services

DCPP

Initialised prediction  
(2025-2036)

ScenarioMIP

High scenario

Medium scenario

Low scenario 1

Low scenario 2

Low scenario 3

Process understanding

AerChemMIP

piClim-X

hist-piSLCF/hist-piAer

SSPX-SLCF

C4MIP

1pctCO2-bgc

1pctCO2-rad

esm-flat10-cdr

esm-flat10-zec

CFMIP

amip-p4k

amip-piForcing

abrupt-2xCO2

abrupt-0p5CO2

DAMIP

hist-nat

hist-aer

hist-GHG

GeoMIP

newGeoMIP

LMIP

land-hist

PMIP

LIGabrupt

RFMIP

piClim-histaer

piClim-aer

piClim-histall

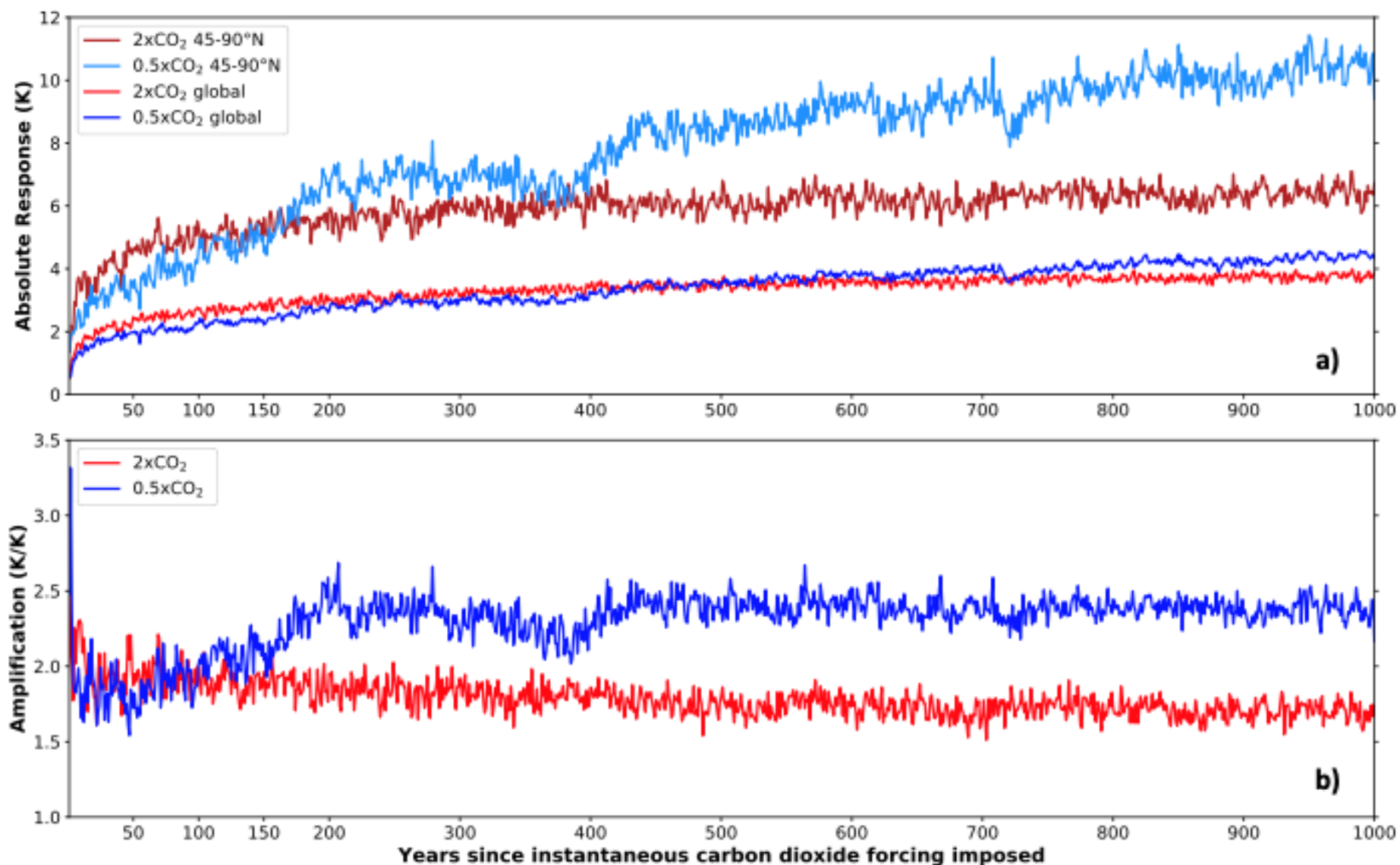
**Challenge: Many experiments only run 150 years (per CMIP protocols).**

**Solution: Extend existing experiments to provide context on climate sensitivity constraints. (e.g., <https://www.longrunmip.org/>).**

