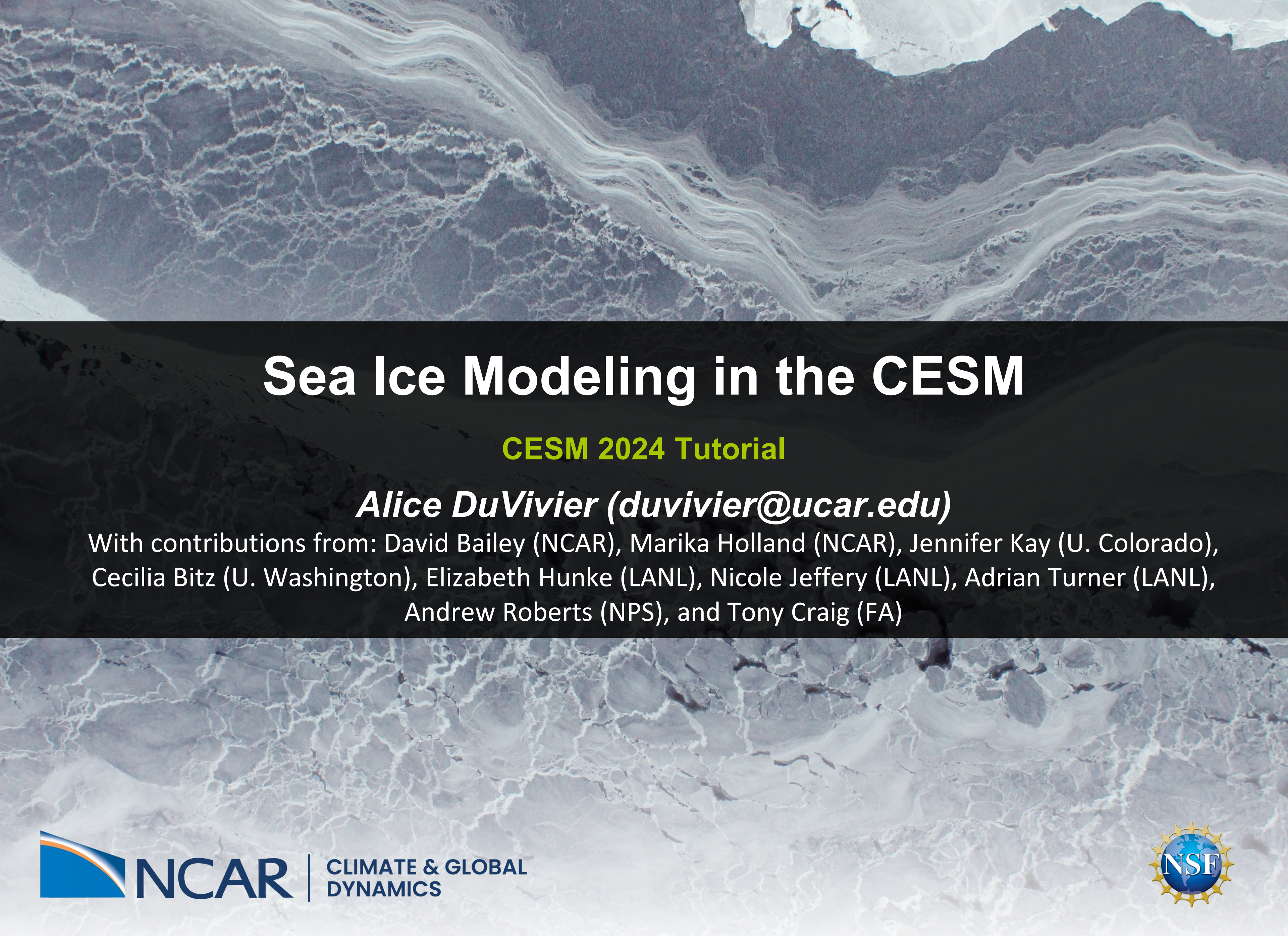

Sea Ice Modeling in the CESM

CESM 2024 Tutorial

Alice DuVivier (duvivier@ucar.edu)

With contributions from: David Bailey (NCAR), Marika Holland (NCAR), Jennifer Kay (U. Colorado), Cecilia Bitz (U. Washington), Elizabeth Hunke (LANL), Nicole Jeffery (LANL), Adrian Turner (LANL), Andrew Roberts (NPS), and Tony Craig (FA)

NCAR | CLIMATE & GLOBAL DYNAMICS

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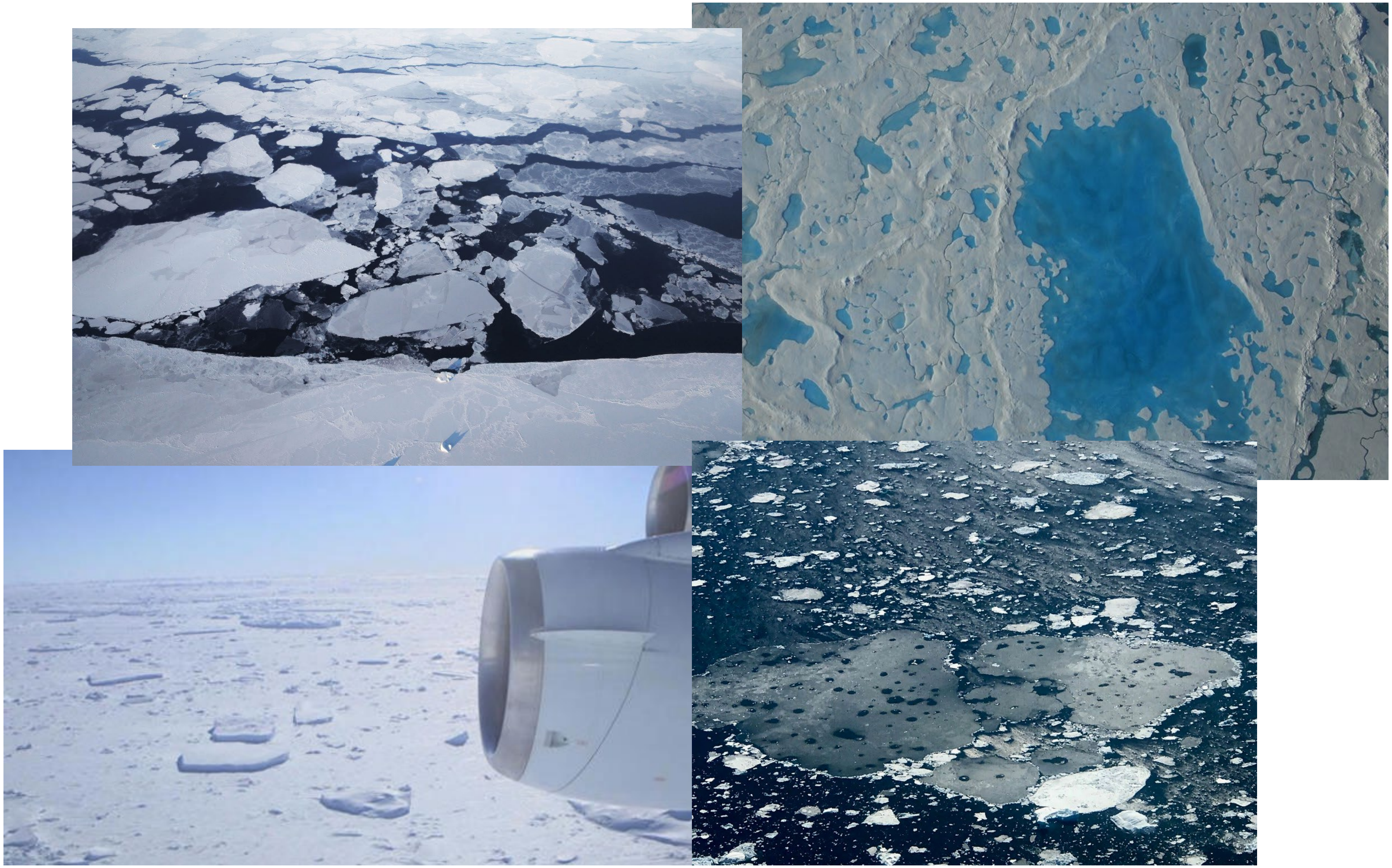
What is Sea Ice?

Sea Ice is frozen sea water that forms seasonally



Photos from NASA Operation IceBridge

Arctic vs. Antarctic

Arctic

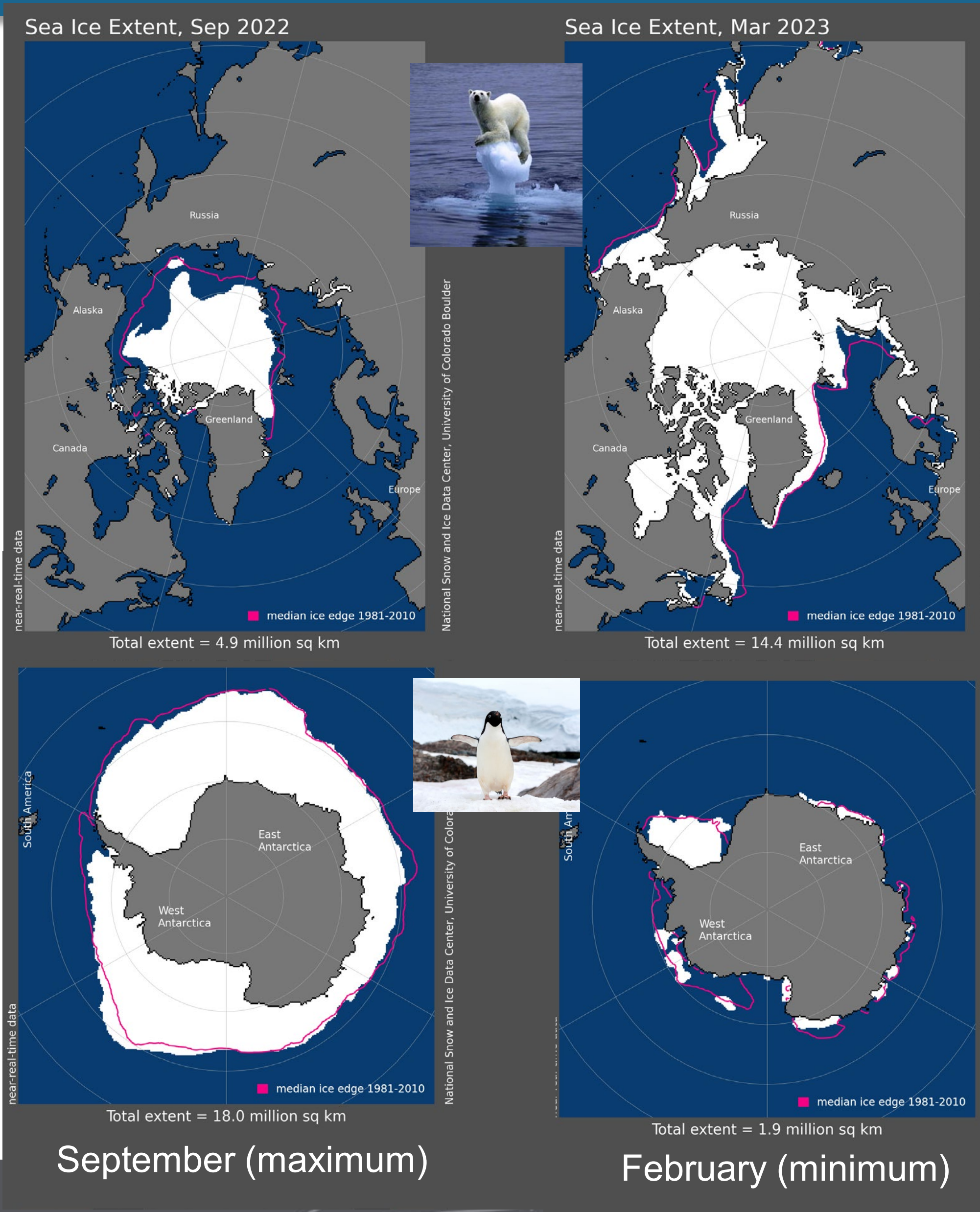
- Ocean bounded by land → ice converges at land, thick!
- Extent seasonal cycle: ~ 5→12 x10^6 km^2
- Land boundaries & ocean heat determine winter extent

Antarctic

- Unbounded → ice in free drift
- Extent seasonal cycle: ~ 2→15 x10^6 km^2
- Ocean heat determines winter extent

September (minimum)

March (maximum)

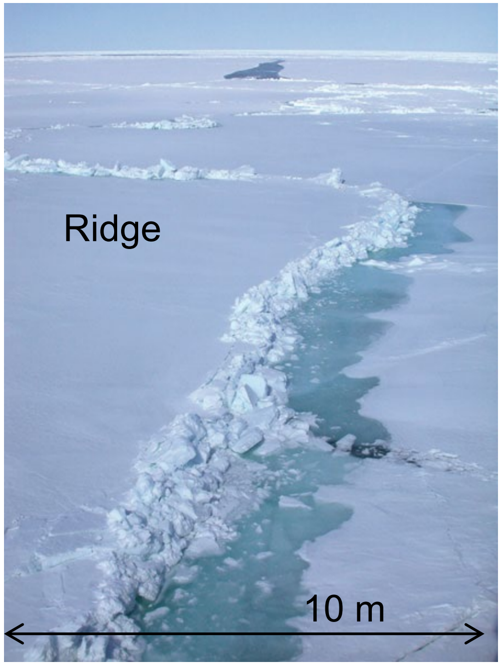


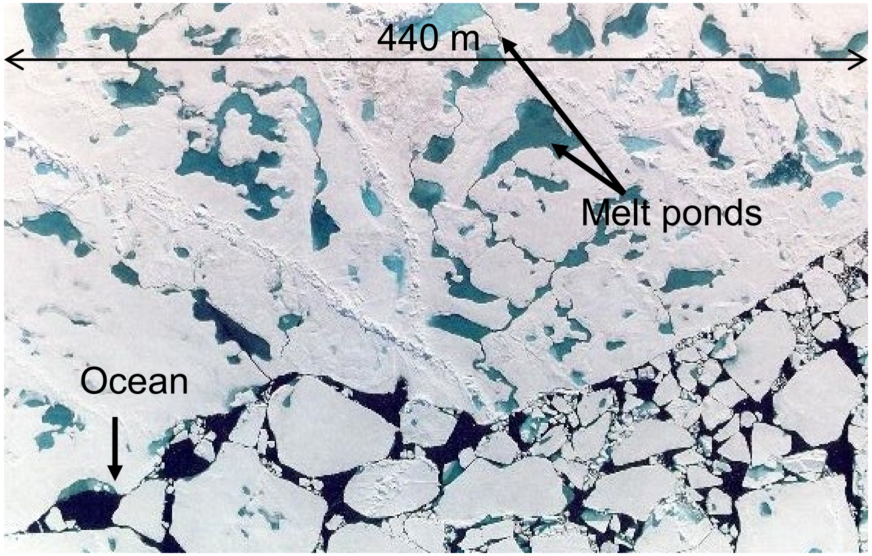
September (maximum)

February (minimum)

Figures from NSIDC

Sea ice Cover



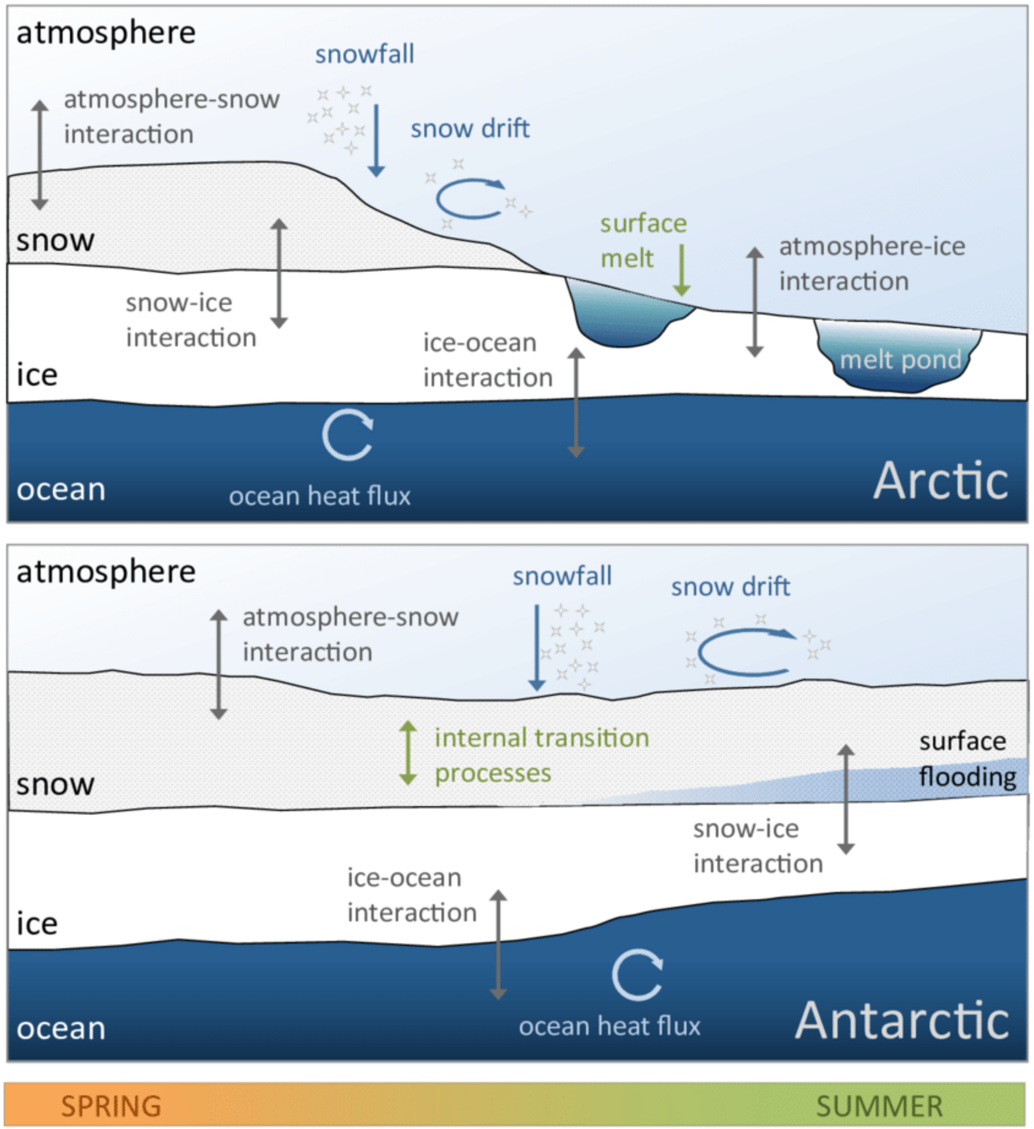


← Photos from Feltham, 2008 by Hajo Eicken

↑ Photo from Don Perovich

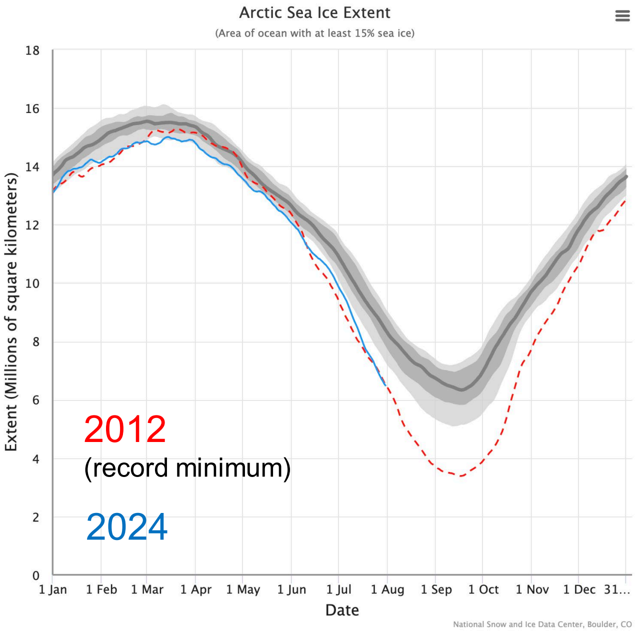
- Heterogeneous – lots of subgridscale variability
  - Leads, ridges, melt ponds, floes, albedo, snow cover, etc.
- Individual floes of varying size can form a continuous cover
- Thickness on the order of meters

