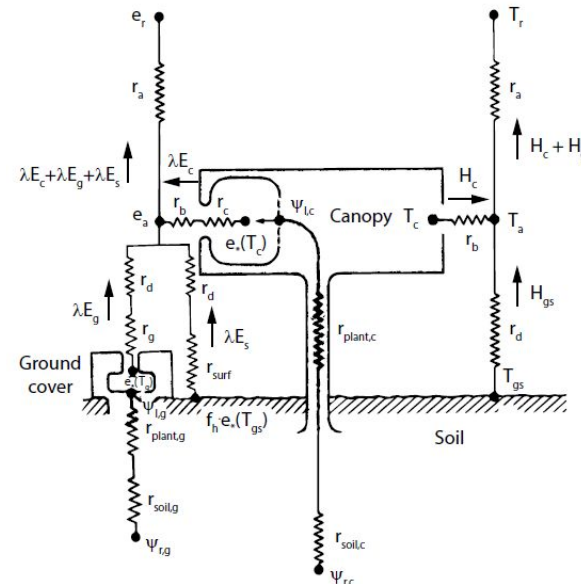




Thoughts on modeling plant canopies: Insights from the Canopy Horizontal Array Turbulence Study (CHATs)

Gordon Bonan, Sean Burns & Edward Patton
NSF National Center for Atmospheric Research
Boulder, Colorado, USA

CESM Land Model Working Group
24 February 2025



Multilayer canopy

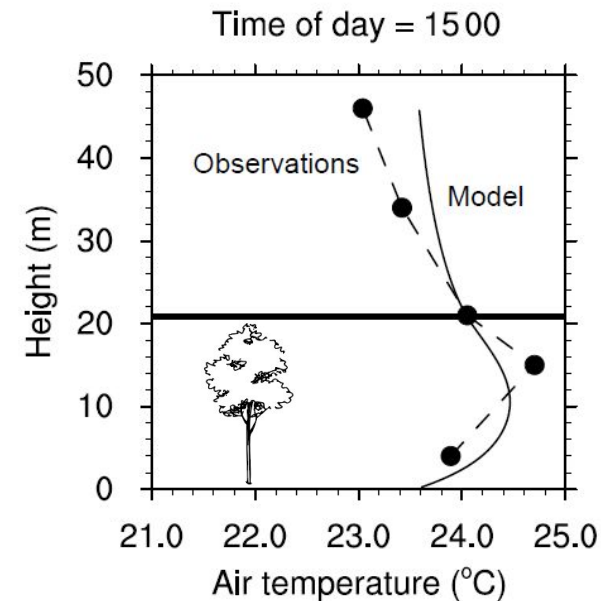
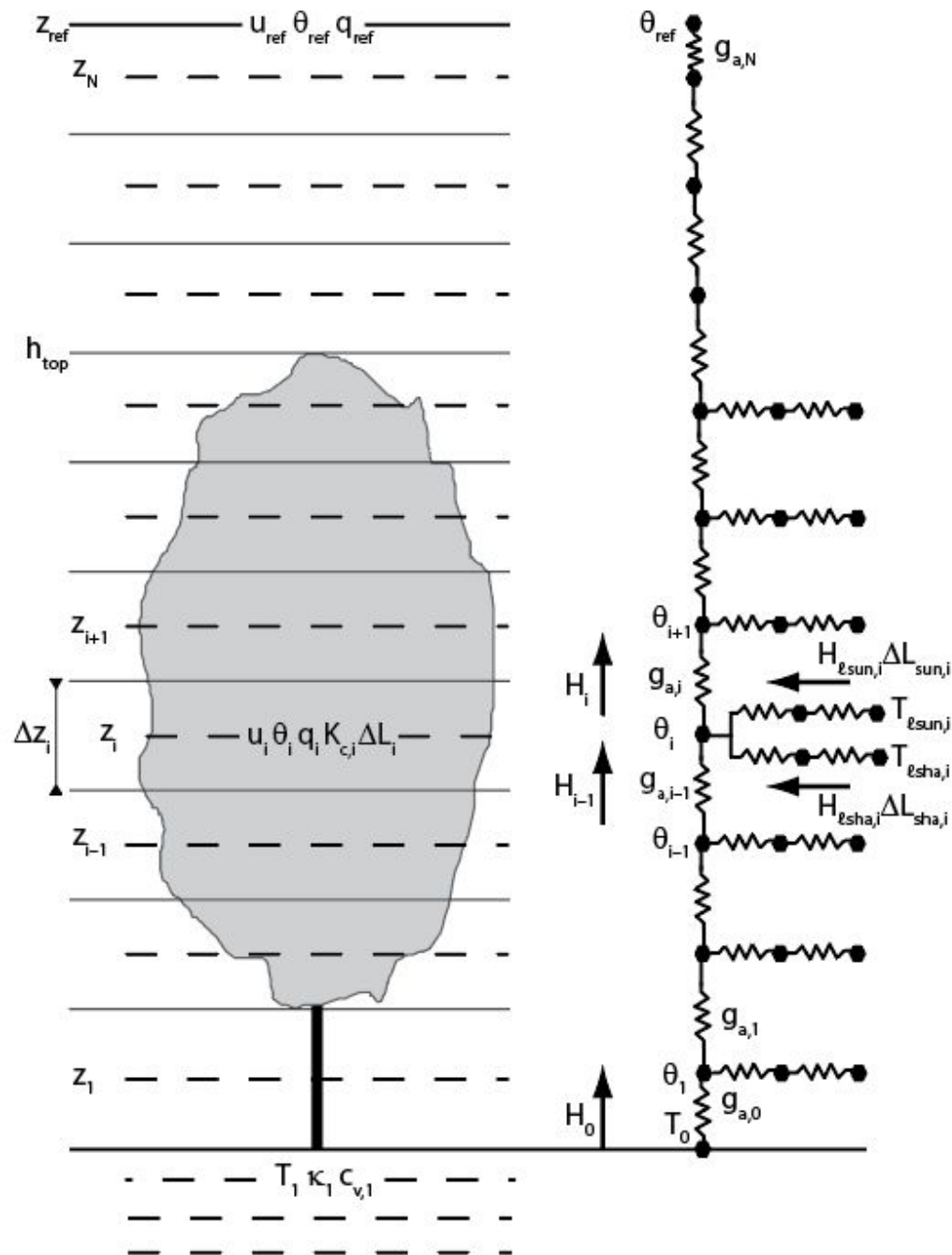
The physics and physiology of the multilayer canopy are simpler and more consistent with theory (and directly observable) than is the CLM5/6 big-leaf canopy (with many ad-hoc parameterizations and much technical debt), and the model enables new science