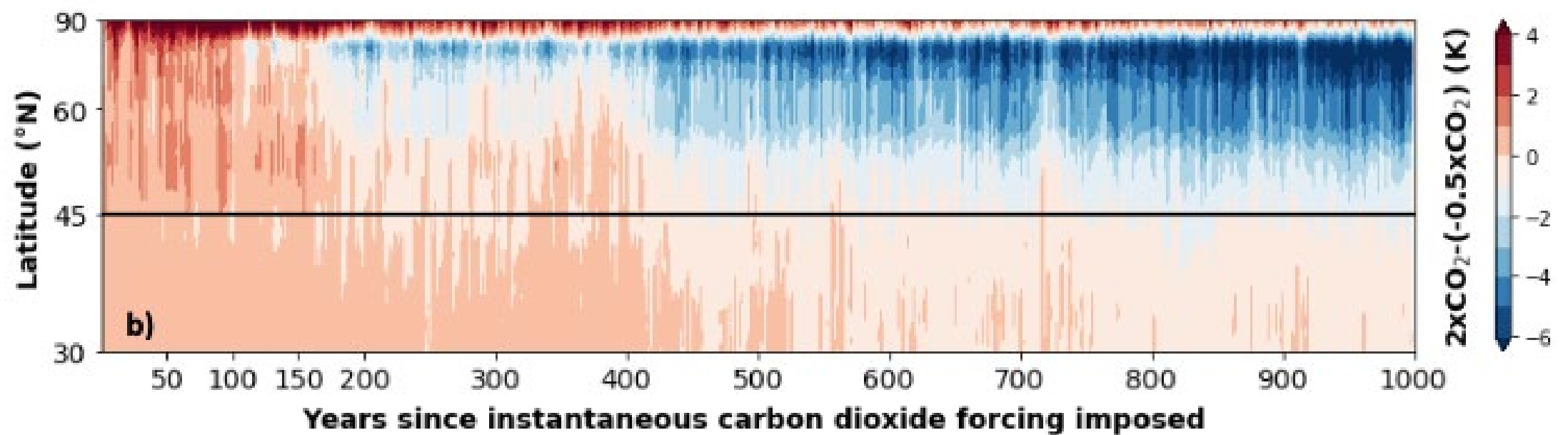
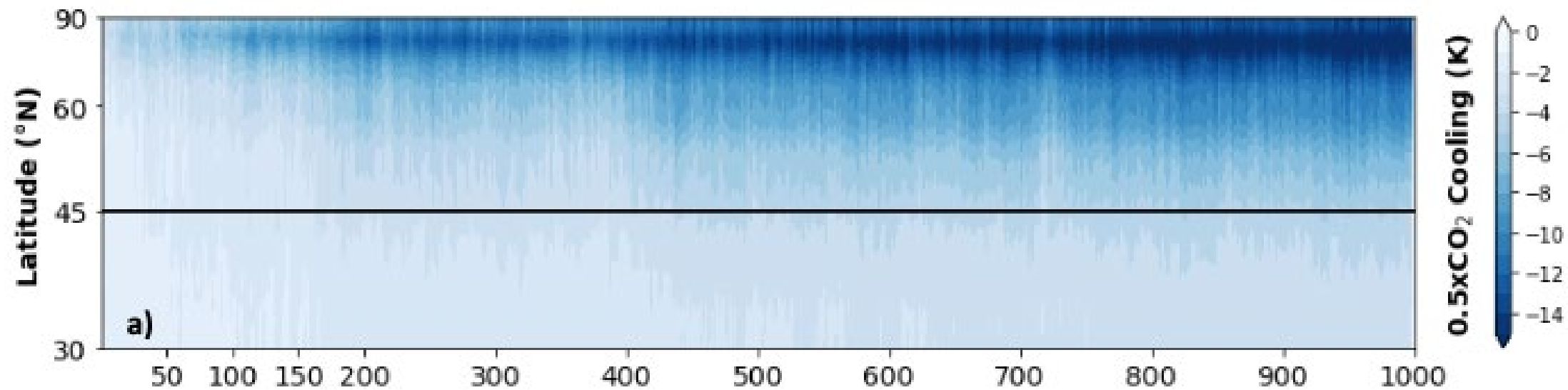
	Description	Reference	tau global	tau 45-90 °N	tau 70-90 °N
<b>CNT</b>	Fully coupled 1850 control	Kay et al. 2015	n/a	n/a	n/a
<b>2xCO<sub>2</sub></b>	Abrupt 2xCO <sub>2</sub> fully coupled	Frey and Kay 2017 (years 1-150); this work (years 151-1000)	219 yrs.	200 yrs.	200 yrs.
<b>0.5xCO<sub>2</sub></b>	Abrupt 0.5xCO <sub>2</sub> fully coupled	Chalmers et al. (years 1-150); this work (years 151-1000)	567 yrs.	437 yrs.	284 yrs.

**Table 1. Model runs used in this study.** All runs use the Community Earth System Model version 1 Large Ensemble version (CESM1-LE, Kay et al. 2015) at the standard one-degree horizontal resolution. Last four columns on the right provide equilibration timescale (years) for the surface temperature response. The equilibration timescale (tau) found by fitting surface temperature (TS) data over all 1000 years of the simulation to an exponential:  $TS(\text{year})=a\text{EXP}(\text{tau}*\text{year})$ .