Arctic vs. Antarctic – seasonal evolution

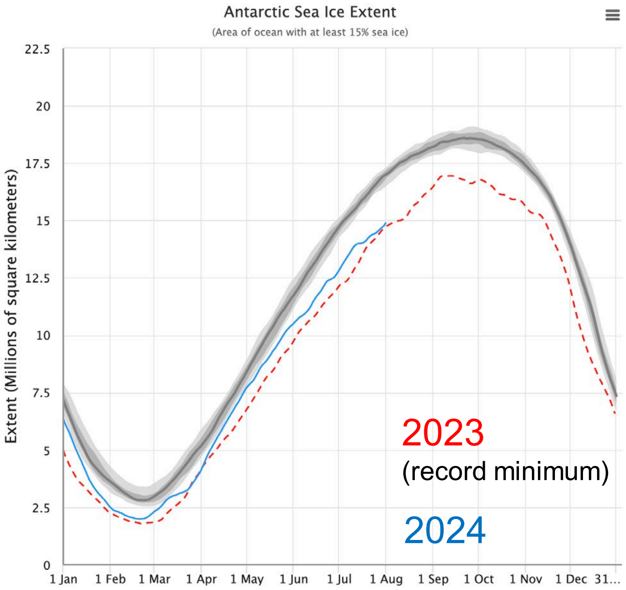


Arndt et al. 2017

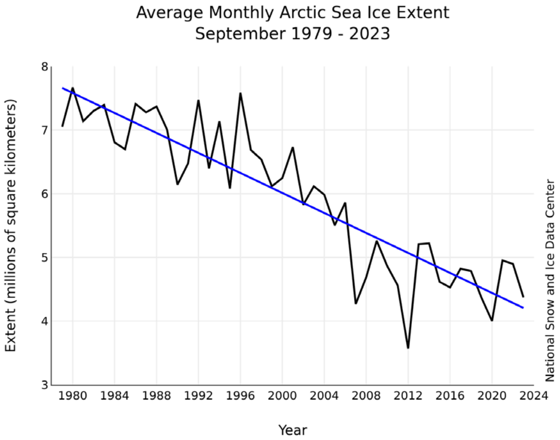
Why do we care about sea ice in climate models?

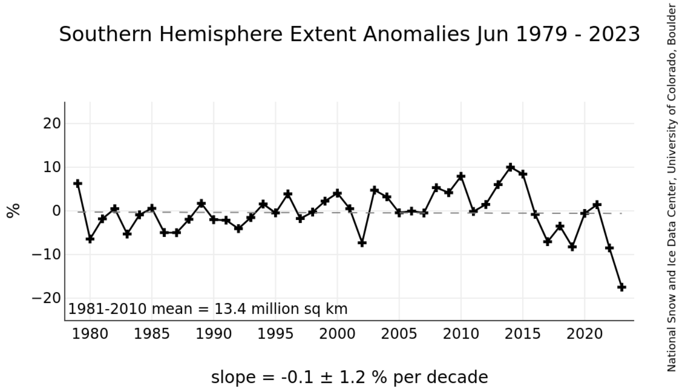


2012
(record minimum)
2024

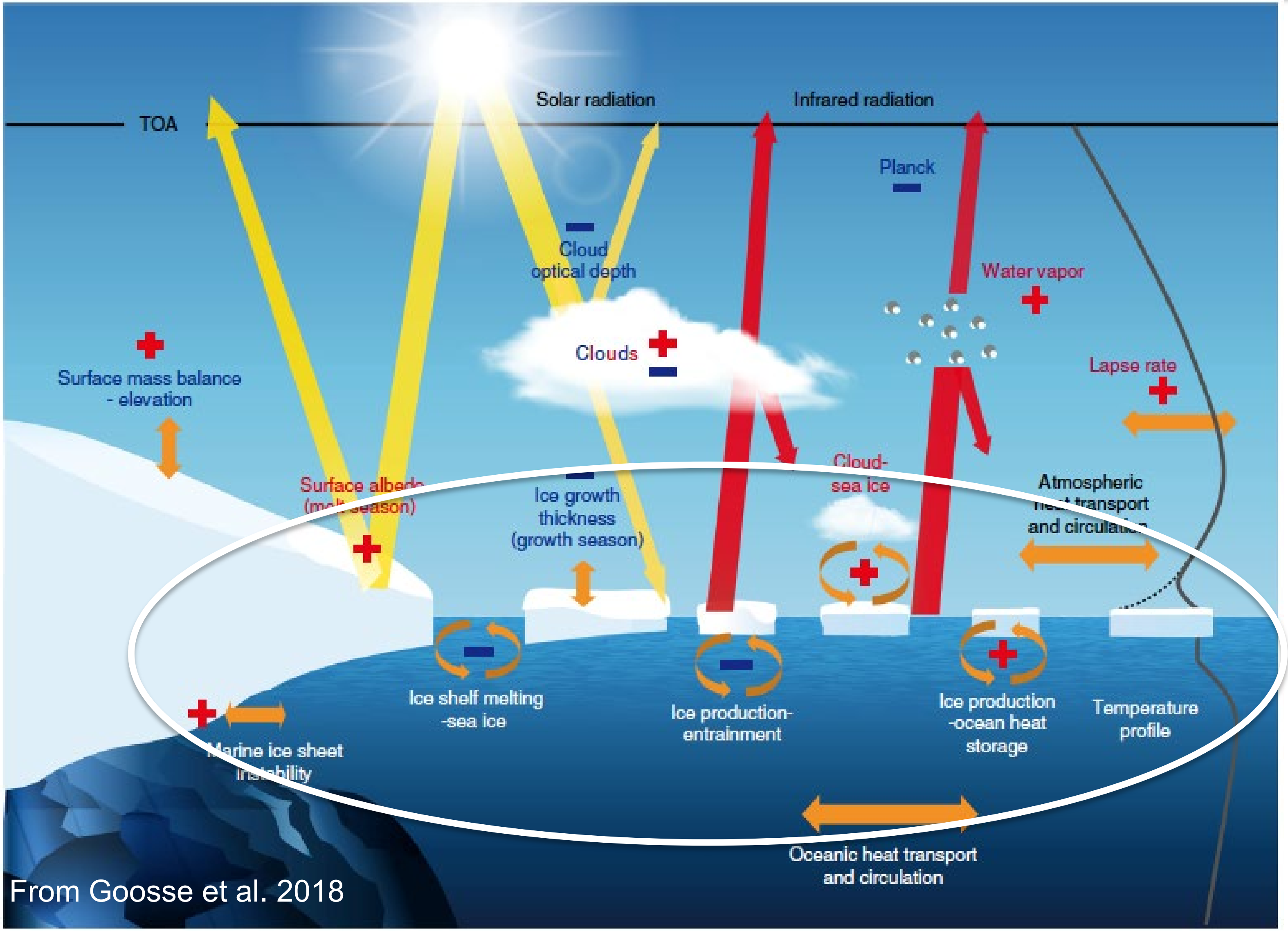


2023
(record minimum)
2024



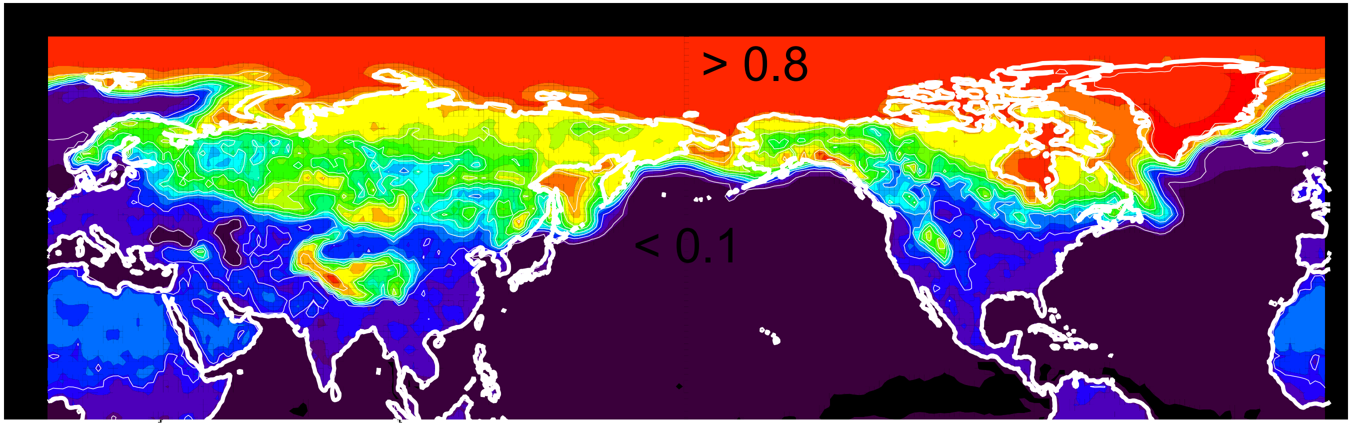


Why sea ice matters: Climate Feedbacks

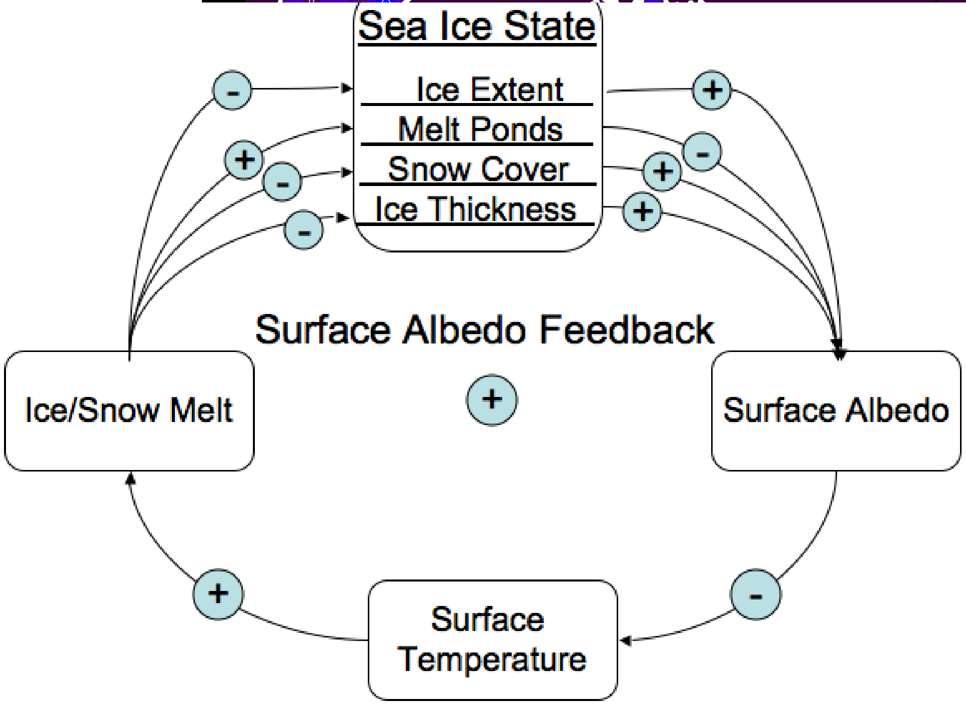


From Goosse et al. 2018

Why sea ice matters: Surface energy budget

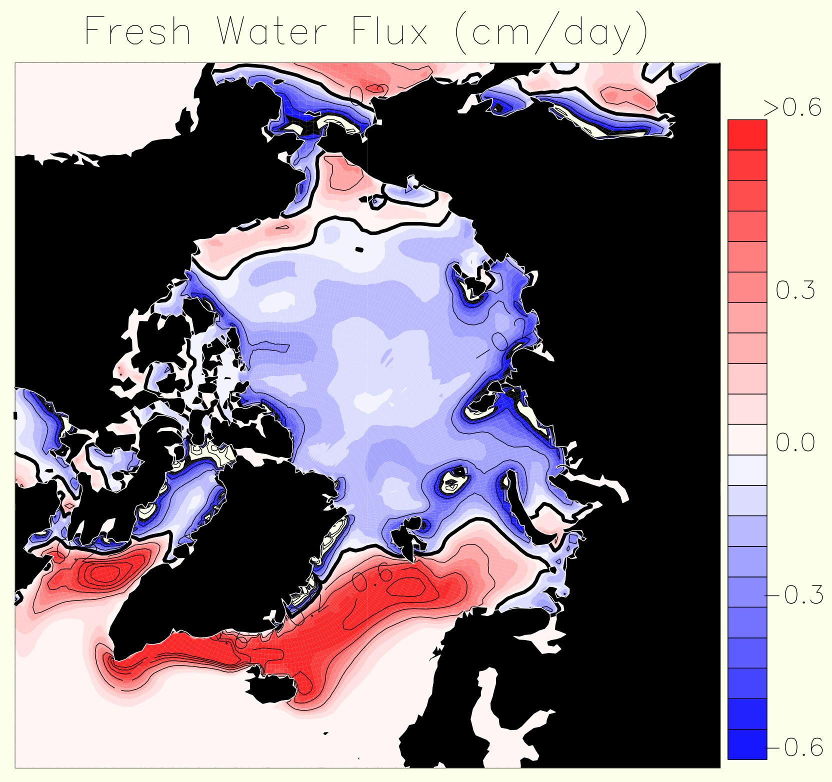


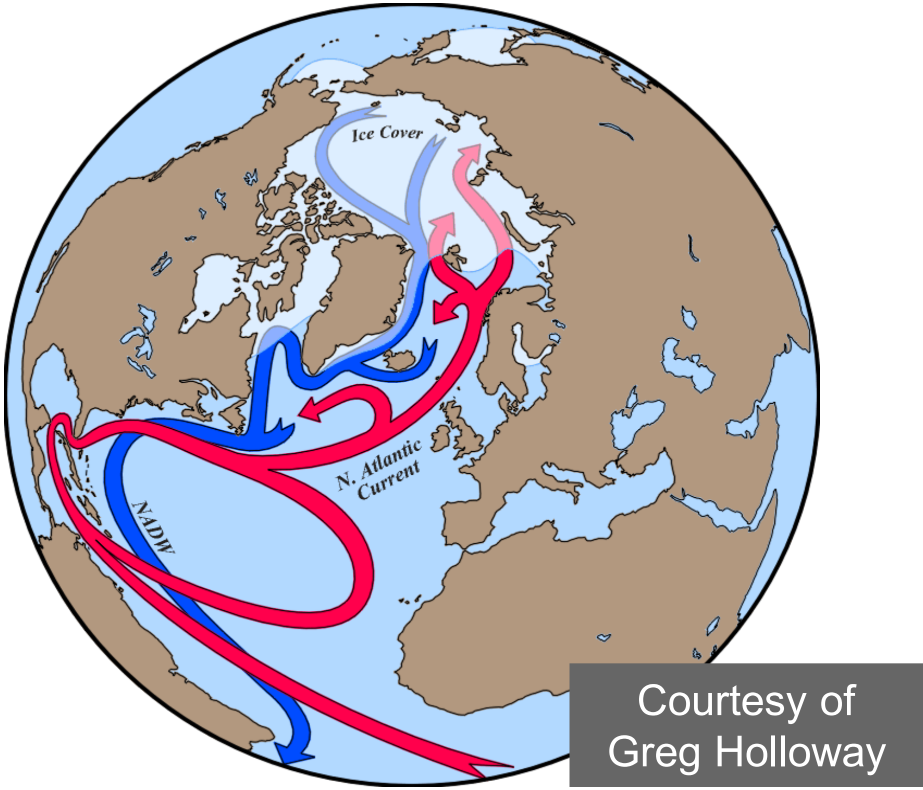
March Mean Surface albedo



- High albedo of sea ice modifies radiative fluxes
- Sea ice insulates ocean from atmosphere influencing turbulent heat & moisture exchange

Why sea ice matters: Hydrological Cycle





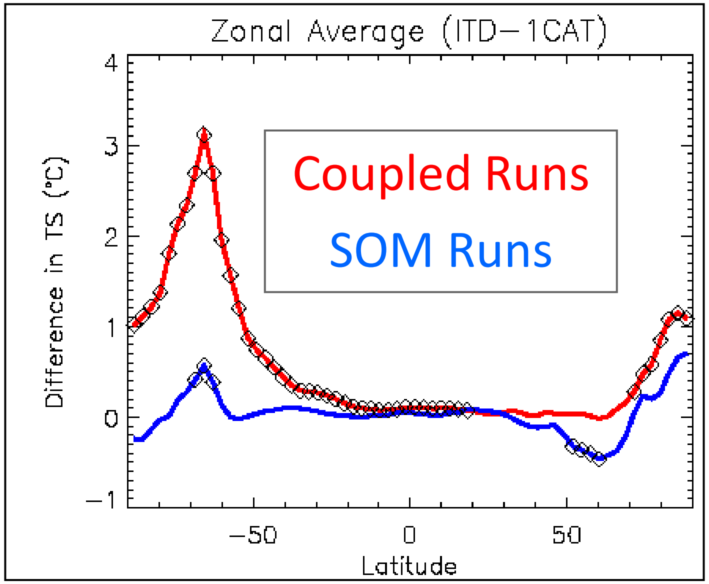
Courtesy of Greg Holloway

- Ice formation leads to salt flux to ocean and relatively fresh ice
- Ice melt releases freshwater back to the ocean
- Can modify ocean circulation

What do we need in a sea ice model for climate applications?