Bonan, Williams et al. (2014) *Geosci. Model Dev.*, 7, 2193-2222

Bonan, Patton, et al. (2018) *Geosci. Model Dev.*, 11, 1467-1496

Bonan, Patton, Finnigan, et al. (2021) *Agric. For. Meteorol.*, 306, 108435

Bonan, Burns & Patton (2025), in prep.



University of
Michigan
Biological Station

Observational dataset

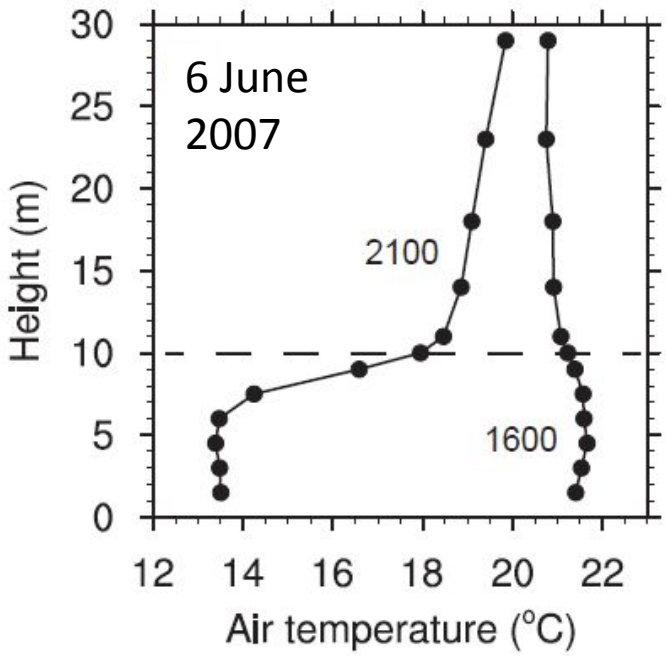
THE CANOPY HORIZONTAL ARRAY TURBULENCE STUDY

BY EDWARD G. PATTON, THOMAS W. HORST, PETER P. SULLIVAN, DONALD H. LENSCHOW, STEVEN P. ONCLEY, WILLIAM O. J. BROWN, SEAN P. BURNS, ALEX B. GUENTHER, ANDREAS HELD, THOMAS KARL, SHANE D. MAYOR, LUCIANA V. RIZZO, SCOTT M. SPULER, JIELUN SUN, ANDREW A. TURNIPSEED, EUGENE J. ALLWINE, STEVEN L. EDBURG, BRIAN K. LAMB, RONI AVISSAR, RONALD J. CALHOUN, JAN KLESSL, WILLIAM J. MASSMAN, KYAW THA PAW U, AND JEFFREY C. WEIL