**Long runs more important for cooling than for warming!**

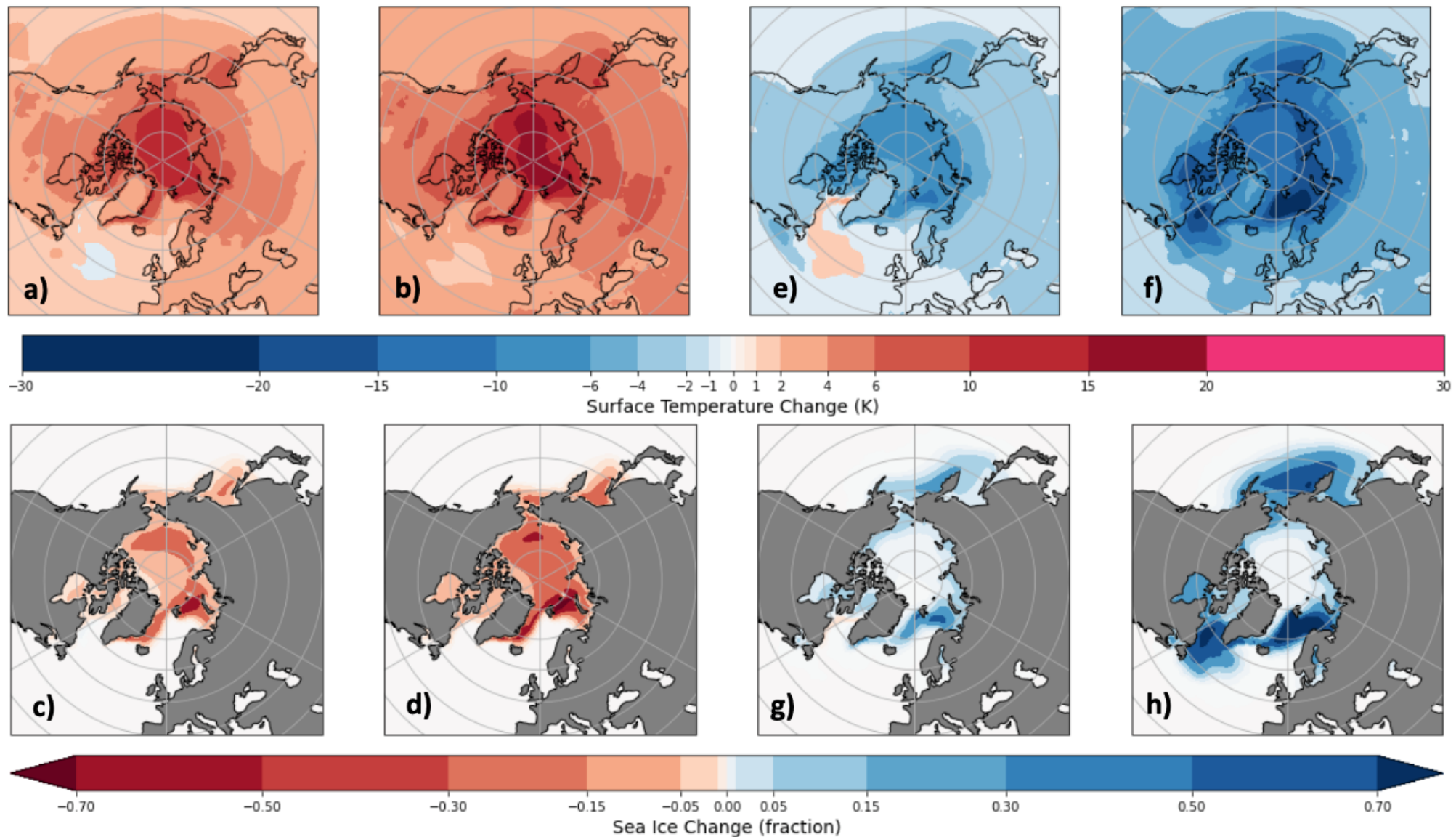


**Figure 1.** Time series of annual mean surface temperature responses to 2xCO<sub>2</sub> and 0.5xCO<sub>2</sub> instantaneous carbon dioxide forcing: a) Global and high northern latitudes (45-90 °N) absolute response; b) Amplification, i.e., high northern latitude values divided by global values. All responses are relative to the 1850 control run (CNT years 400-1399, see Table 1).



Years since instantaneous carbon dioxide forcing imposed





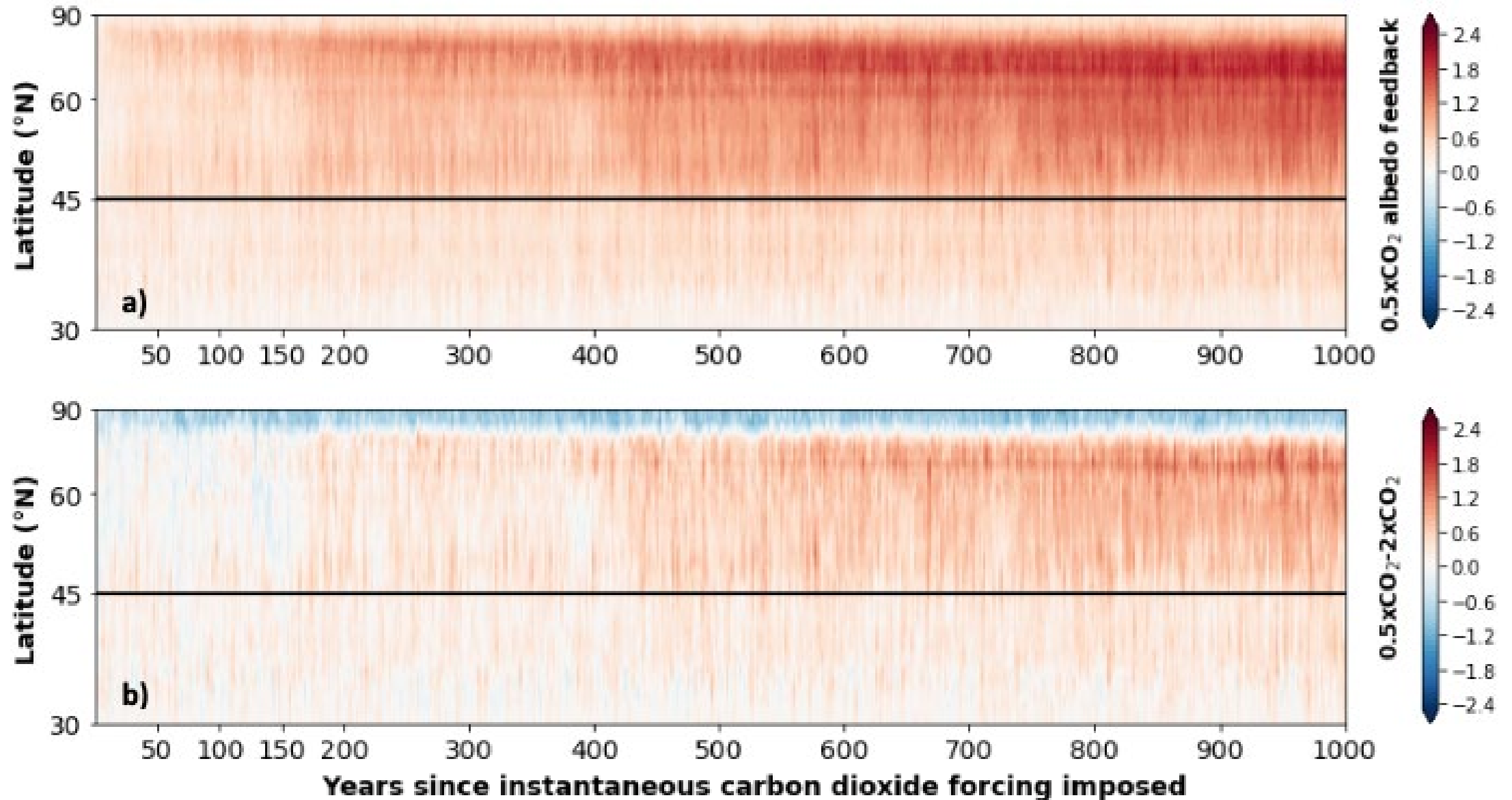
**Figure 3. Polar maps of annual high northern latitude response: a) 2xCO<sub>2</sub> surface temperature early (years 1-150), b) 2xCO<sub>2</sub> surface temperature late (years 850-100), c) 0.5xCO<sub>2</sub> surface temperature early, d) 0.5xCO<sub>2</sub> surface temperature late, e-h) as in a-d) but for sea ice. All responses are relative to the 1850 control (years 400-1399 CNT, Table 1).**