- Model which simulates a reasonable mean state/variability of sea ice
  - Concentration, thickness, mass budgets
- Realistically simulates ice-ocean-atmosphere exchanges of heat and moisture
- Realistically simulates response to climate perturbations - key climate feedbacks

Sea Ice models used in climate simulations



(Holland et al., 2006)

Two primary components

- Dynamics
  - Solves force balance to determine motion
- Thermodynamics
  - Solves for vertical ice temperature profile
  - Vertical/lateral melt and growth rates
- Ice Thickness Distribution (some models)
  - Sub-gridscale parameterization
  - Accounts for high spatial heterogeneity in ice

CICE (pronounced "sice"): The CICE Consortium Model

- CESM2 uses the CICE V5.1.2 (Hunke et al.)
  - Full documentation available online: http://www.cesm.ucar.edu/models/cesm2.0/sea-ice/
- Current CICE development is through the international CICE Consortium
  - https://github.com/CICE-Consortium/

Dynamics

Sea Ice Model - Dynamics

- Force balance between wind stress, water stress, internal ice stress, Coriolis and stress associated with sea surface slope
- Ice treated as a continuum with an effective large-scale rheology describing the relationship between stress and deformation
- Ice freely diverges (no tensile strength)
- Ice resists convergence and shear

(e.g. Hibler, 1979)

$$\underset{\substack{\uparrow\\ \text{Total derivative}}}{\mathrm{m}\frac{\mathbf{D}\mathbf{u}}{\mathrm{Dt}}} = \underset{\substack{\uparrow\\ \text{Coriolis}}}{-\mathrm{mf}\mathbf{k}\times\mathbf{u}} + \underset{\substack{\uparrow\\ \text{Air}\\ \text{stress}}}{\boldsymbol{\tau}_{\mathrm{a}}} + \underset{\substack{\uparrow\\ \text{Ocean}\\ \text{stress}}}{\boldsymbol{\tau}_{\mathrm{w}}} - \underset{\substack{\uparrow\\ \text{Sea Surface}\\ \text{Slope}}}{\mathrm{mg_r}\nabla \mathrm{Y}} + \underset{\substack{\uparrow\\ \text{Internal}\\ \text{Ice Stress}}}{\nabla\cdot\sigma}$$

Air-Ice and Ocean-Ice Stress

Air-Ice Stress

$$\vec{\tau}_a = \frac{\rho_a u^{*2}\vec{U}_a}{|\vec{U}_a|}, \qquad u^* = c_u\left|\vec{U}_a\right|$$

Ocean-Ice Stress

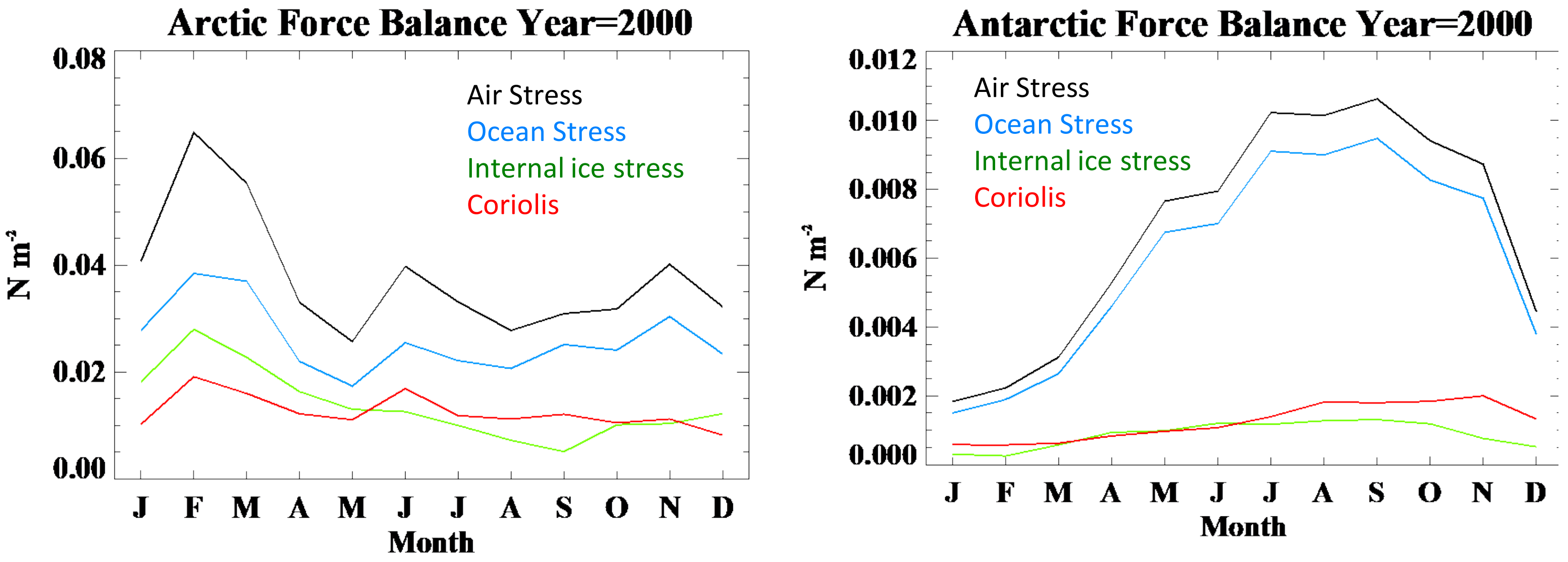
$$\vec{\tau}_w = c_w\rho_w\left|\vec{U}_w - \vec{u}\right|\left[\left(\vec{U}_w - \vec{u}\right)\cos\theta + \hat{k}\times\left(\vec{U}_w - \vec{u}\right)\sin\theta\right]$$

(e.g. Hibler, 1979)

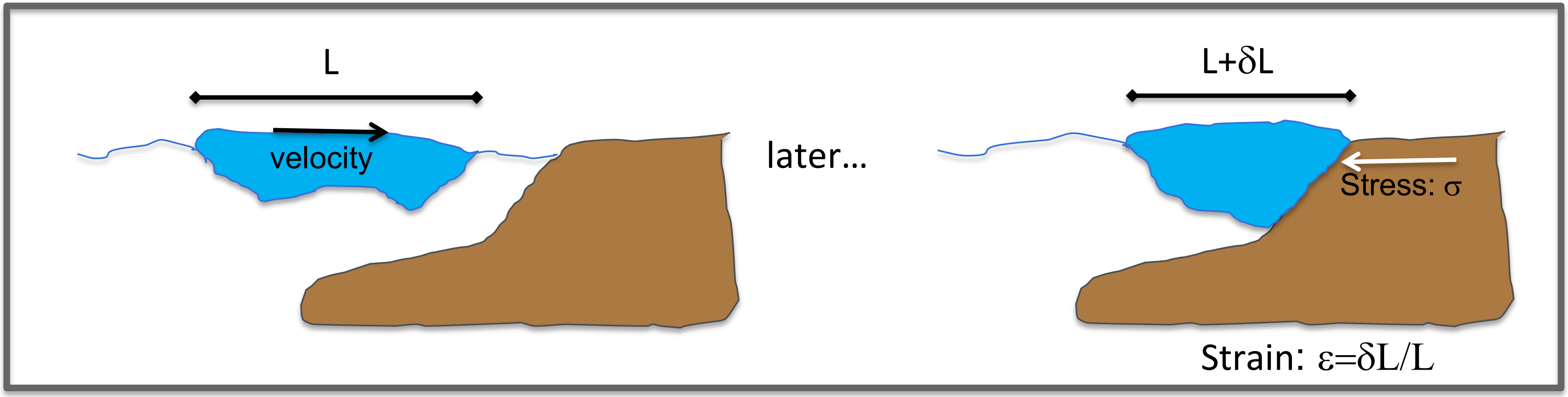
$$\underset{\substack{\uparrow\\\text{Total derivative}}}{m\frac{D\mathbf{u}}{Dt}} = \underset{\substack{\uparrow\\\text{Coriolis}}}{-mf\mathbf{k}\times\mathbf{u}} + \underset{\substack{\uparrow\\\text{Air}\\\text{stress}}}{\boldsymbol{\tau}_a} + \underset{\substack{\uparrow\\\text{Ocean}\\\text{stress}}}{\boldsymbol{\tau}_w} - \underset{\substack{\uparrow\\\text{Sea Surface}\\\text{Slope}}}{mg_r\nabla Y} + \underset{\substack{\uparrow\\\text{Internal}\\\text{Ice Stress}}}{\nabla\cdot\sigma}$$

Simulated Force Balance

- Arctic: Air stress largely balanced by ocean stress. Internal ice stress has smaller role
- Antarctic: Ice in nearly free drift - weak internal ice stress



Internal Ice stress



- Stress causes ice to deform, but volume is conserved.
- Need to relate ice stress (σ) to ice strain rate (ε) → area of active research.

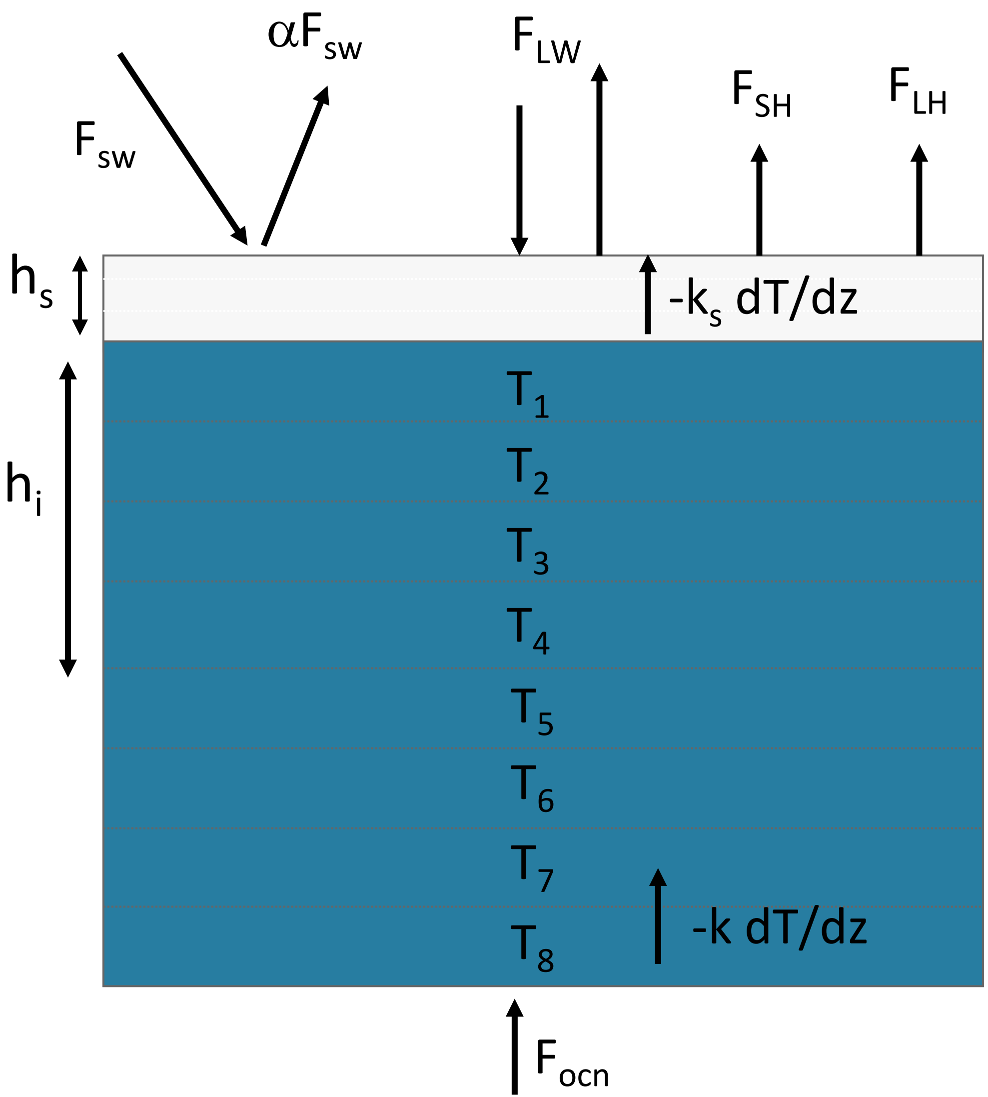
(e.g. Hibler, 1979)

$$m\frac{Du}{Dt} = -mf\mathbf{k}\times\mathbf{u} + \boldsymbol{\tau}_a + \boldsymbol{\tau}_w - mg_r\nabla Y + \boxed{\nabla\cdot\sigma}$$

Total derivative ↑ Coriolis ↑ Air stress ↑ Ocean stress ↑ Sea Surface Slope ↑ Internal Ice Stress

Thermodynamics

Sea ice thermodynamics



- Calculate top and basal growth/melt
- CESM 2: 8 sea ice thickness categories and 3 snow layers. (CESM1: 4 and 1 respectively)

Top surface flux balance

$$(1-\alpha)F_{SW} + F_{LW} - \sigma T^4 + F_{SH} + F_{LH} + k\frac{\partial T}{\partial z} = -q\frac{dh}{dt}$$

Vertical heat transfer (conduction)

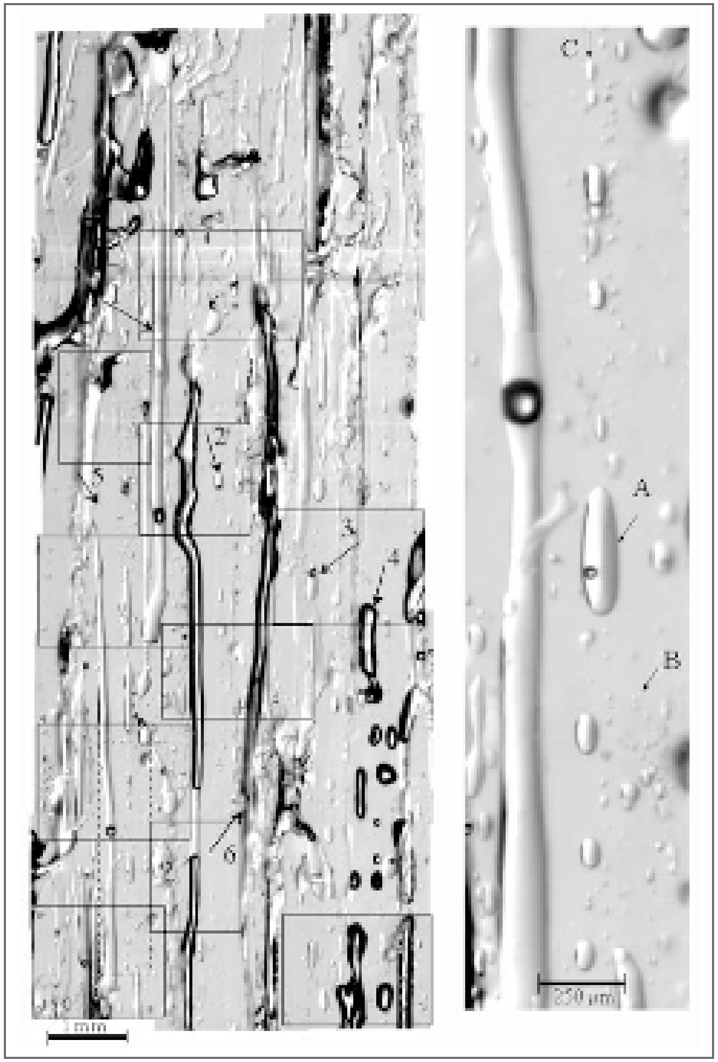
$$\rho c\frac{\partial T}{\partial t} = \frac{\partial}{\partial z}k\frac{\partial T}{\partial z} + Q_{SW}$$

SEA ICE

Bottom surface flux balance

$$F_{ocn} - k\frac{\partial T}{\partial z} = -q\frac{dh}{dt}$$

Thermodynamics: Vertical Heat Transfer



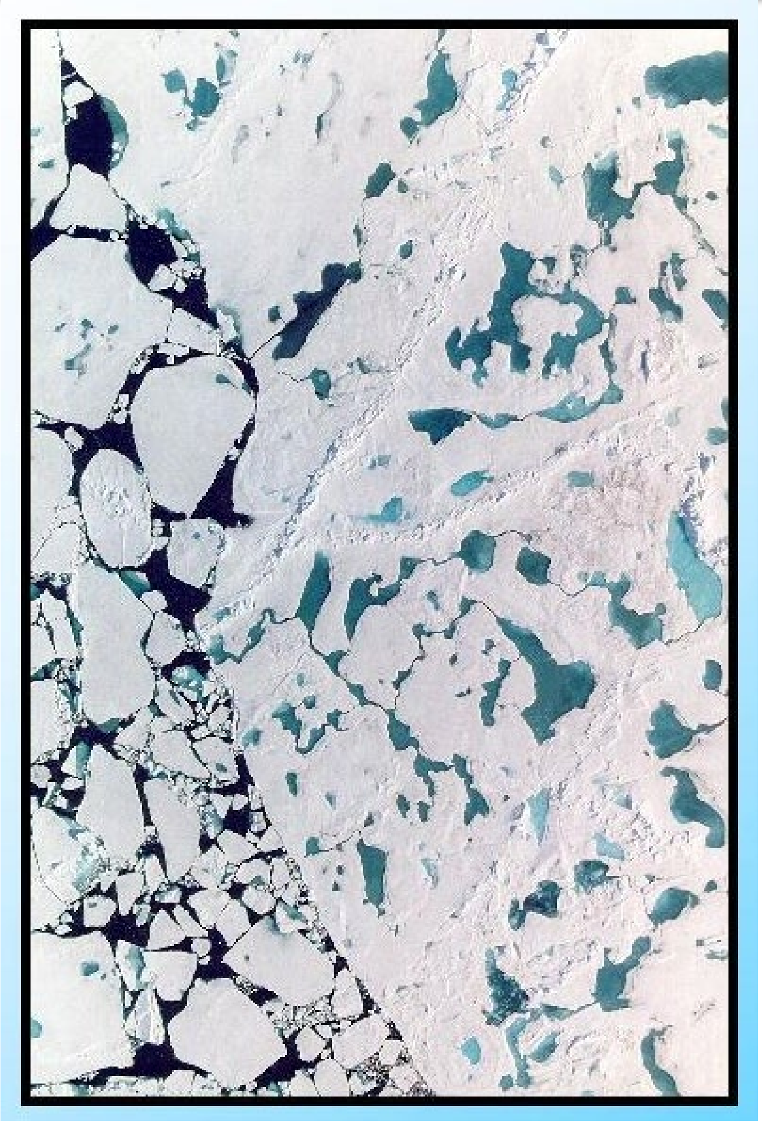
(from Light, Maykut, Grenfell, 2003)