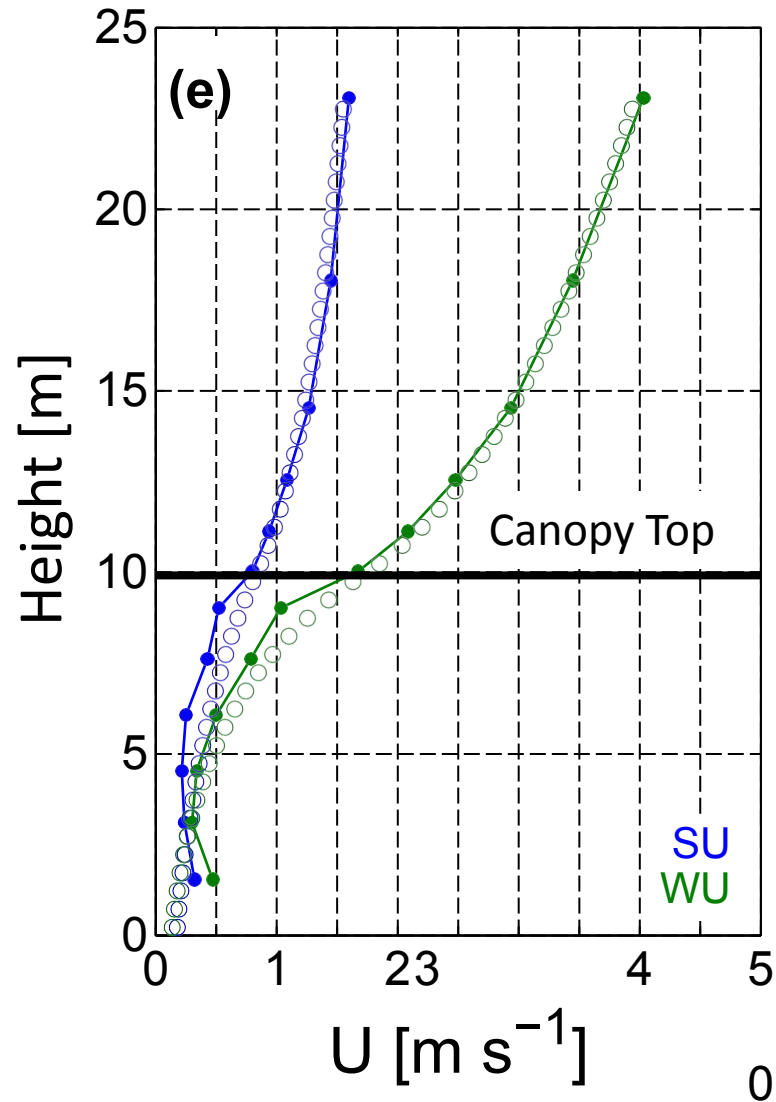
Patton et al. (2011) *BAMS*, 92, 593-611

mid-March – mid-June 2007

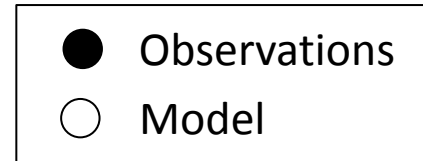
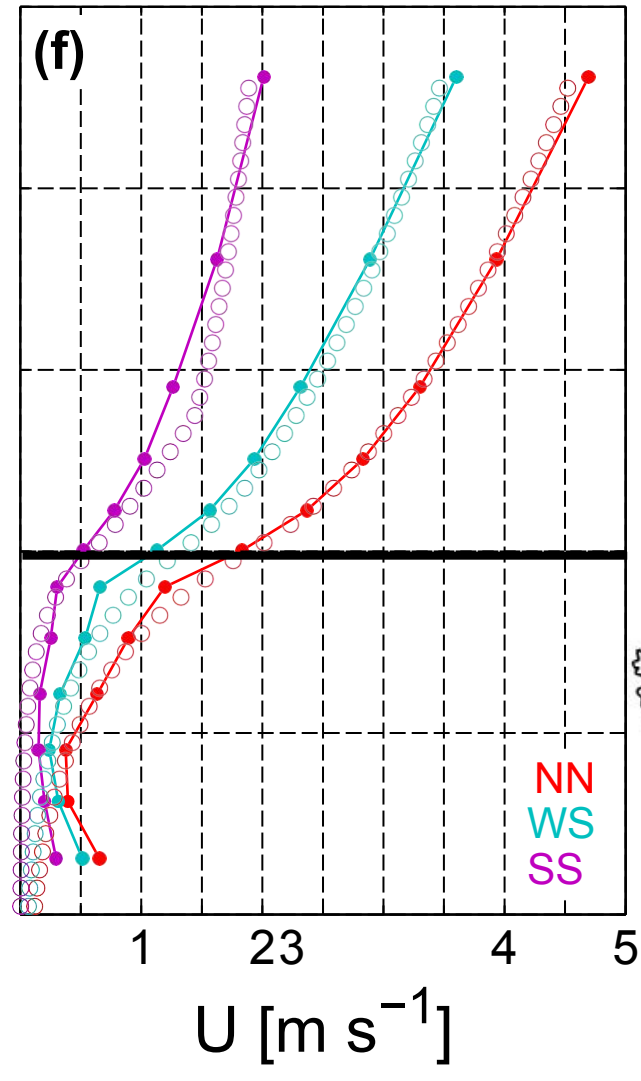