# Why more **cooling** than **warming**?

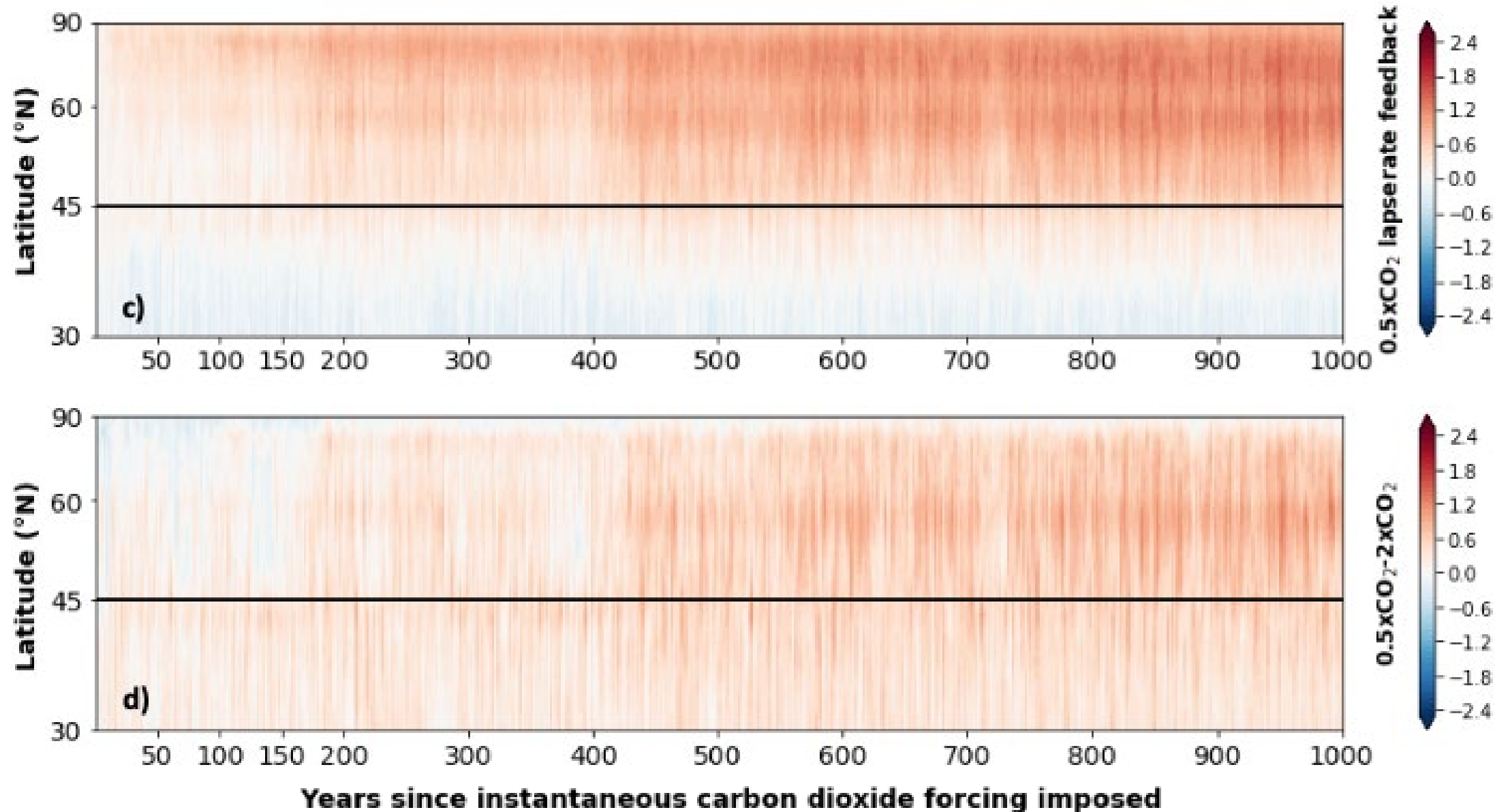
It's not the CO<sub>2</sub> radiative forcing.  
*2xCO<sub>2</sub> forcing >> 0.5x CO<sub>2</sub> forcing*

Description	CNT_2×CO <sub>2</sub>	CNT_0.5×CO <sub>2</sub>
Global mean surface temperature response (K)	2.75 ± 0.03	−2.28 ± 0.03
ERF_fSST (W m <sup>−2</sup> )	3.82 ± 0.05	−3.44 ± 0.06
ERF_reg (W m <sup>−2</sup> )	3.84 ± 0.69	−3.26 ± 0.58

# It's the sea-ice associated feedbacks: **surface albedo** and lapse rate



# It's the sea-ice associated feedbacks: surface albedo and **lapse rate**





# Summary

- After 150 years – CESM1 shows more 2xCO<sub>2</sub> warming than 0.5xCO<sub>2</sub> cooling due to stronger 2xCO<sub>2</sub> radiative forcing and more positive 2xCO<sub>2</sub> feedbacks.
- After 1000 years – CESM1 shows more 0.5xCO<sub>2</sub> cooling than 2xCO<sub>2</sub> warming due to feedbacks associated with sea ice expansion (positive lapse rate and surface albedo feedbacks).
- How do we use 0.5xCO<sub>2</sub> experiments (with LGM) to constrain climate sensitivity?

Support provided by University of Colorado  
CIRES, NSF CAREER (1554659), NSF OPP  
(2233420).





**FYI**

**For Your Information**



# Climate models have long shown greater-than-global Arctic surface warming due to unique positive feedbacks (surface albedo, lapse rate)