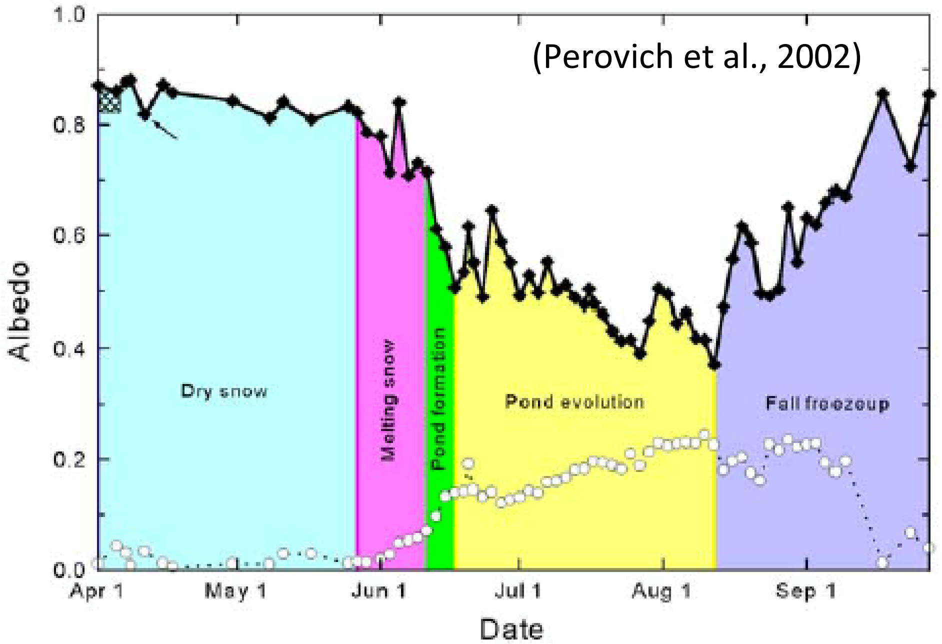
$$\rho c \frac{\partial T}{\partial t} = \frac{\partial}{\partial z} k \frac{\partial T}{\partial z} + Q_{SW}$$

SEA ICE

- Heat capacity and conductivity are functions of T/S of ice
- Solve to get temperature **and** salinity profiles using mushy layer thermodynamics (Turner and Hunke 2015; new in CESM2)
- Assume pockets/channels are brine filled and they are in thermal equilibrium with ice
- Assume non-varying ice density

Albedo



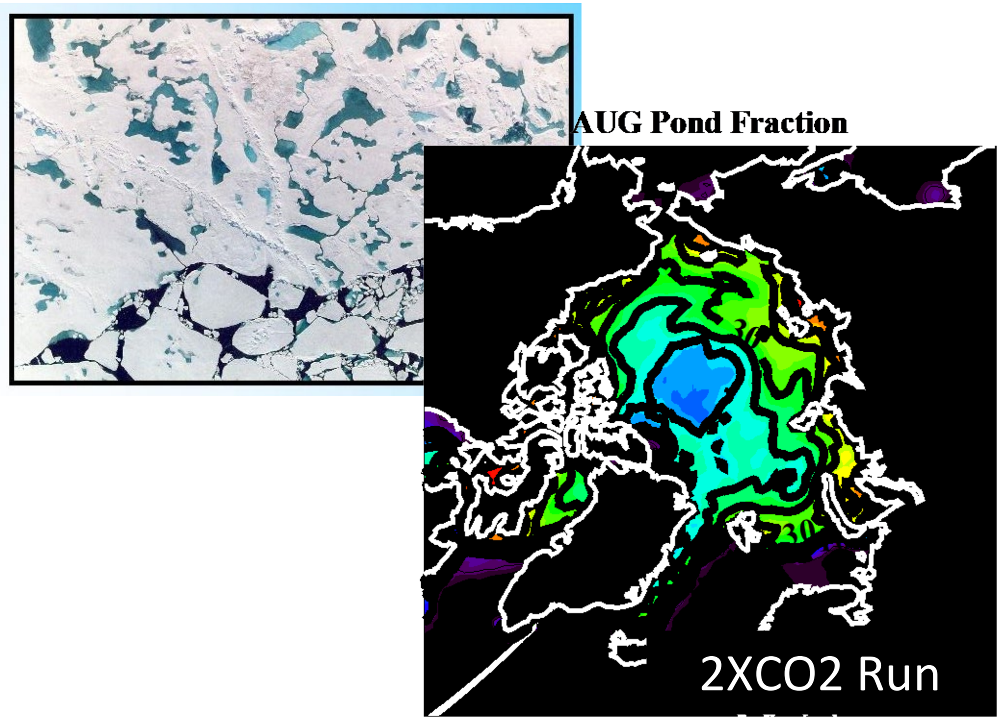


Often the parameterized sea ice albedo depends on characteristics of surface state (snow, temp, ponding, h_i).

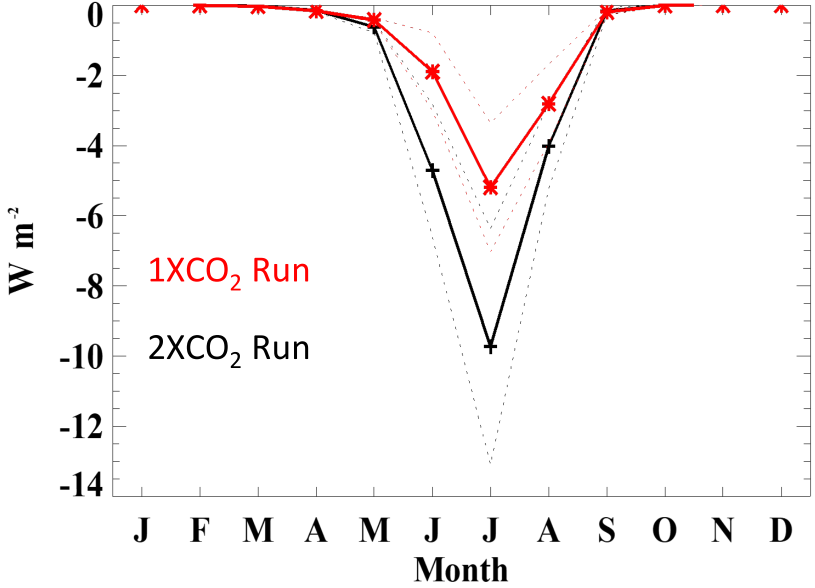
Surface ice albedo is only for fraction of gridcell covered by ice.

Melt Pond Parameterization

- Only influences radiation and has big influence on surface forcing
- Ponds evolve over time and are carried as tracers on the ice
- CESM2 pond evolution takes into account if sea ice is deformed (level ponds)

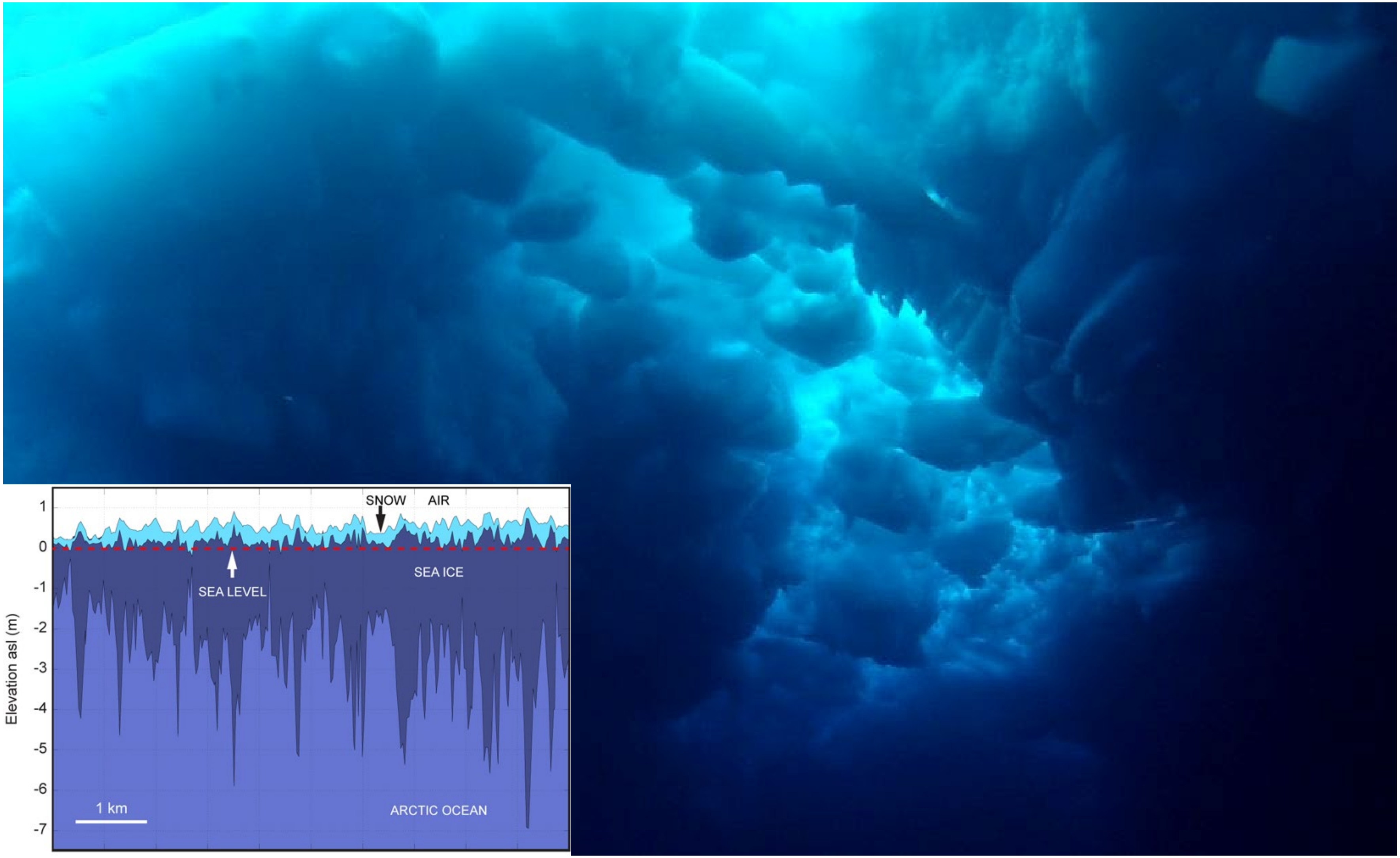


Arctic Basin Average SW flux (pond – nopond)



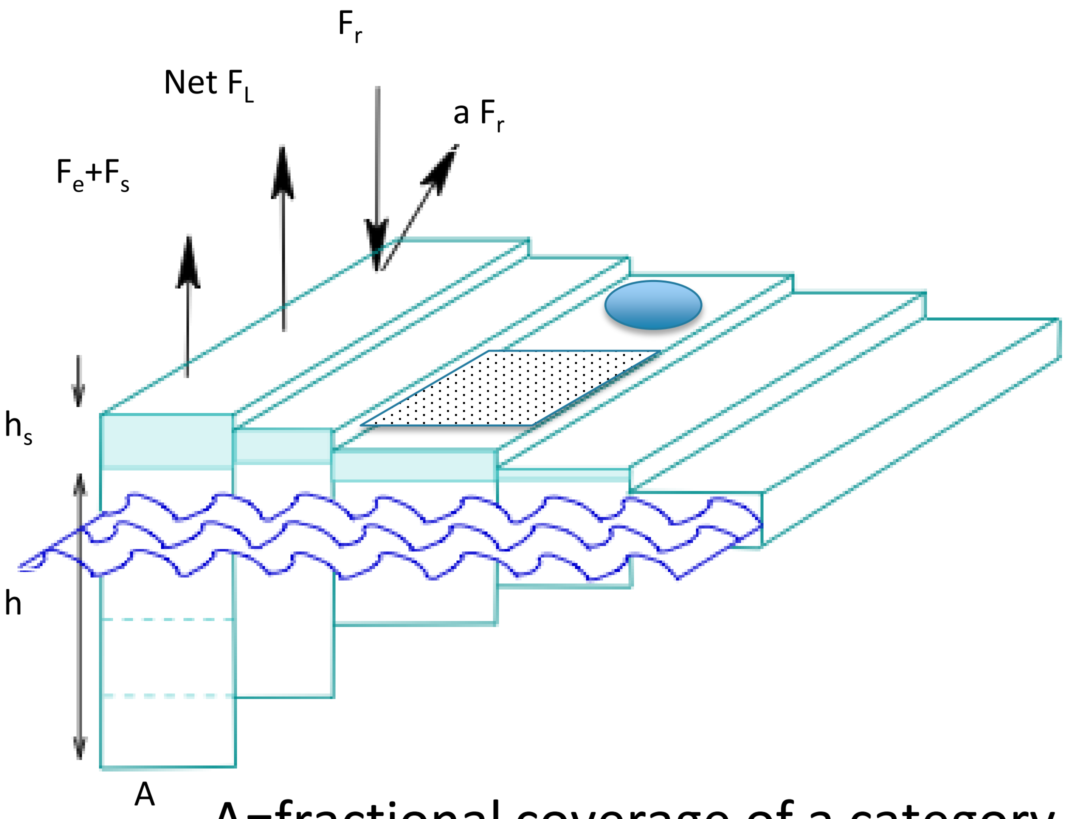
Holland, M. M., et al. 2012: Improved sea ice shortwave radiation physics in CCSM4

Ice Thickness Distribution



Ice Thickness Distribution

- Represents high spatial heterogeneity of sea ice
- CESM uses five ice “categories”



A=fractional coverage of a category

For each category, keep track of:

- Fractional area per grid cell
- Volume per grid cell
- Enthalpy per grid cell
- Surface temperature
- Snow and melt pond areas
- Aerosol contents
- Etc.

Ice Thickness Distribution

Ice thickness distribution g(x,y,h,t) evolution equation from Thorndike et al. (1975)

$$\frac{\partial g}{\partial t} = -\frac{\partial}{\partial h}(fg) + L(g) - \nabla \bullet (\vec{v}g) + \Psi(h,g,\vec{v})$$

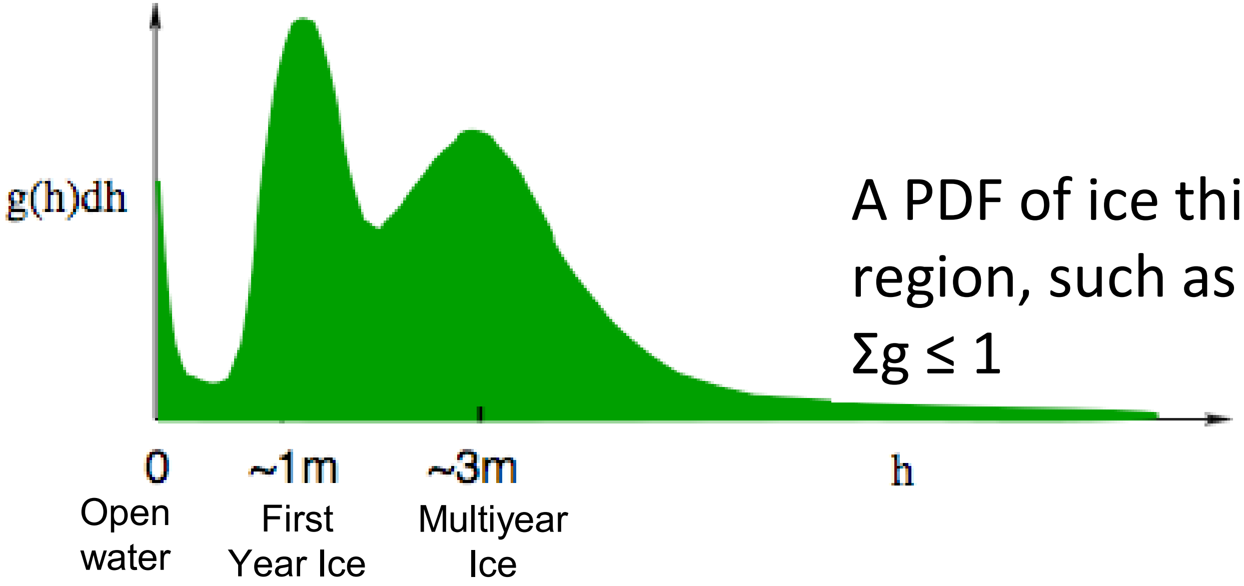
Ice Growth

Lateral Melt

Convergence

Mechanical Redistribution

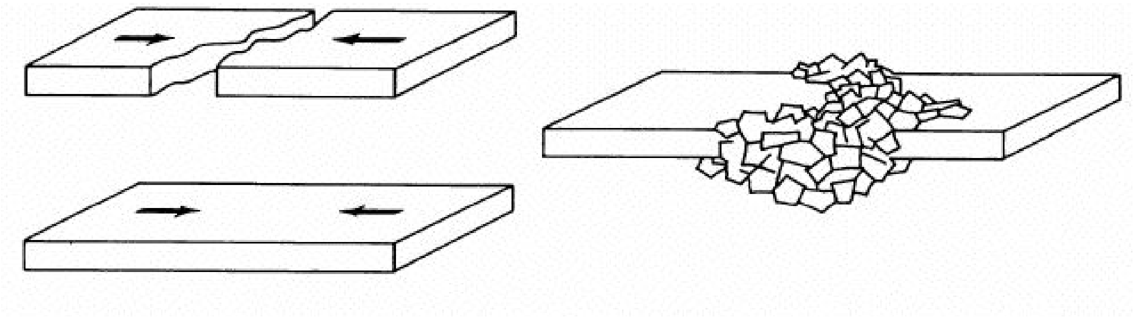
g(h)dh is the fractional area covered by ice of thickness h to h+dh

