# Disentangling warming and circulation influences on precipitation

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# Disentangling warming and circulation influences on precipitation using nudging experiments

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### Challenges of quantifying drivers of precipitation change

- While there is **high confidence** that anthropogenic forcing has contributed to mean multi-decadal precipitation changes across several regions (i.e., western Africa and southeastern South America)...
- ... it is challenging to robustly assess the magnitude of relative contributions of greenhouse gas forcing (including stratospheric ozone depletion) and different species of aerosols because of:
  - 1. The large role of internal variability
  - 2. Observational uncertainty
  - 3. Model uncertainty
  - 4. Forcing uncertainty

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## Two mechanisms contribute to precipitation change under global warming

#### Thermodynamic

- These changes follow Clausius-Clapeyron suggesting that a warmer atmosphere can hold more water vapor
  - "wet get wetter and dry get drier"

#### Dynamic

- Result from shifts in atmospheric circulation which affect the horizontal and vertical transports of water vapor
- Modeling studies indicate that increasing greenhouse gas concentrations leads to:
  - An expansion of the tropical Hadley cell and subtropical dry zones
  - A poleward shift in storm tracks

### **Driving research questions**

- 1. Can we better understand the role of internal variability on precipitation change by quantifying the relative contributions of **dynamic** vs. **thermodynamic** changes in the CESM2 Large Ensemble?
- 2. Using idealized experiments, what is the influence of warming alone on precipitation change?

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## Quantifying dynamic vs. thermodynamic contributions to precipitation change

- Using CESM2 Large Ensemble:
  - Daily 500 hPa vertical velocity ( $\omega$ )
  - Daily precipitation
  - Consider the mid-latitudes (broadly; 20°-70°)
  - Compare two periods:
    - Historical  $\rightarrow$  1981-2000
    - Future → 2081-2100
  - NH winter (DJF)
  - Fully coupled ocean-atmosphere model
  - 50 members with CMIP6 forcing
  - SSP3-7.0



Danabasoglu et al. 2019 JAMES; Edwards 2010 WIREs Climate Change

## Quantifying dynamic vs. thermodynamic contributions to precipitation change

- **1**. Obtain the PDF of  $\omega$  (PDF $_{\omega}$ ) for a historical and future period
- 2. Composite daily precipitation for each  $\omega$  bin for both periods (P<sub> $\omega$ </sub>)

Example for all NH mid-latitudes

## Quantifying dynamic vs. thermodynamic contributions to precipitation change

- **1**. Obtain the PDF of  $\omega$  (PDF $_{\omega}$ ) for a historical and future period
- 2. Composite daily precipitation for each  $\omega$  bin for both periods (P<sub> $\omega$ </sub>)
- 3. 'Dynamic change' is the change in mean precipitation due to the change in  $PDF_{\omega}$ :

 $\mathsf{P}_{\omega[\text{historical}]}(\mathsf{PDF}_{\omega[\text{future}]} - \mathsf{PDF}_{\omega[\text{historical}]})$ 

4. 'Thermodynamic change' is the change in the expected precipitation for a given  $\omega$ :

 $PDF_{\omega[historical]} (P_{\omega[future]} - P_{\omega[historical]})$ 

### **Ensemble mean NH winter precipitation changes**

#### Key changes:

- Pî over land
- P↓ over Mediterranean regions
- P↓ equatorward of storm tracks
- P↑ poleward of storm tracks





What spatial patterns lead to the variability in projected precipitation change?





Total precipitation change

#### Total precipitation change



Thermodynamic change



#### Total precipitation change



Thermodynamic change



Total precipitation change



### Total precipitation change, ensemble mean removed

Thermodynamic change



## What are the large-scale atmospheric circulation changes across these SOMs?



SOM1 has a strong negative trend that allows the eastward extension and strengthening of the subtropical jet

## Question 1: Can we better understand the role of internal variability on precipitation change by quantifying the relative contributions of dynamic vs. thermodynamic changes in the CESM2 Large Ensemble?

- Why do some ensemble members see an expansion and strengthening of the jet while others do not?
- Can this help us rule out some of the scenarios in CESM2?
  - Similar to questions posed by Grise (2022; *GRL*)

500 hPa geopotential heights (zonal mean removed) change, ensemble mean removed SOM<sub>3</sub> SOM1 20 55°N 15 40°N 10 25°N 5 meters 0 SOM7 SOM9 -5 55°N -1040°N -1525°N -20 120°W 150°W 120°W 150°W 90°W 90°W

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### CAM6 experimental design

#### AMIPnudged\_hist

- Winds nudged to ERA5 reanalysis every 6 hours from 850 hPa to the top of the atmosphere
- 1979-2014
- Prescribed historical SSTs and sea ice extent

#### AMIPnudged\_fut

- Winds nudged to ERA5 reanalysis every 6 hours from 850 hPa to the top of the atmosphere
- 1979-1998
- Prescribed SSTs and sea ice extent taken from the CESM2-LE mean from 2081-2100



### Next steps

- What are the precipitation trends in each nudged run?
- What does the tropical Pacific look like for each nudged run?
- Can we use idealized simulations where we nudge horizontal winds to better understand hydroclimate trends?



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