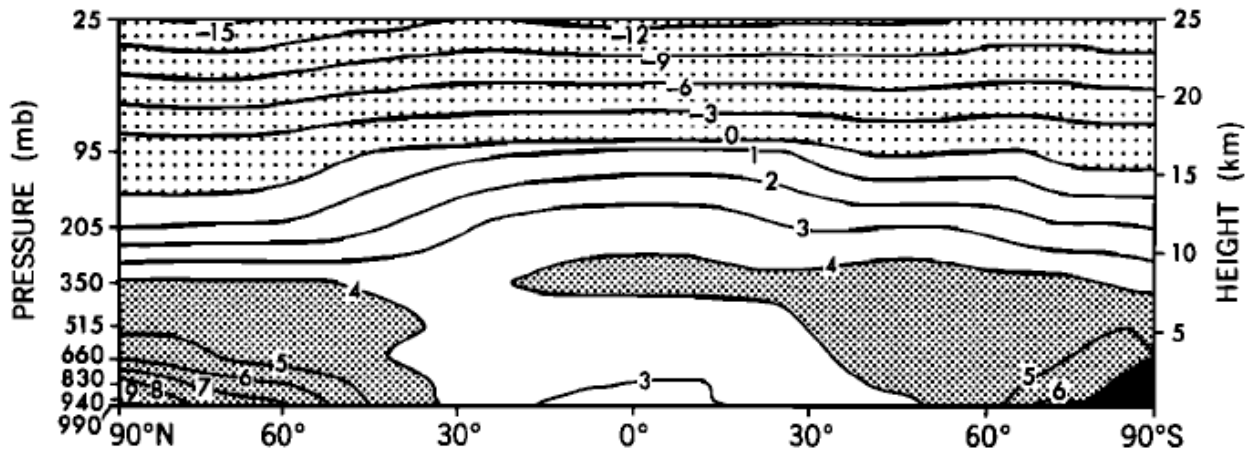


Fig. 14. Latitude height distribution of the zonal mean difference in annual mean temperature (degrees Kelvin) of the model atmosphere between the  $4 \times \text{CO}_2$  and  $1 \times \text{CO}_2$  experiment.