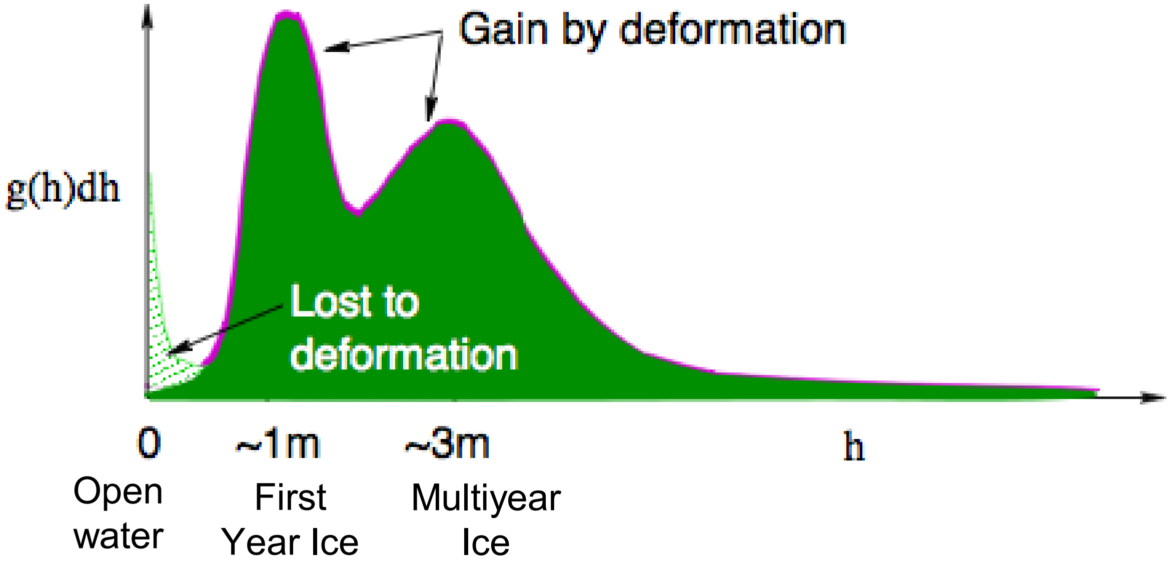


A PDF of ice thickness h in a region, such as a grid cell. $0 \leq \Sigma g \leq 1$

Ice Thickness Distribution: impact of convergence

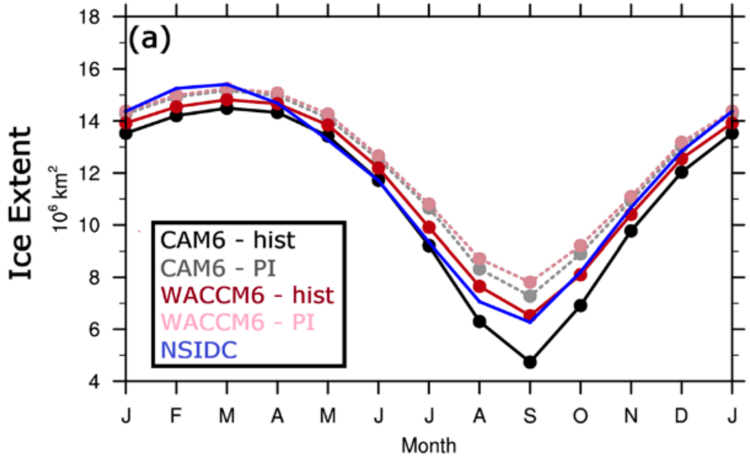
Mechanical redistribution: Transfer ice from thin part of distribution to thicker categories