Wind speed (May 2007)

Unstable conditions



Near-neutral / stable



Stability classes

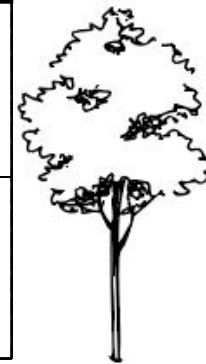
SU: strongly unstable

WU: weakly unstable

NN: near-neutral

WS: weakly stable

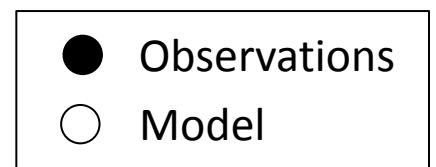
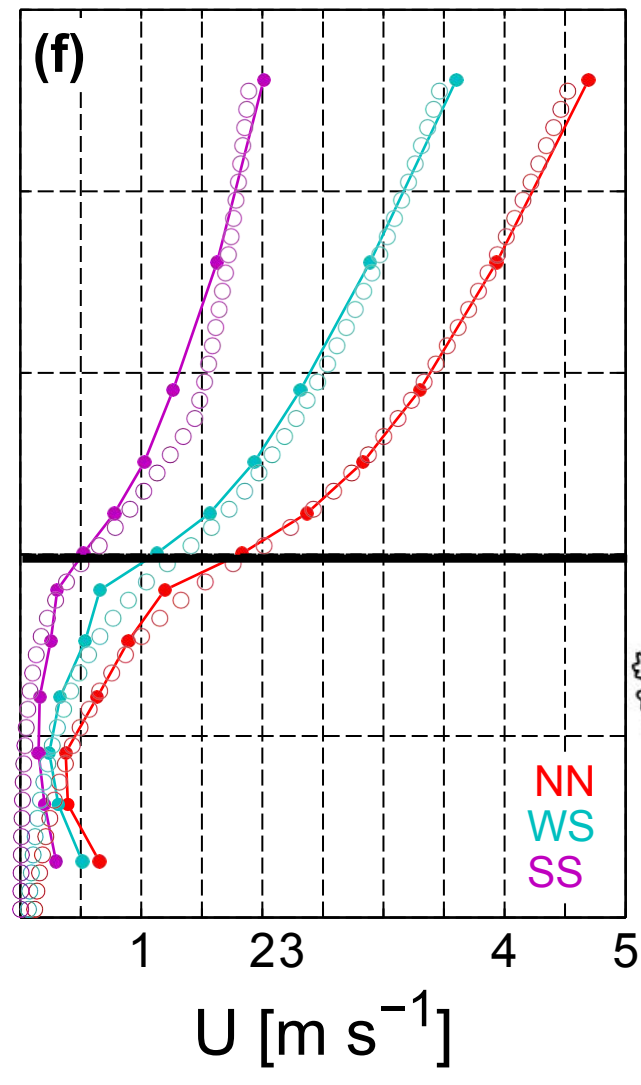
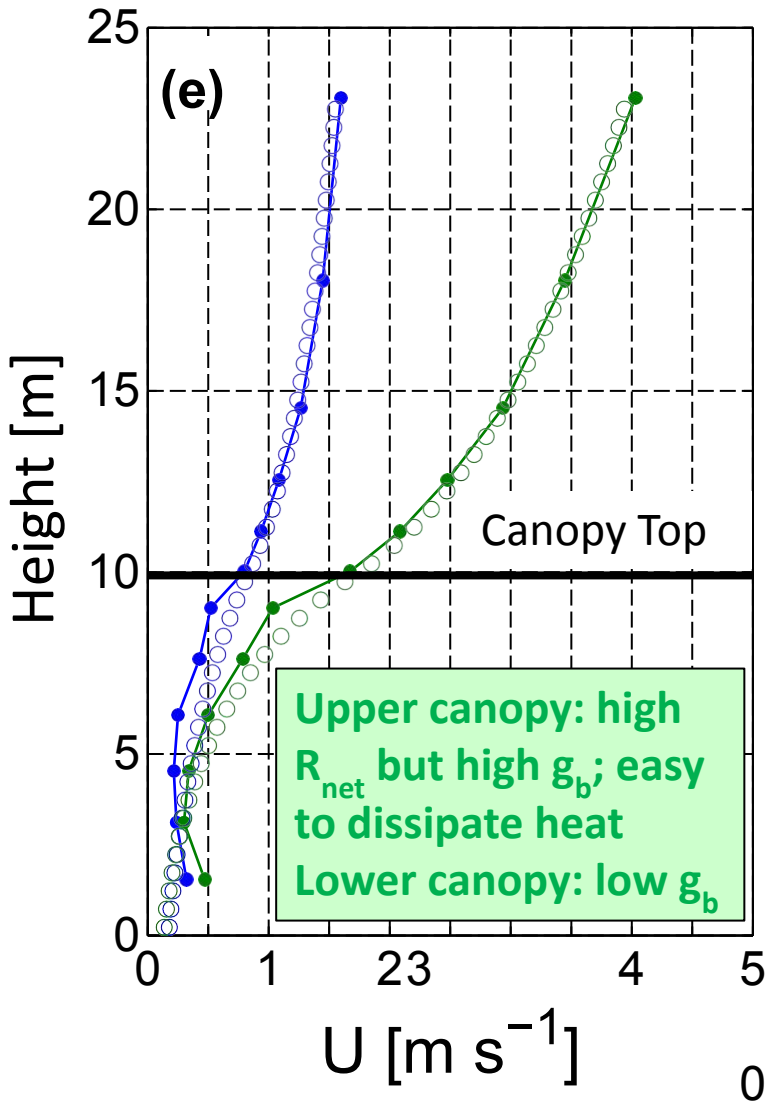
SS: strongly stable