Figures from Manabe and Stouffer 1980  
Similar results in many more recent studies.

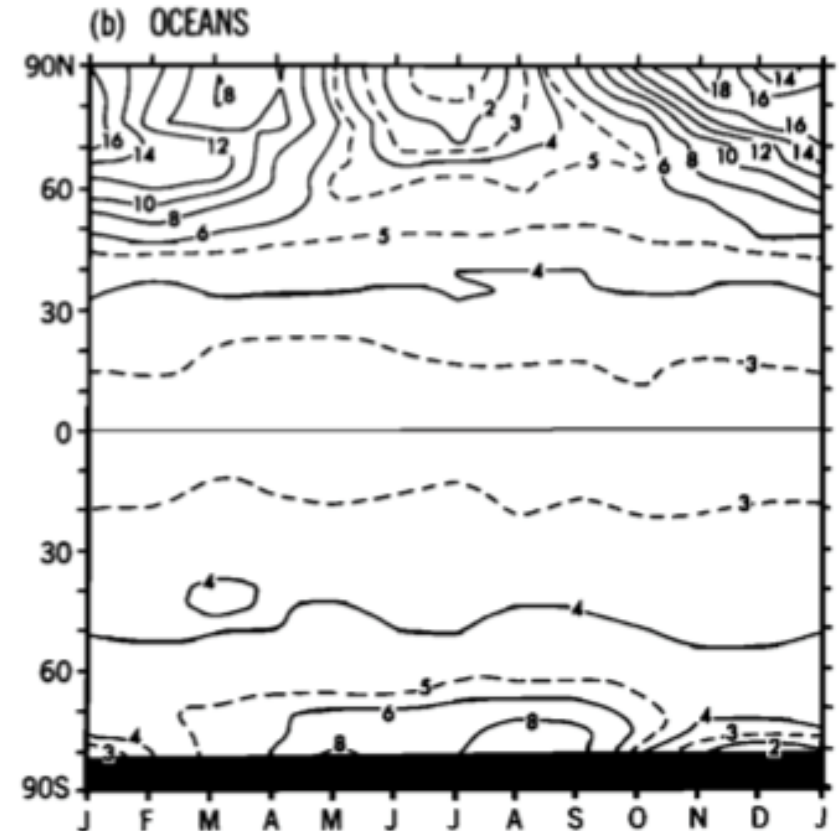
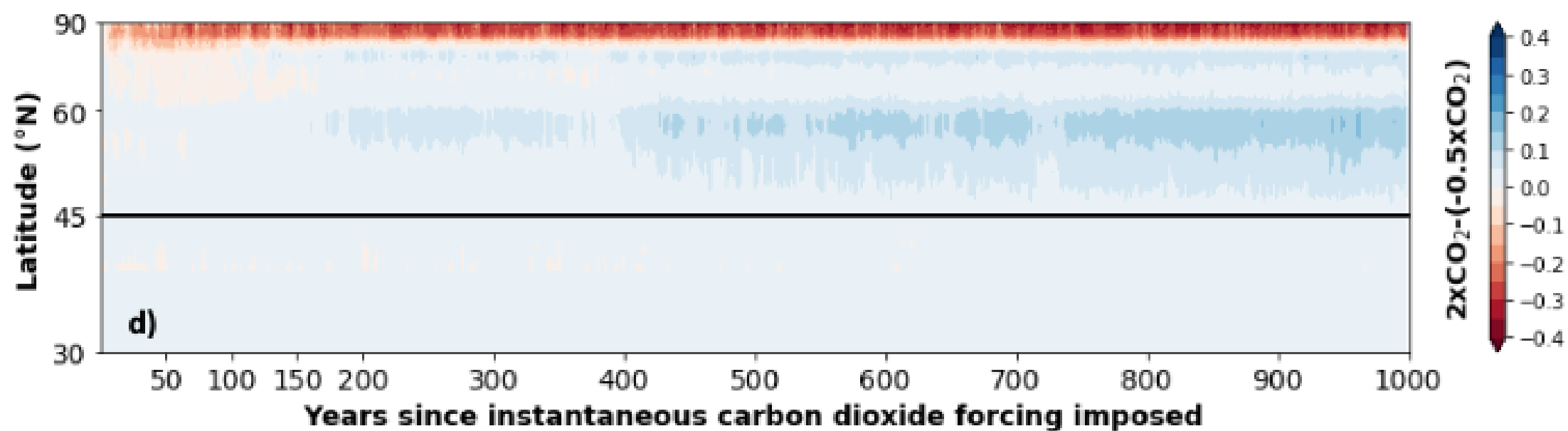
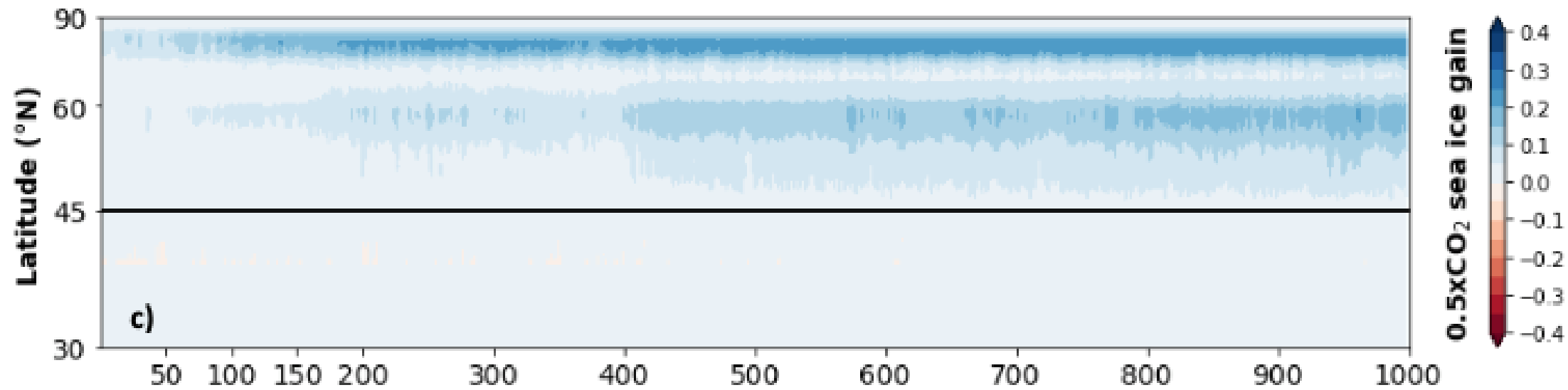
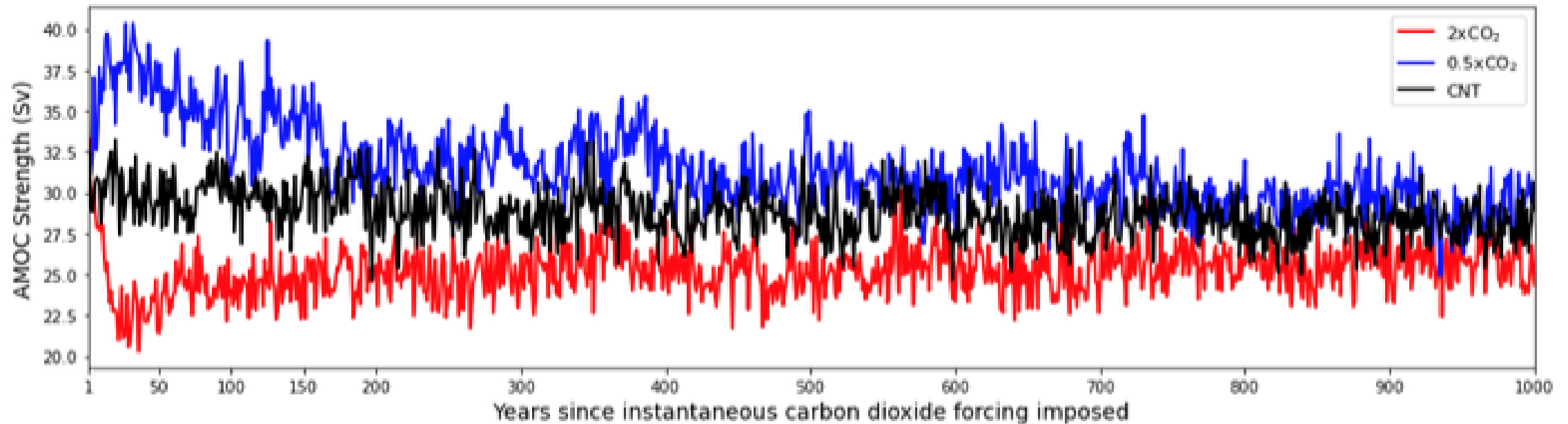