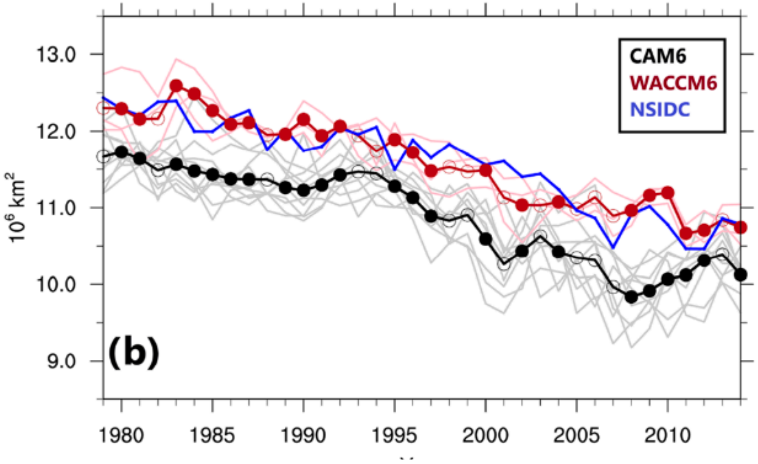


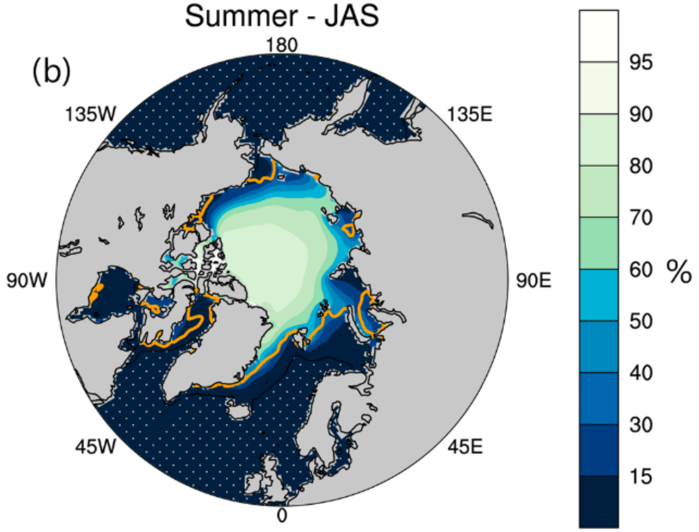


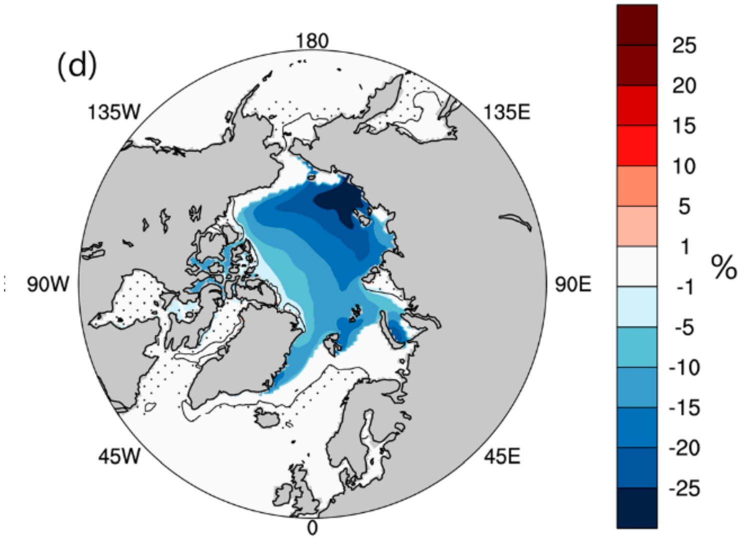
CESM2 Science Highlights

CESM2 Historical (1979-2014) Arctic Sea Ice Extent



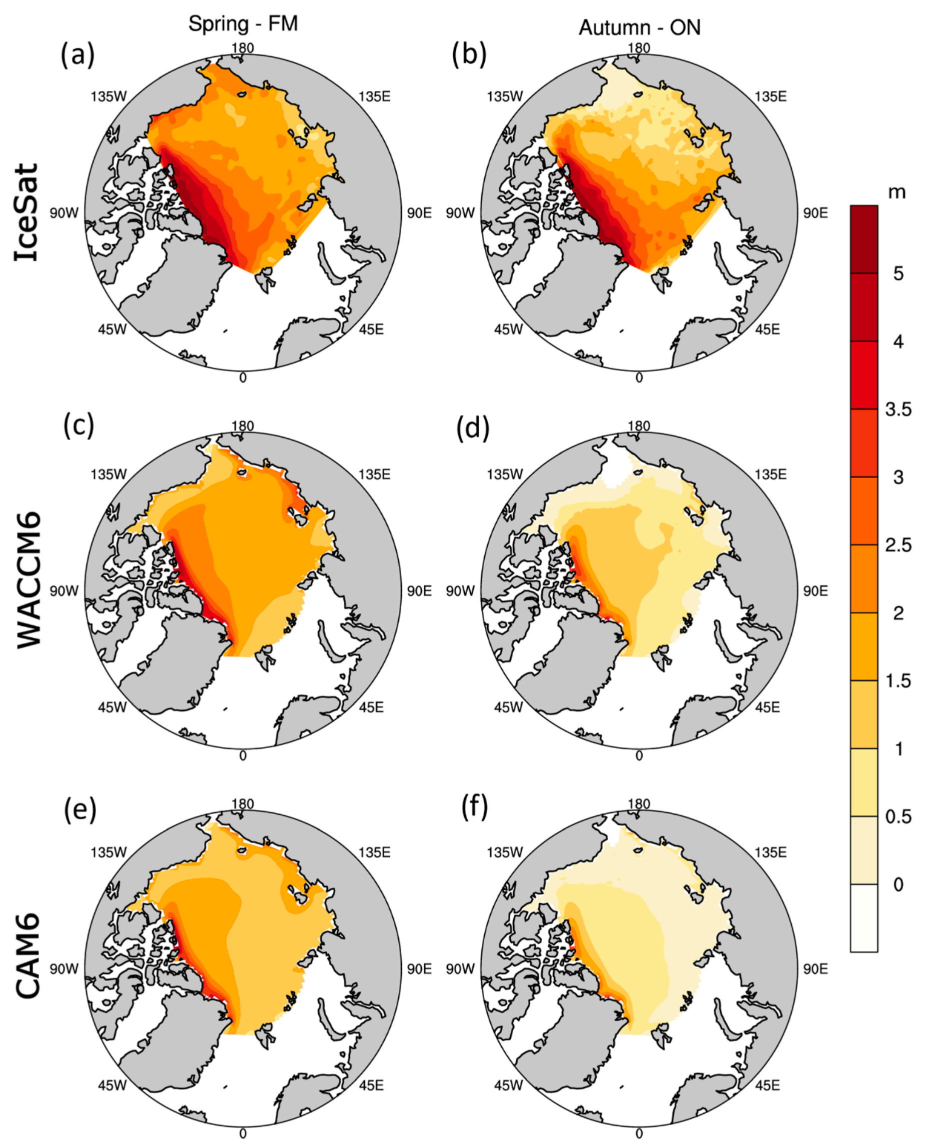






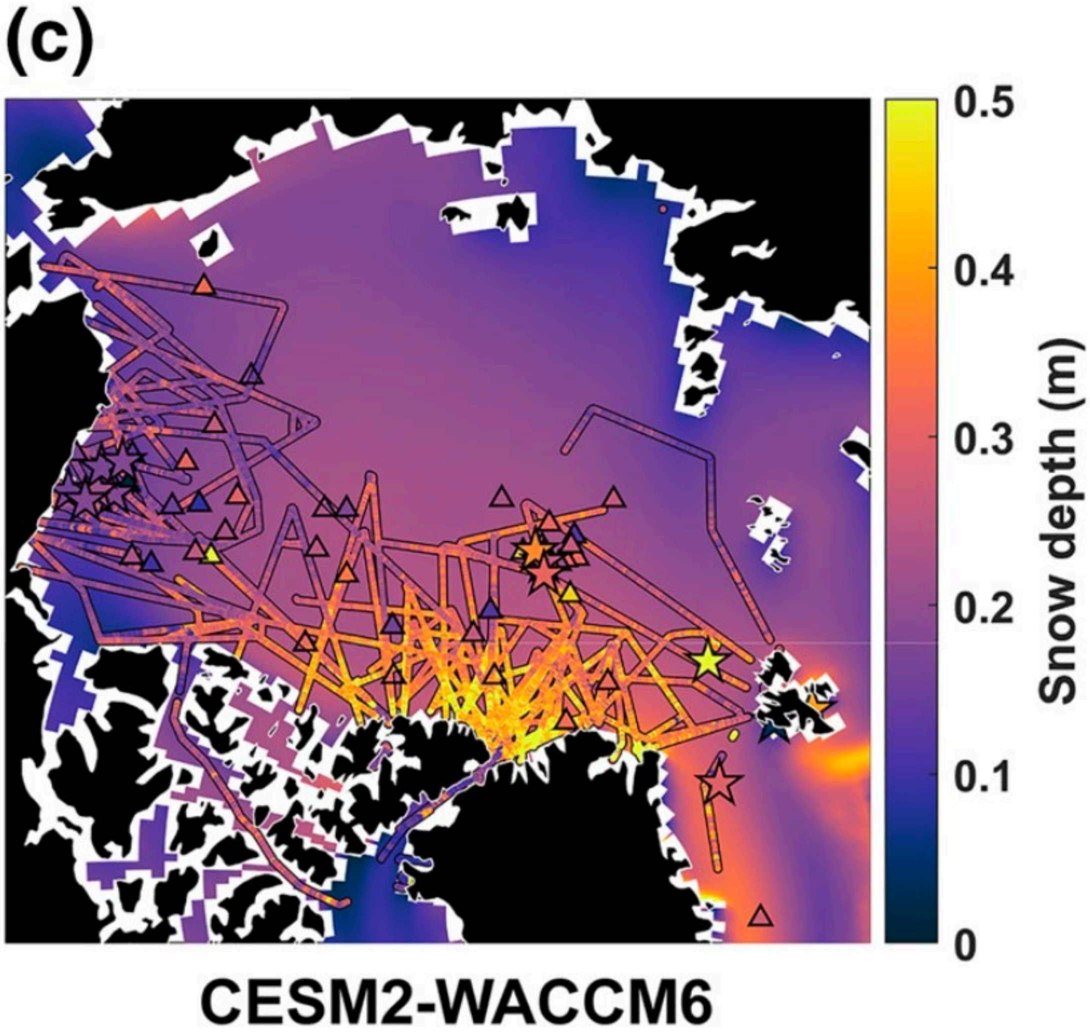
DuVivier et al. 2020

CESM2 Historical (2001-2005) Arctic Sea Ice Thickness



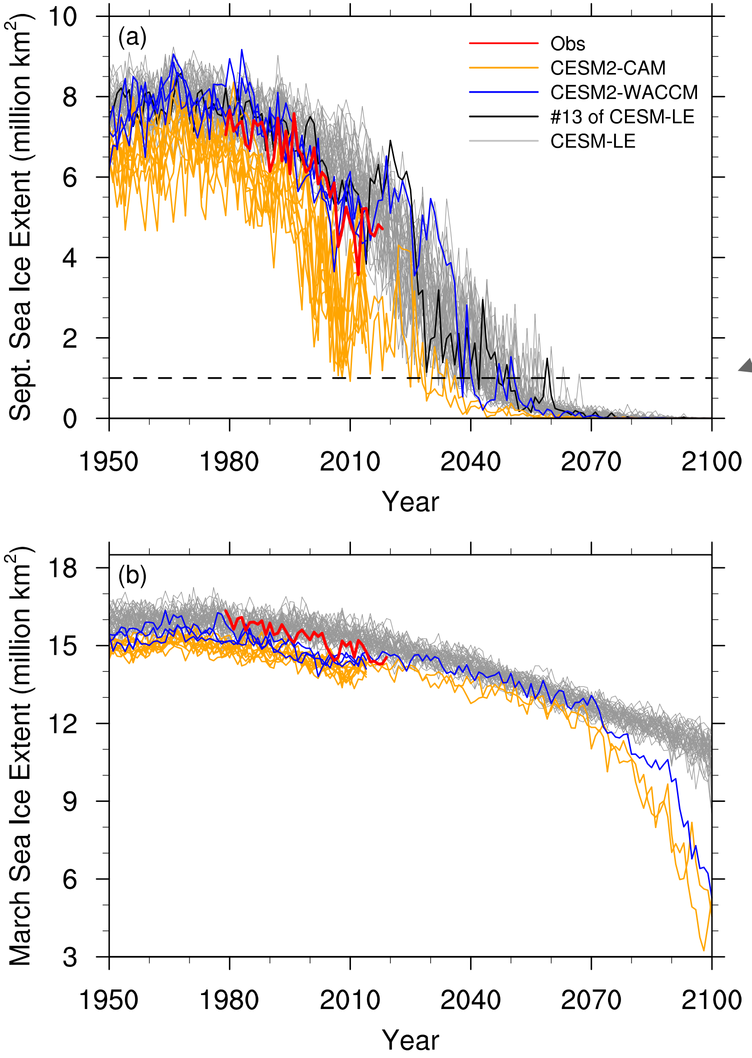
DuVivier et al. 2020

CESM2 Historical Arctic Sea Ice Snow



Webster et al. 2020

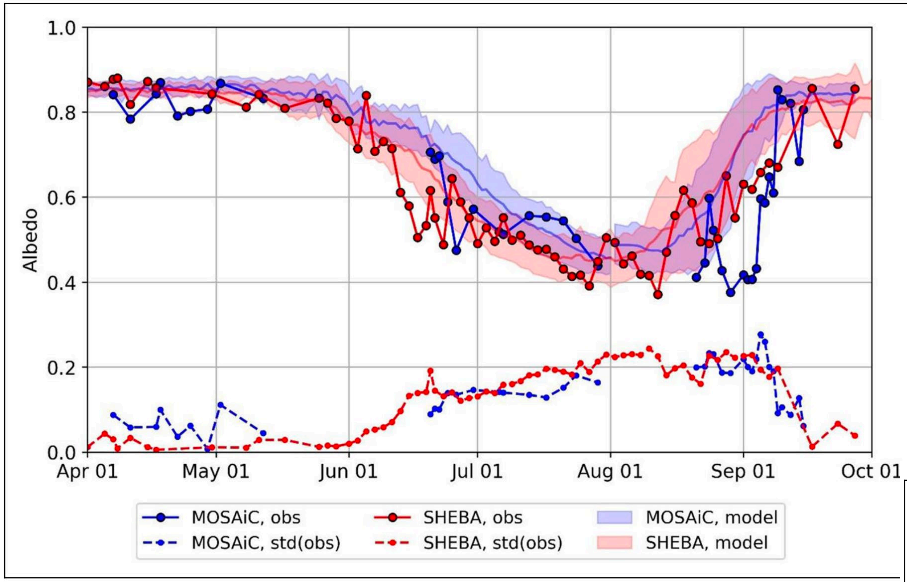
CESM2 Arctic Sea Ice Extent Projections



"Ice Free" Arctic

DeRepentigny et al. 2022

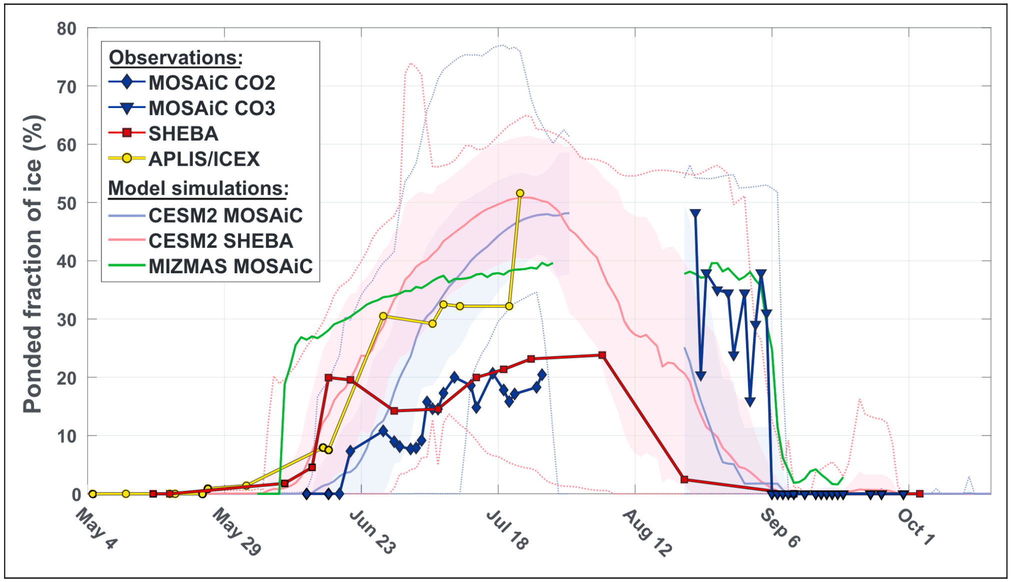
CESM2 Arctic Sea Ice – Melt ponds



Light et al. 2022

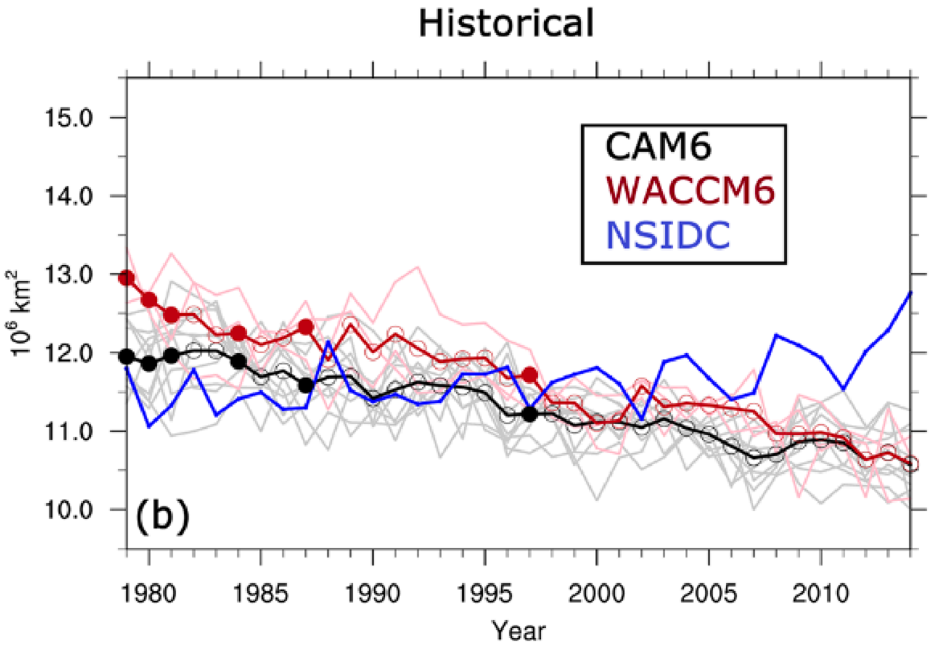
Ponds are too extensive, but not deep enough.

→ Compensating biases lead to reasonable albedo.



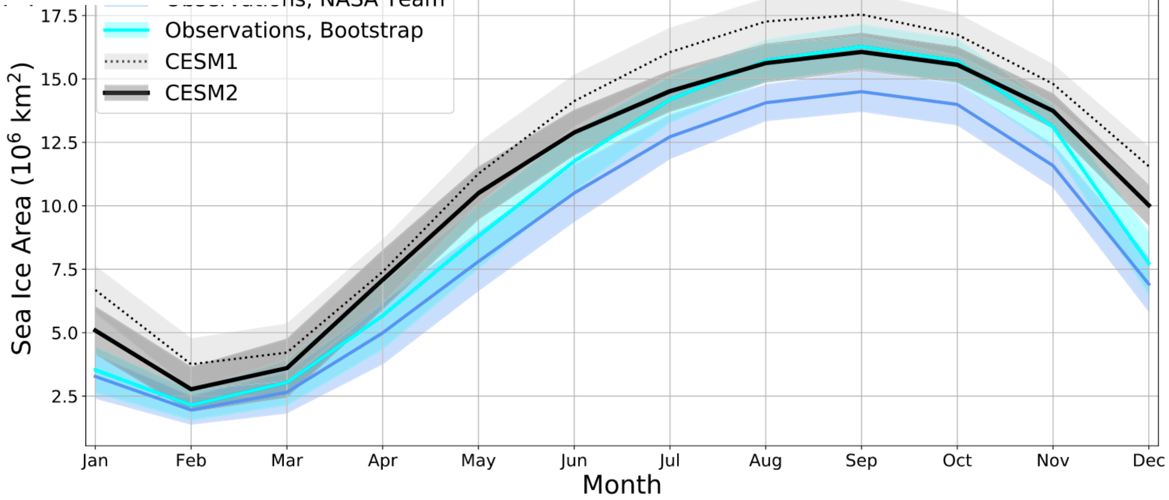
Webster et al. 2022

CESM2 Antarctic Sea Ice

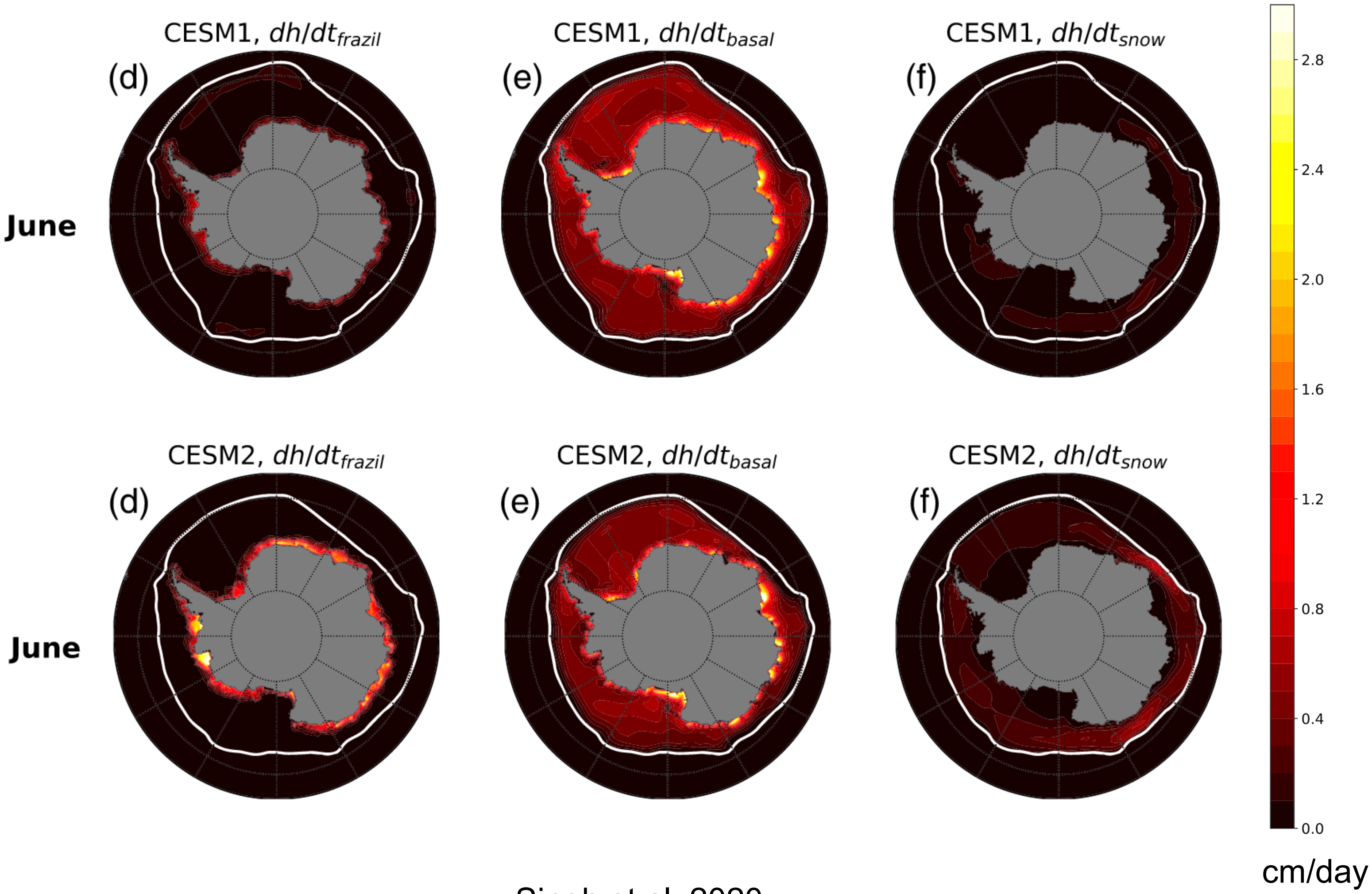


DuVivier et al. 2020

Singh et al. 2020



CESM2 Antarctic Sea Ice – growth processes are changing



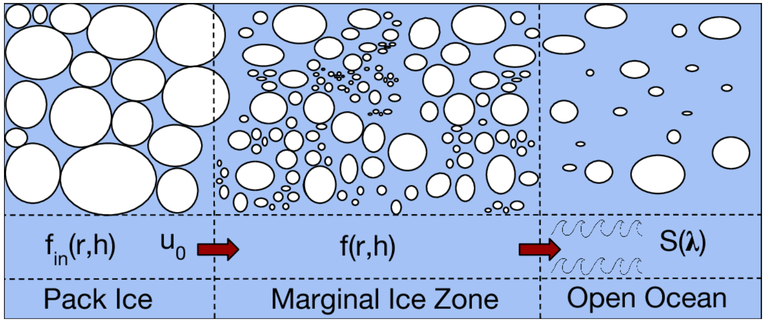
Singh et al. 2020

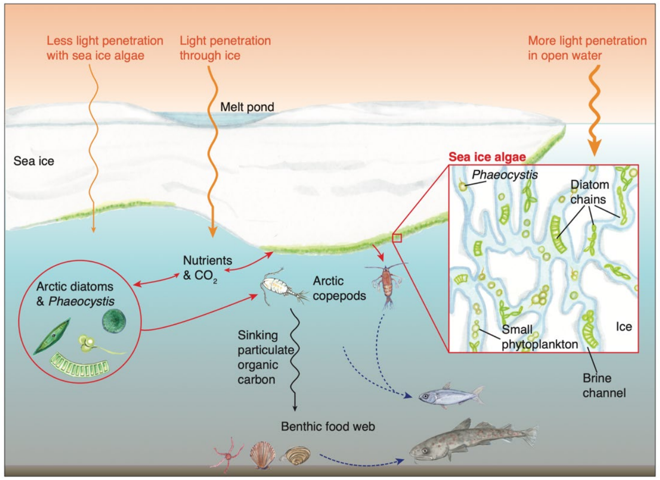
Summary

- CICE in CESM2
  - EVP dynamics
  - Sophisticated mushy layer thermodynamics (Turner and Hunke 2015)
  - 8 sea ice vertical levels; 3 snow vertical levels
  - Sub-gridscale ice thickness distribution – 5 categories
  - Level ice ponds (Hunke et al. 2013)
  - Salinity dependent freezing point

In development for CESM3 and beyond…

- Wave-ice interactions for floe size distribution
- Sea ice biogeochemistry
- Snow evolution physics
- Melt pond physics
- And more!





Thank You

Questions?

dbailey@ucar.edu

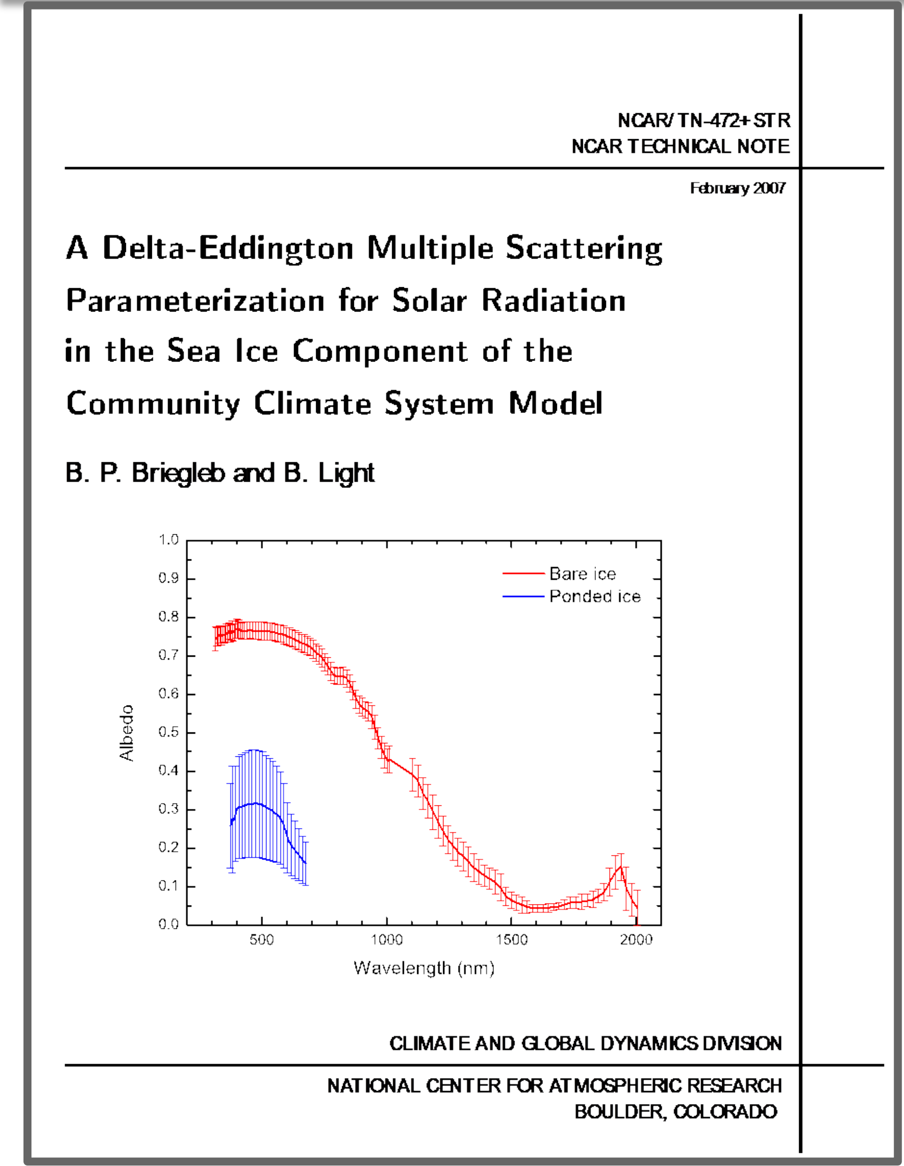
http://assets.nydailynews.com/polopoly_fs/1.17149.1313675705!/img/httpImage/image.jpg_gen/derivatives/gallery_635/gal-natures-best-08-jpg.jpg

Sea Ice models used in climate simulations

Two primary components

- Dynamics
  - Solves force balance to determine sea ice motion
- Thermodynamics
  - Solves for vertical ice temperature profile
  - Vertical/lateral melt and growth rates

Delta Eddington Solar Radiation parameterization



NCAR/TN-472+STR
NCAR TECHNICAL NOTE

February 2007

A Delta-Eddington Multiple Scattering Parameterization for Solar Radiation in the Sea Ice Component of the Community Climate System Model

B. P. Briegleb and B. Light

CLIMATE AND GLOBAL DYNAMICS DIVISION

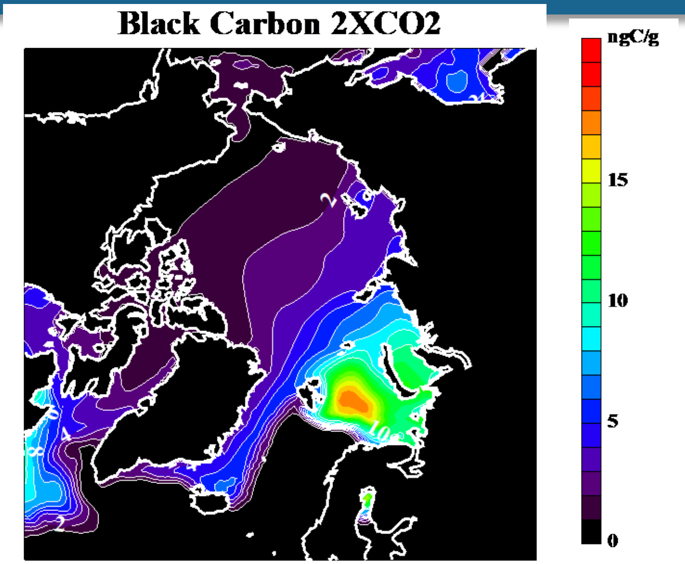
NATIONAL CENTER FOR ATMOSPHERIC RESEARCH
BOULDER, COLORADO

- Inherent optical properties define scattering and absorption properties for snow, sea ice, and absorbers.
- Calculate base albedo and then modify.
- Explicitly allows for included absorbers (e.g. algae, carbon, sediment) in sea ice
- Accounts for melt ponds, snow grain sizes, etc.
- Used in CESM1 <u>and</u> CESM2