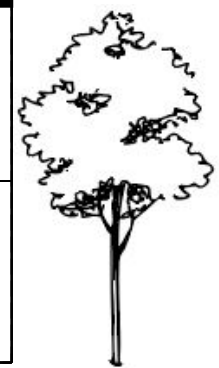
Wind speed (May 2007)

Unstable conditions

Near-neutral / stable



- Stability classes**
- SU: strongly unstable
 - WU: weakly unstable
 - NN: near-neutral
 - WS: weakly stable
 - SS: strongly stable



Analytical solution to the one-dimensional (vertical) momentum conservation equation

CLM5/6 canopy physics: wind speed in canopy

1984

$$U_{af} = u_*$$

“The assumption that $U_{af} = u_*$ in general **is of adequate accuracy for the present model**, considering the uncertainty in the other aspects of the foliage conductance.”

Dickinson (1984) in *Climate Processes and Climate Sensitivity* (Hansen & Takahashi, eds.; AGU)

2025: CLM6

CLM6 has a bulk wind speed in the canopy (u_{af}) that is equal to u_* . u_{af} is used to calculate leaf boundary layer conductance.

$$u_{af}(p) = \underbrace{um(p) * \sqrt{1._r8 / (ram1(p) * um(p))}}_{\text{Simplifies to } u_*}$$

Introduces an under-canopy wind speed (u_{uc})

```
! empirical undercanopy wind speed
uuc(p) = min(0.4_r8, (0.03_r8 * um(p) / ustar(p)))
```

CLM5/6 canopy physics: wind speed in canopy

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CLM6: multiple wind speeds

uaf: u_* is a measure of turbulence not wind
uuc: introduced with biomass heat storage
u10 (10-m wind) and fire spread (?)