Fig. 16. Latitude-time distribution of zonal mean difference in surface air temperature (degrees Kelvin) between the  $4 \times \text{CO}_2$  and  $1 \times \text{CO}_2$  experiment. (a) Oceans and continents, (b) oceans, and (c) continents. The difference is computed at the lowest finite difference level located at about 70 m above the earth's surface.

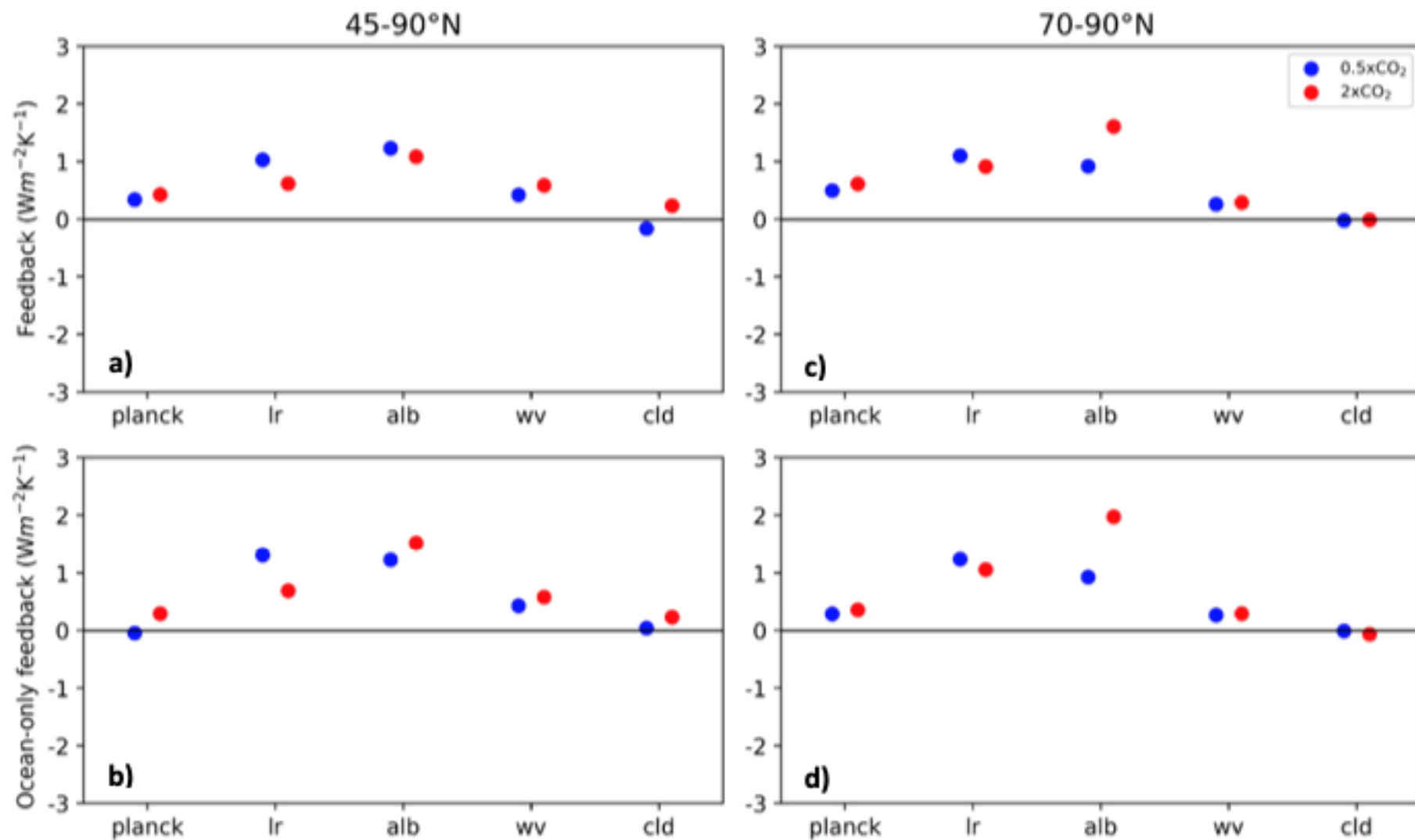


Years since instantaneous carbon dioxide forcing imposed



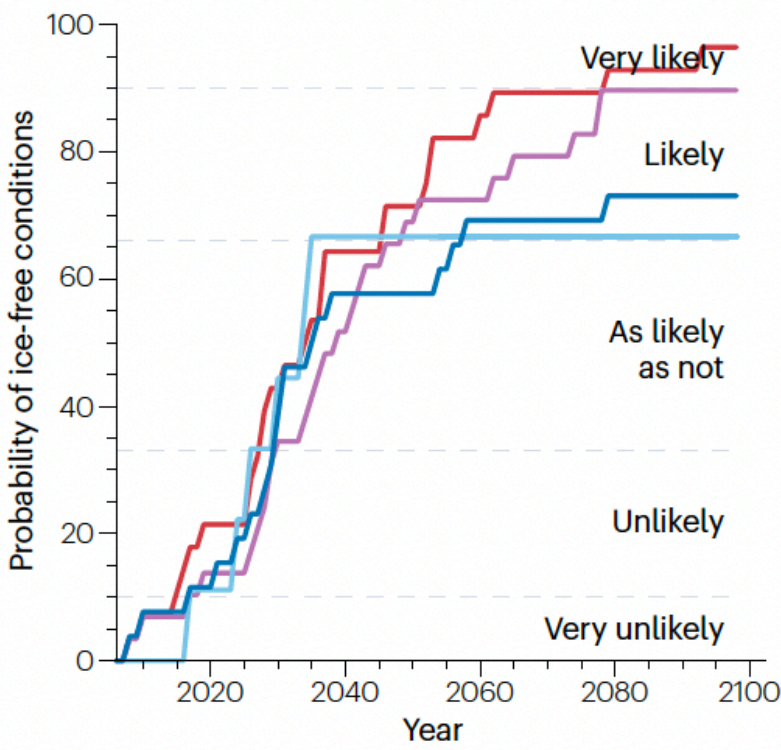


**Figure 4. Timeseries of Annual Mean Max Atlantic Meridional Overturning Circulation (AMOC) strength.** AMOC strength defined as the maximum meridional overturning in Sverdrups (Sv) in the North Atlantic (30 to 55 degrees N and 800-2,000 meters depth). CNT plotted for years 400-1399.

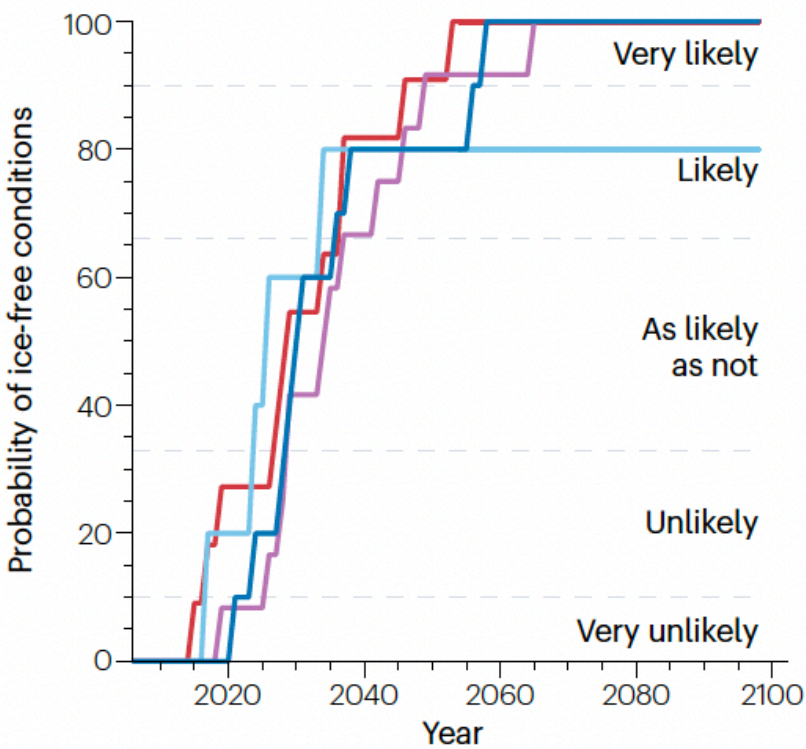


**Figure 5. High Northern Hemisphere Radiative Feedback Parameters: a) 45-90 °N, b) 70-90 °N, c-d) as in a-b) but for ocean only.** Feedback parameters calculated for years 850-999 with CNT years 400-549 as a reference climate for the Planck (“planck”), lapse rate (“lr”), surface albedo (“alb”), water vapor (“wv”) and cloud (“cld”) feedbacks.