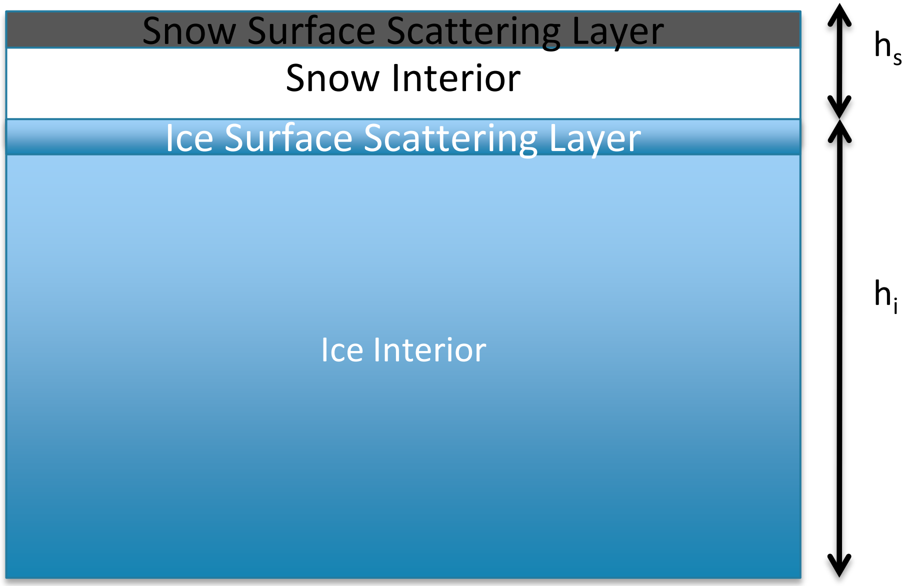
Aerosol deposition and cycling

- Aerosol (e.g. dust, black carbon) deposition and cycling now included.
- These are deposited from the atmosphere and modified by melt and transport
- ~10% of the impact of melt ponds

With 1850 Aerosol Deposition



Arctic Basin Average



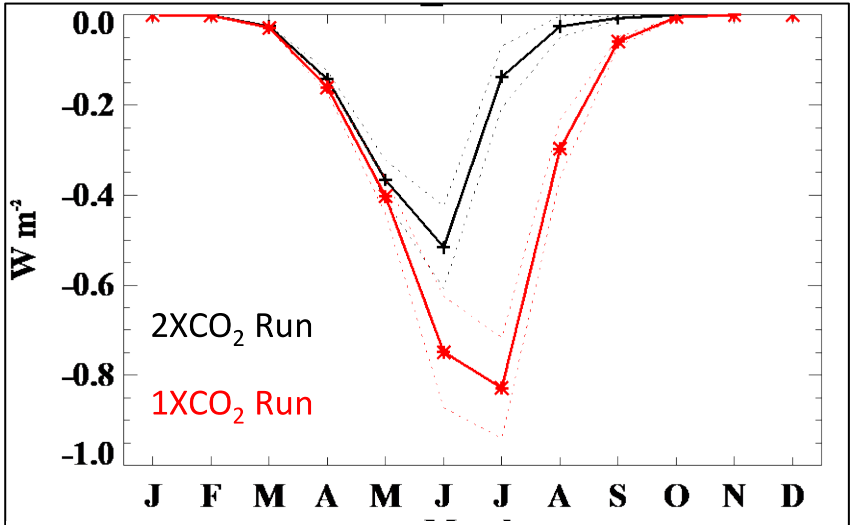
Snow Surface Scattering Layer

Snow Interior

Ice Surface Scattering Layer

Ice Interior

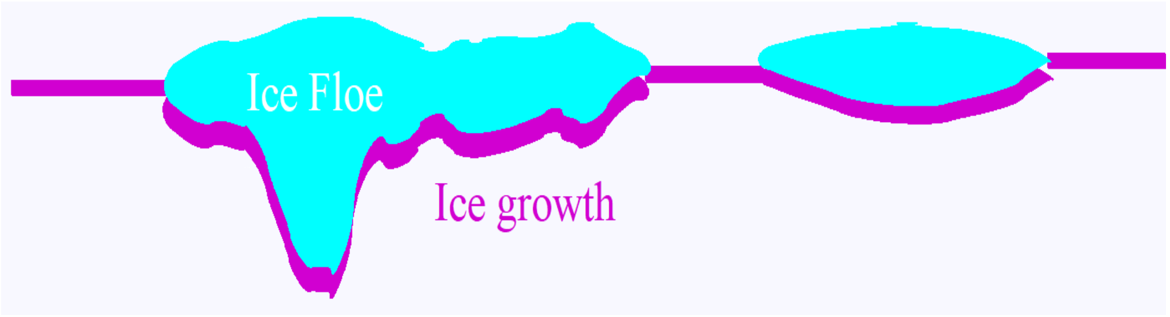
Arctic Basin Average SW flux (aero – noaero)

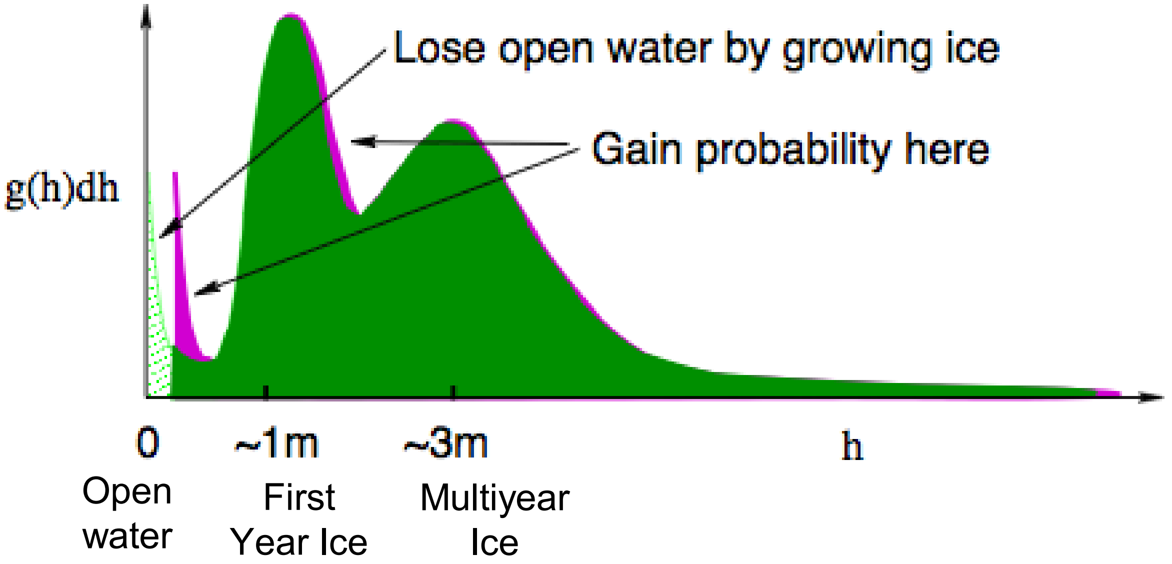


Holland, M. M., et al. 2012: Improved sea ice shortwave radiation physics in CCSM4

Ice Thickness Distribution: impact of ice growth

Lose open water, gain probability of both thin ice and thicker ice





Dynamics used in CESM

CESM uses Elastic Viscous Plastic Model (Hunke and Dukowicz, 1997)

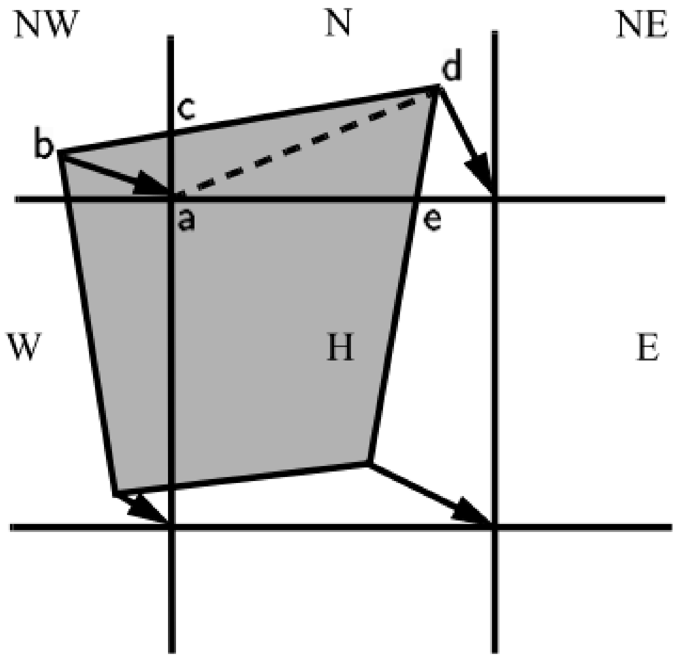
- Ice has no tensile strength but resists convergence and shear with strength dependent on ice state.
- Treats ice as a continuum, based on Viscous-Plastic Rheology (Hibler, 1979)

  →Plastic at normal strain rates and viscous at very small strain rates.

  →A viscous-plastic material creeps along but responds to stresses and strains.
- EVP adds in non-physical elasticity as numerical device for solving equations.

Advection

Would make so many state variables prohibitive, if it weren't for remapping by Lipscomb and Hunke 2004.



Conserved quantities are remapped from the shaded "departure region", which is computed from backward trajectories of the ice motion field.

Assessing Sea Ice Mass Budgets

- Equilibrium Ice Thickness Reached when
  - Ice growth is balanced by ice melt + ice divergence
  - Illustrative to consider how different models achieve this balance and how mass budgets change over time

$$\frac{d\bar{h}}{dt} = \Gamma_h - \nabla \bullet (\vec{u}h)$$

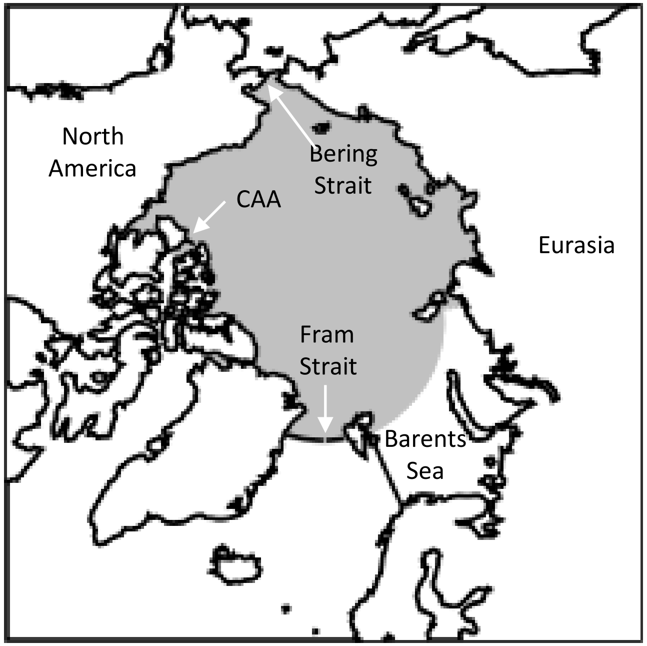
Ice volume change

Thermodynamic source

Divergence

Climate model archive of monthly averaged ice thickness and velocity

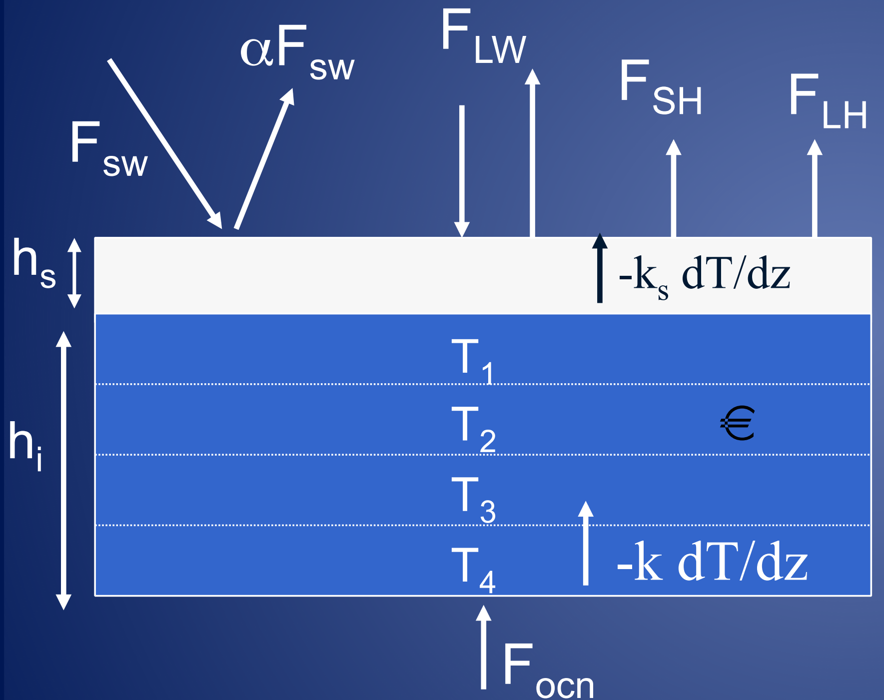
Assess Arctic ice volume, transport through Arctic straits, and solve for ice growth/melt as residual



Holland et al., 2010

Sea ice loss is modified by climate feedbacks

- Fundamental sea ice thermodynamics gives rise to a number of important feedbacks



Balance of fluxes at surface

$$\boxed{(1-\alpha)F_{SW}} + F_{LW} - \sigma T^4 + F_{SH} + F_{LH} + k\frac{\partial T}{\partial z} = -q\frac{dh}{dt}$$

Vertical heat transfer (conduction, SW absorption)

Balance of fluxes at ice base

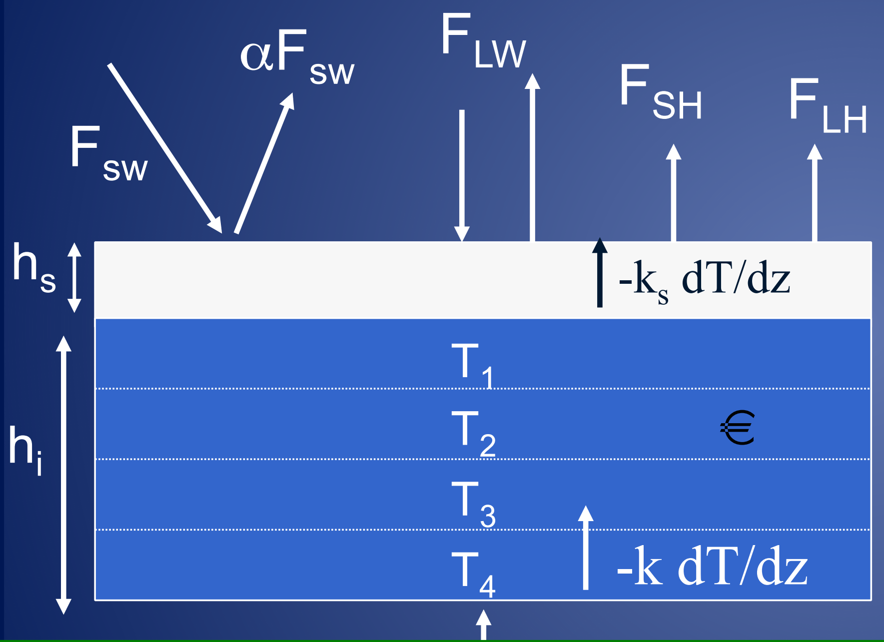
$$\boxed{F_{ocn}} - k\frac{\partial T}{\partial z} = -q\frac{dh}{dt}$$

Surface albedo changes modify SW absorption in ice and ocean heat flux

Ice loss lowers albedo – positive feedback

Ice mass budgets affected by climate feedbacks

- Fundamental sea ice thermodynamics gives rise to a number of important feedbacks



Balance of fluxes at surface

$$(1-\alpha)F_{SW} + F_{LW} - \sigma T^4 + F_{SH} + F_{LH} + k\frac{\partial T}{\partial z} = -q\frac{dh}{dt}$$

Vertical heat transfer (conduction, SW absorption)

Balance of fluxes at ice base

$$F_{ocn} - k\frac{\partial T}{\partial z} = -q\frac{dh}{dt}$$

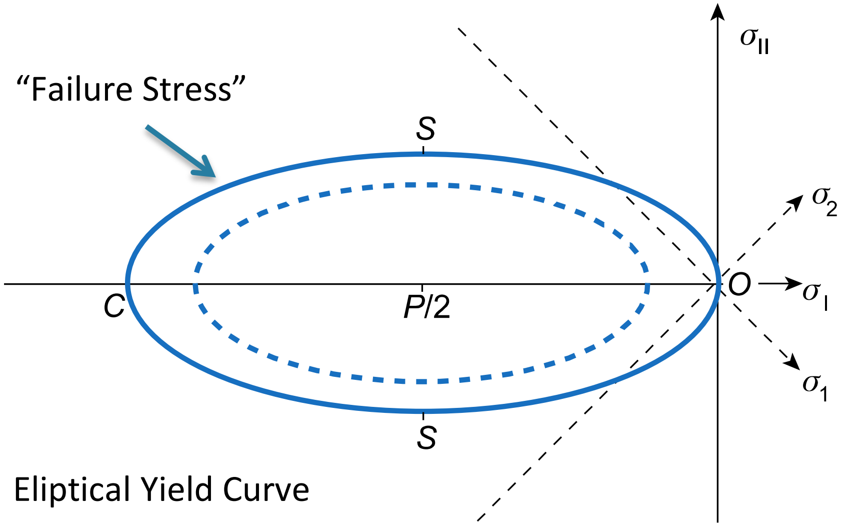
Heat conduction related to vertical temperature gradient

Causes ice growth to vary as 1/h

Has a stabilizing effect on ice thickness since thin ice grows more rapidly

Sea Ice Model - Dynamics

- Internal Ice Stress
  - Use variant of Viscous-Plastic Rheology (Hibler, 1979)
  - Treats ice as a continuum - plastic at normal strain rates and viscous at very small strain rates.
  - Ice has no tensile strength (freely diverges) but resists convergence and shear (strength dependent on ice state)



Eliptical Yield Curve

Elastic-Viscous-Plastic Model

EVP model uses explicit time stepping by adding elastic waves to constitutive law (Hunke and Dukowicz, 1997)

Ice Thickness Distribution

Ice thickness distribution g(x,y,h,t) evolution equation from Thorndike et al. (1975)