2025: CLM6

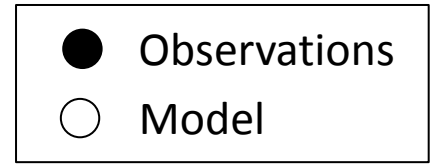
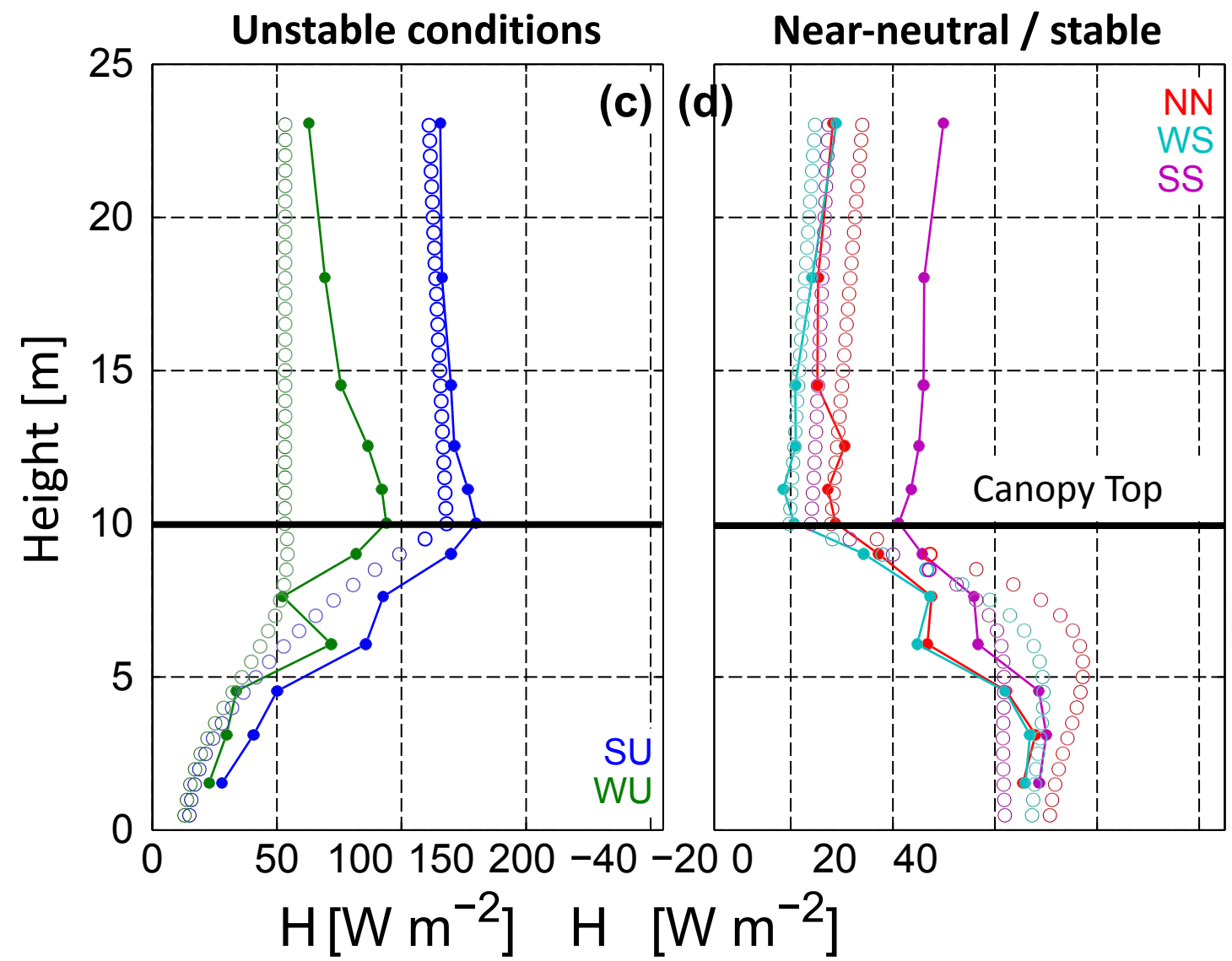
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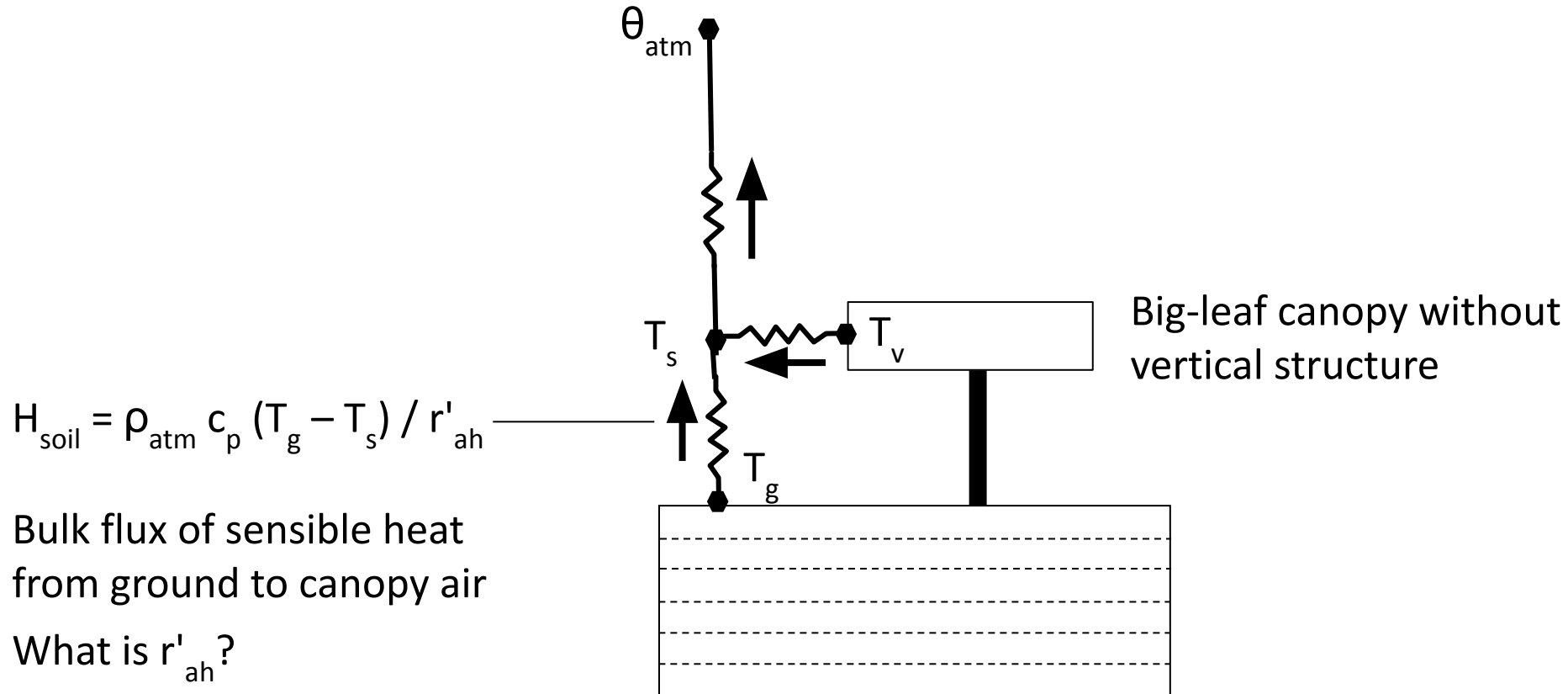
Sensible heat flux