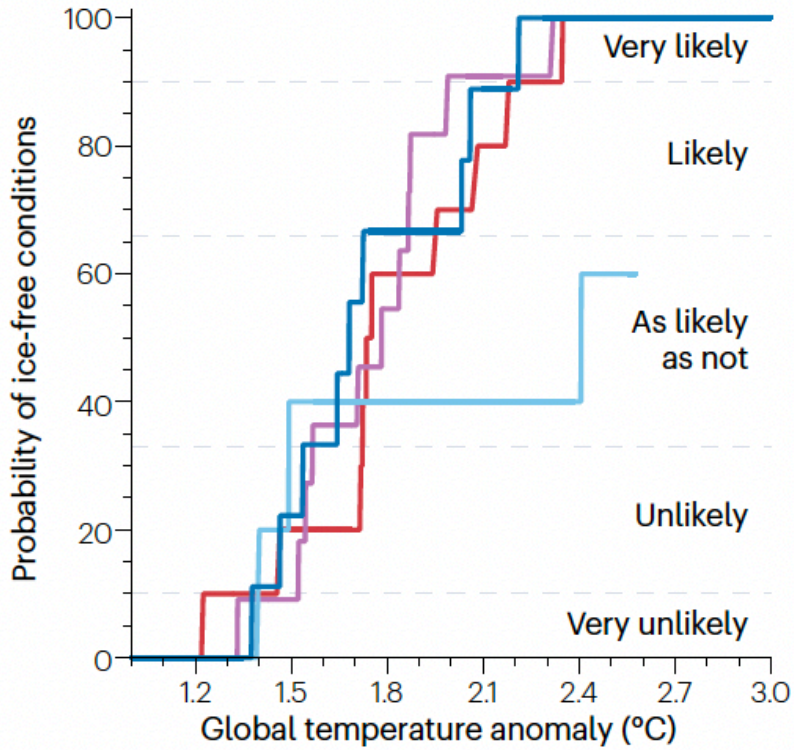
**b** Cumulative probability of first ice-free conditions in September in CMIP6 models



**c** Cumulative probability of first ice-free conditions in September in selected CMIP6 models



**d** Cumulative probability of first ice-free conditions in September in selected CMIP6 models in terms of global warming



*“Although definition differences and model selection influence the specifics of ice-free predictions, **they all indicate that ice-free conditions tend to occur at least once by 2050 under all assessed scenarios**, and become a frequent occurrence thereafter under all scenarios except SSP1-1.9.”*

Figure 3 from Jahn, Holland, and Kay (2024)



### a Year of first September ice-free conditions

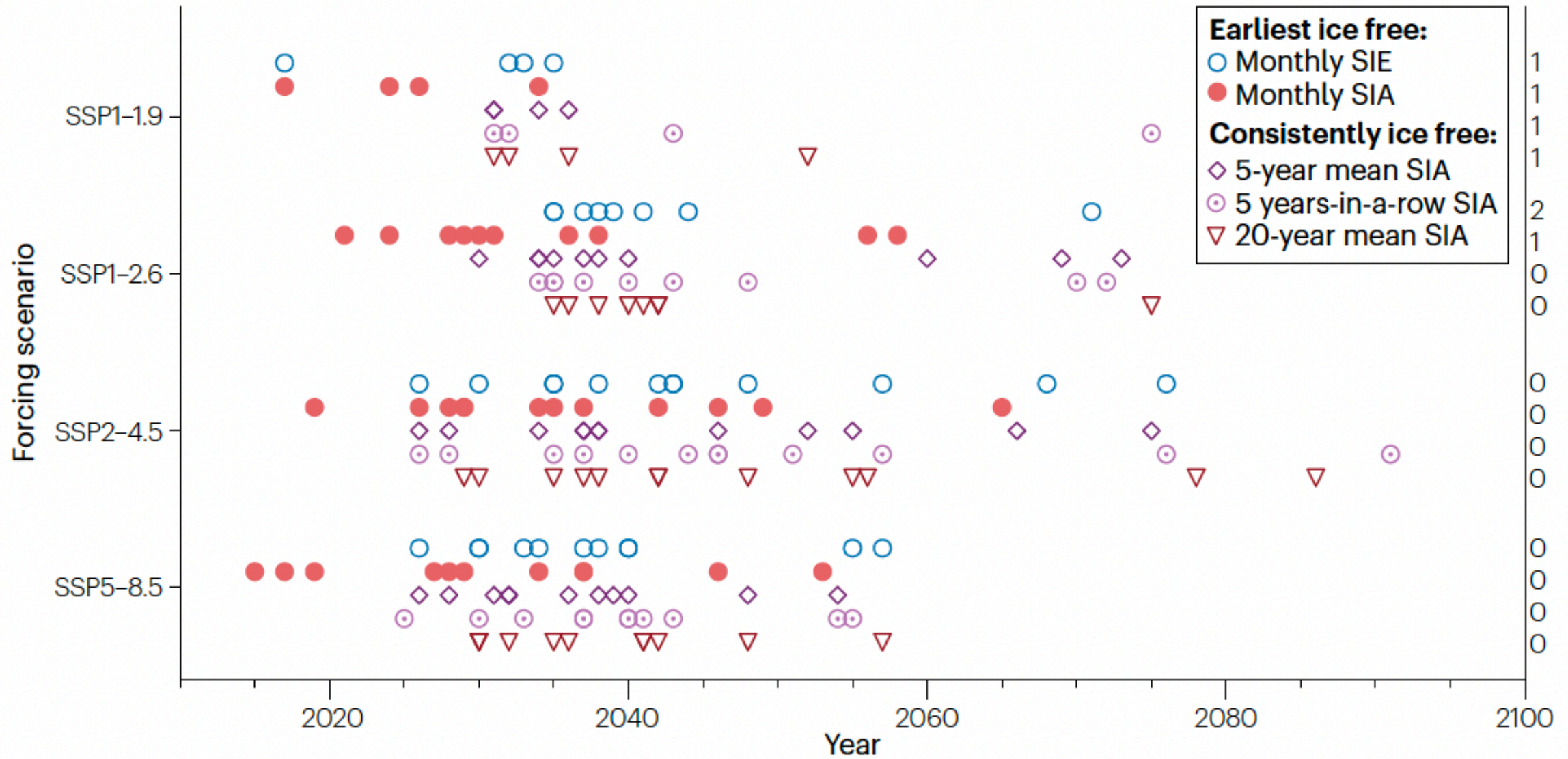


Figure 3 from Jahn, Holland, and Kay (2024)