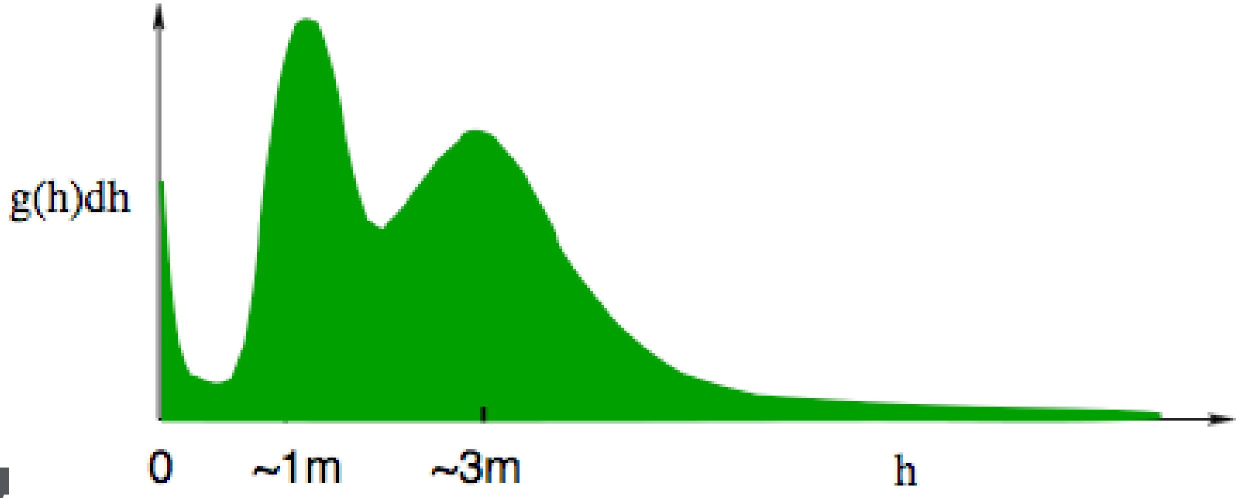
$$\frac{\partial g}{\partial t} = -\frac{\partial}{\partial h}(fg) + L(g) - \nabla \bullet (\vec{v}g) + \Psi(h,g,\vec{v})$$

Ice Growth

Lateral Melt

Convergence

Mechanical Redistribution

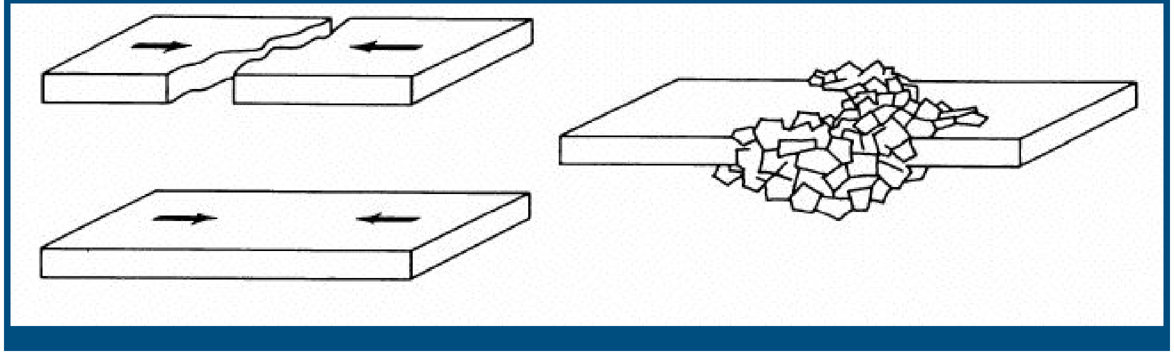


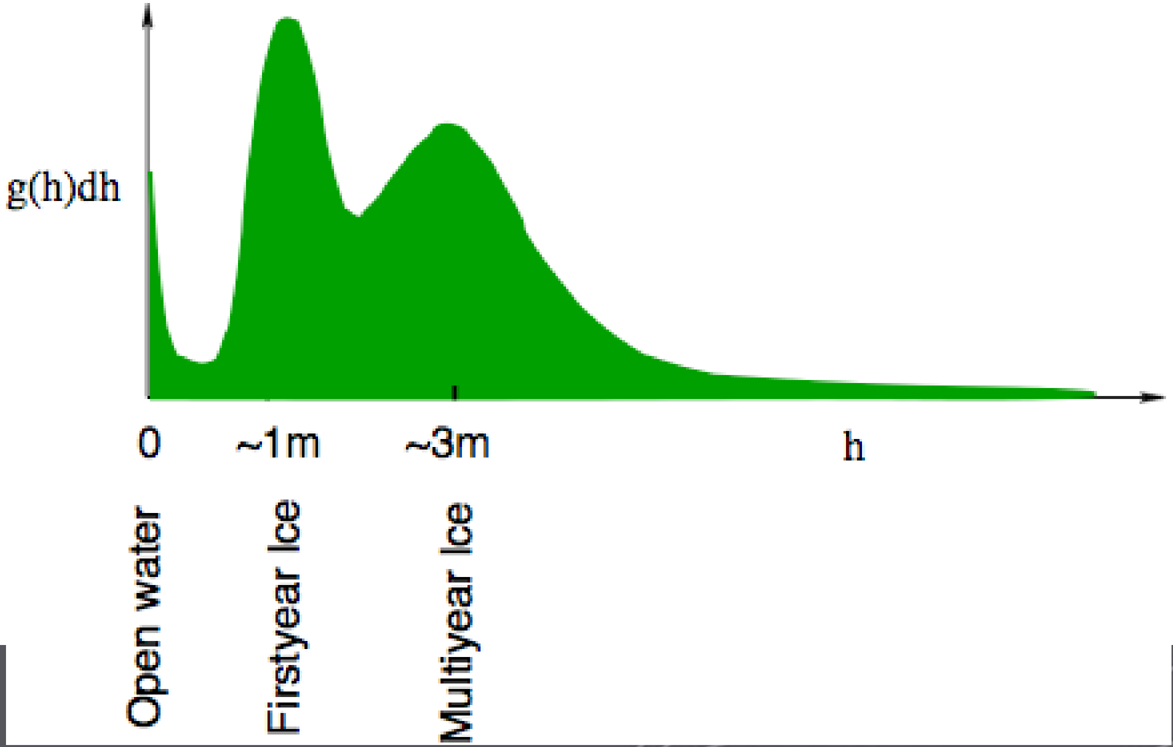
g(h)dh is the fractional area covered by ice of thickness h to h+dh

Y = Mechanical redistribution

Transfers ice from thin part of distribution to thicker categories

Converging hypothetical floes





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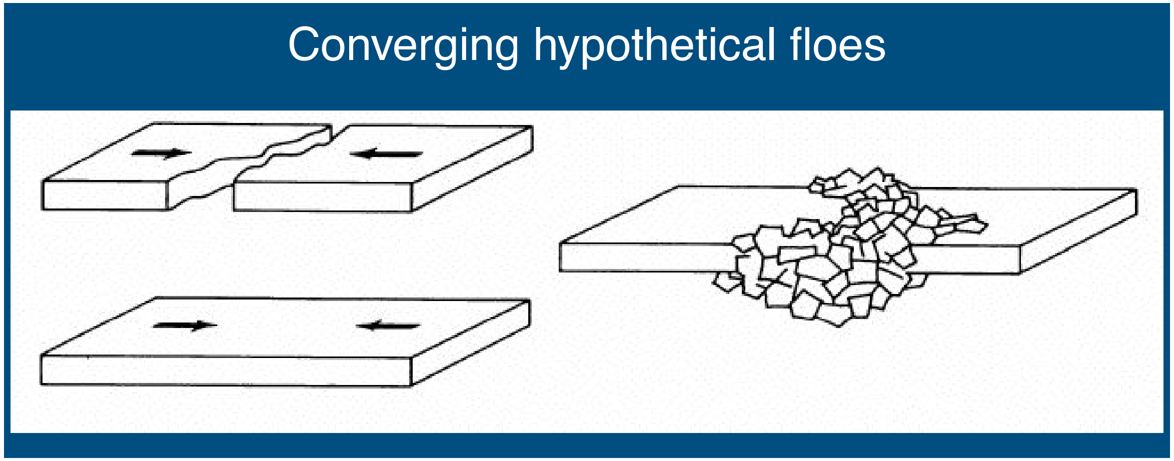
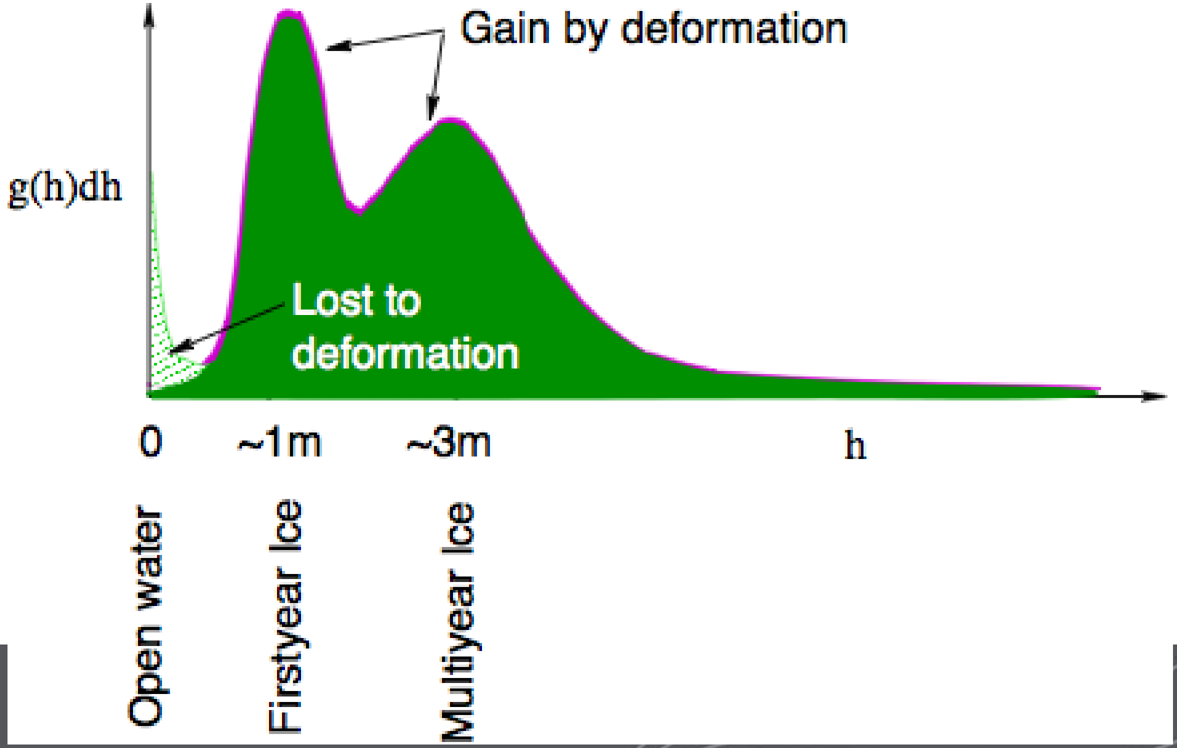
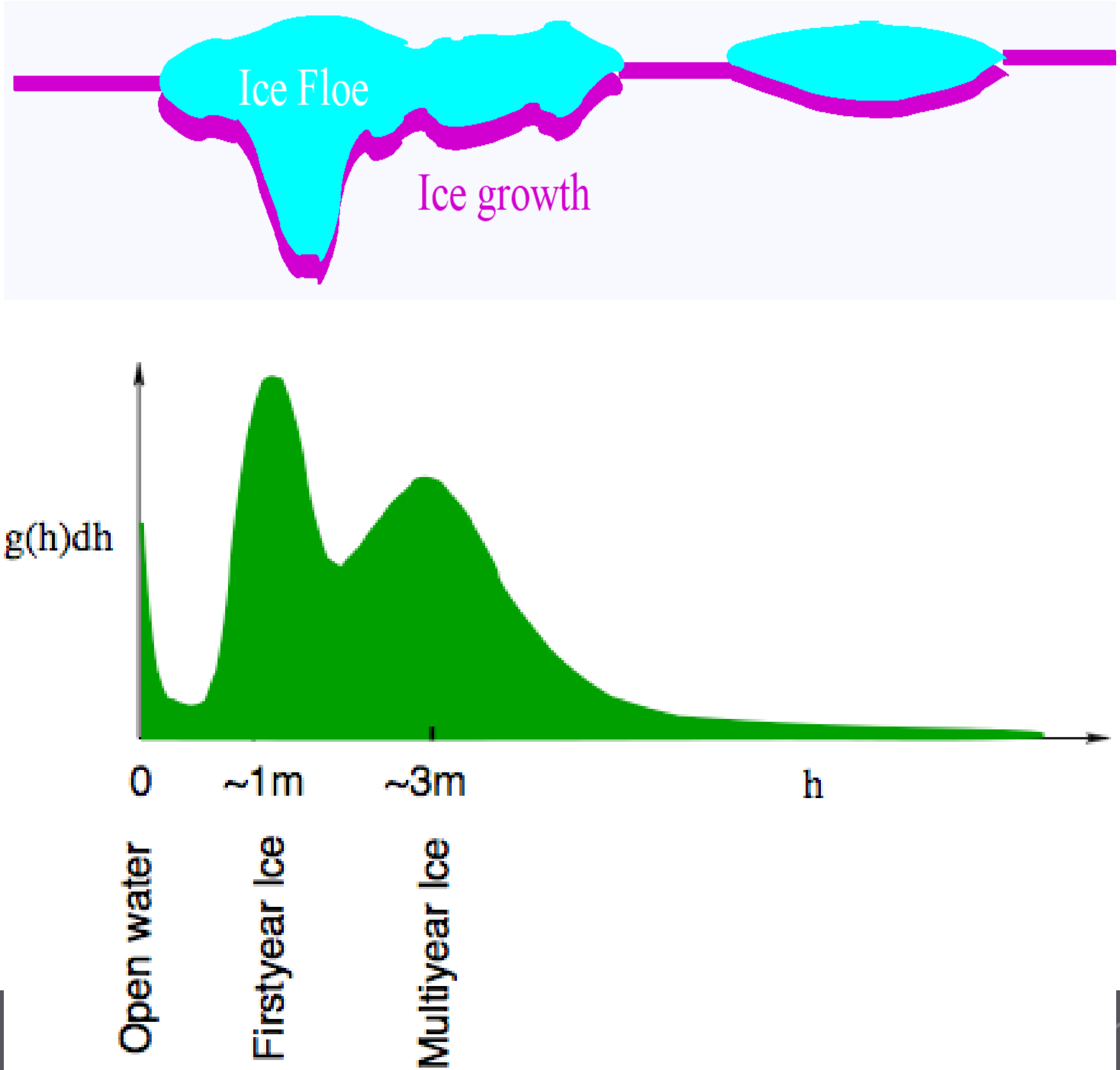


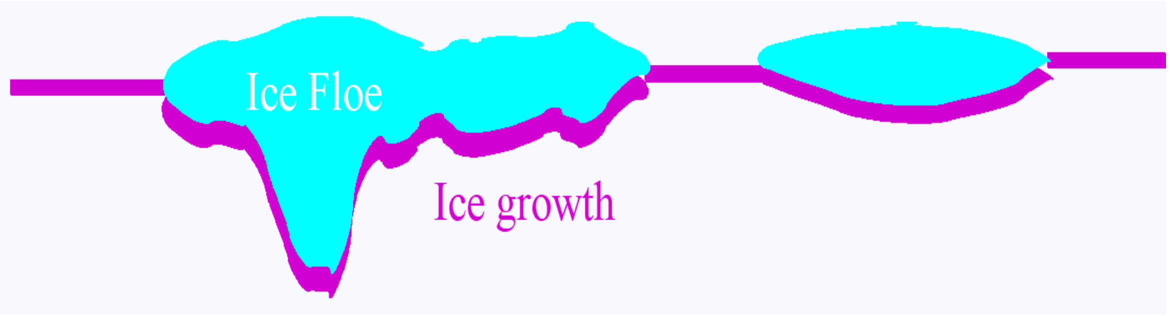
Converging hypothetical floes

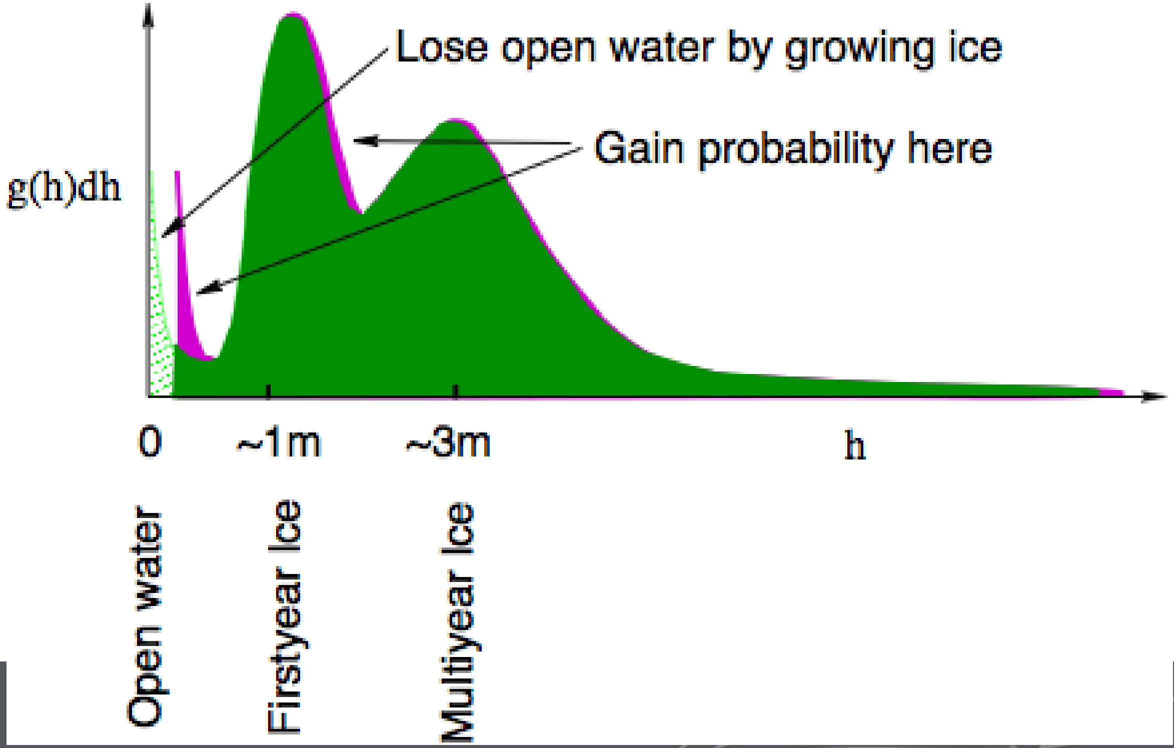


Ice growth:


Ice Floe
Ice growth
g(h)dh
0
~1m
~3m
h
Open water
Firstyear Ice
Multiyear Ice

Ice growth:


Ice Floe
Ice growth


Lose open water by growing ice
Gain probability here
g(h)dh
0
~1m
~3m
h
Open water
Firstyear Ice
Multiyear Ice