- Stability classes**
- SU: strongly unstable
 - WU: weakly unstable
 - NN: near-neutral
 - WS: weakly stable
 - SS: strongly stable

Uses first-order turbulence closure

CLM5/6 perspective of the land surface



CLM5/6 canopy physics: under-canopy turbulence

1984

Under-canopy H

$$H_g = \rho_a c_p C_D u_{af} (T_g - T_{af})$$

Dickinson (1984) in *Climate Processes and Climate Sensitivity* (Hansen & Takahashi, eds.; AGU)

2025: CLM5/6

$$H_{soil} = -\rho_{atm} C_p \frac{(T_s - T_1)}{r'_{ah}} \quad (2.5.90)$$

$$r'_{ah} = r'_{aw} = \frac{1}{C_s U_{av}} \quad (2.5.116)$$

https://escomp.github.io/ctsm-docs/versions/release-clm5.0/html/tech_note/index.html

CLM5/6 canopy physics: under-canopy turbulence

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https://escomp.github.io/ctsm-docs/versions/release-clm5.0/html/tech_note/index.html

<u>Year</u>	<u>Model</u>	<u>Implementation</u>
1993	BATS1e	$C_s = 0.004$
2004	CLM3	$C_s = 0.004$ (dense canopy) → bare soil as PAI = 0
2010	CLM4	C_s modified for under-canopy stability
2018	CLM5	C_s under-canopy stability removed
2025	CLM6	Replace uaf with uuc (biomass heat storage)

CLM5/6 canopy physics: under-canopy turbulence

1984

Under-canopy H

$$H_g = \rho_a c_p C_D u_{af} (T_g - T_{af})$$

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2025: CLM6

```
uaf(p) = um(p)*sqrt( 1._r8/(ram1(p)*um(p)) )
```

```
! empirical undercanopy wind speed
```

```
uuc(p) = min(0.4_r8,(0.03_r8*um(p)/ustar(p)))
```

```
if (use_biomass_heat_storage) then
```

```
! use uuc for ground fluxes (keep uaf for canopy terms)
```

```
rah(p,below_canopy) = 1._r8/(csoilcn*uuc(p))
```

```
else
```

```
rah(p,below_canopy) = 1._r8/(csoilcn*uaf(p))
```

```
endif
```

CLM5/6 canopy physics: under-canopy turbulence

- Conceptual framework is unchanged over 40+ years
- C_s and u change as needed for a particular model configuration
- uuc changes fluxes in a way unrelated to the storage of heat in biomass. It changes the under-canopy turbulence.

2025: CLM6

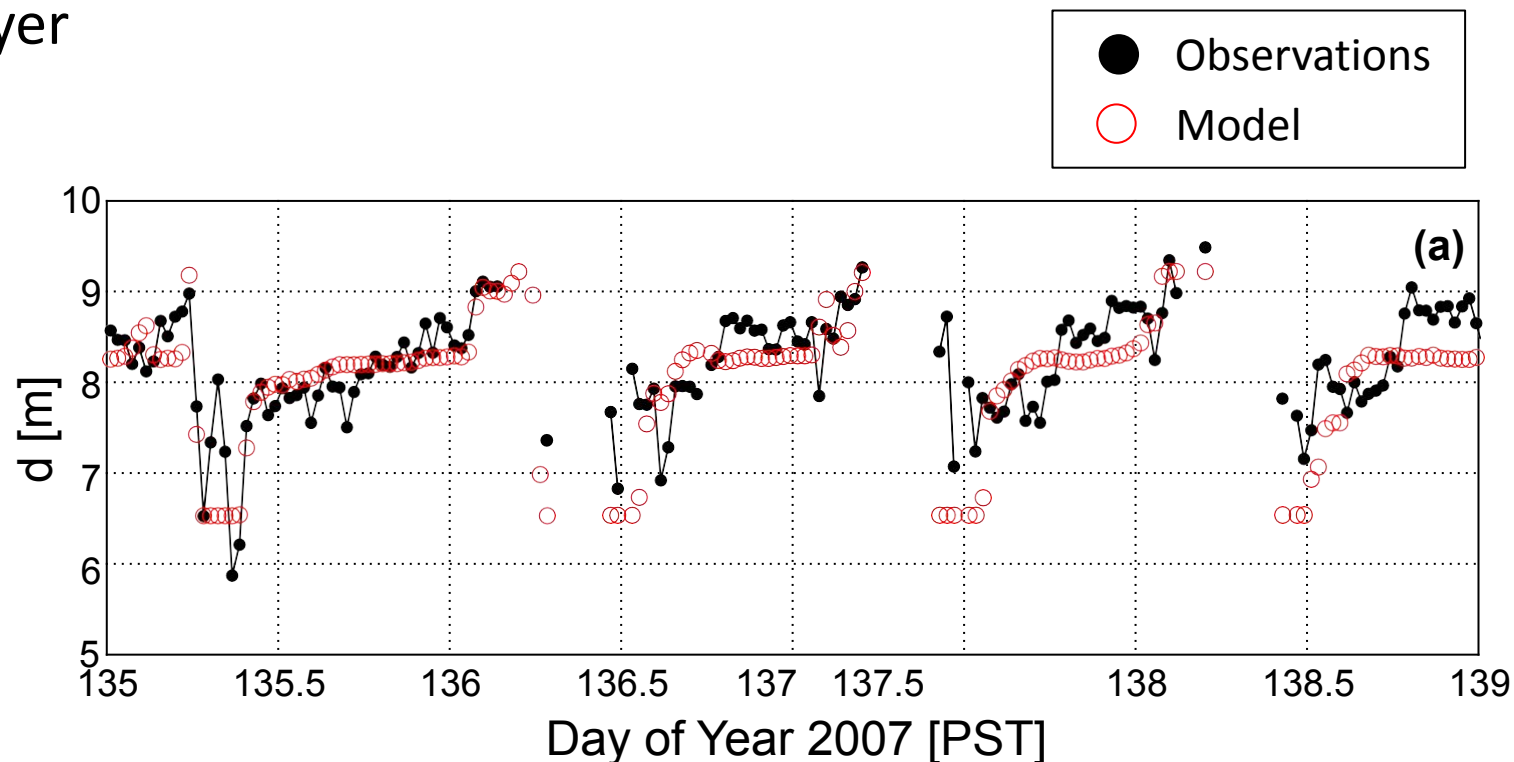
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uaf(p) = um(p)*sqrt( 1._r8/(ram1(p)*um(p)) )

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if (use_biomass_heat_storage) then
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else
  rah(p,below_canopy) = 1._r8/(csoilcn*uaf(p))
endif
```

Canopy turbulence: z_0 & d

- z_0 is an empirical parameter used in MOST
- Multilayer canopy replaces z_0 with observable canopy-imposed parameters (e.g., leaf drag coefficient)
- d emerges from roughness sublayer theory and depends on flow
- Varies by 3 m in a 10-m canopy!



CLM5/6 canopy physics

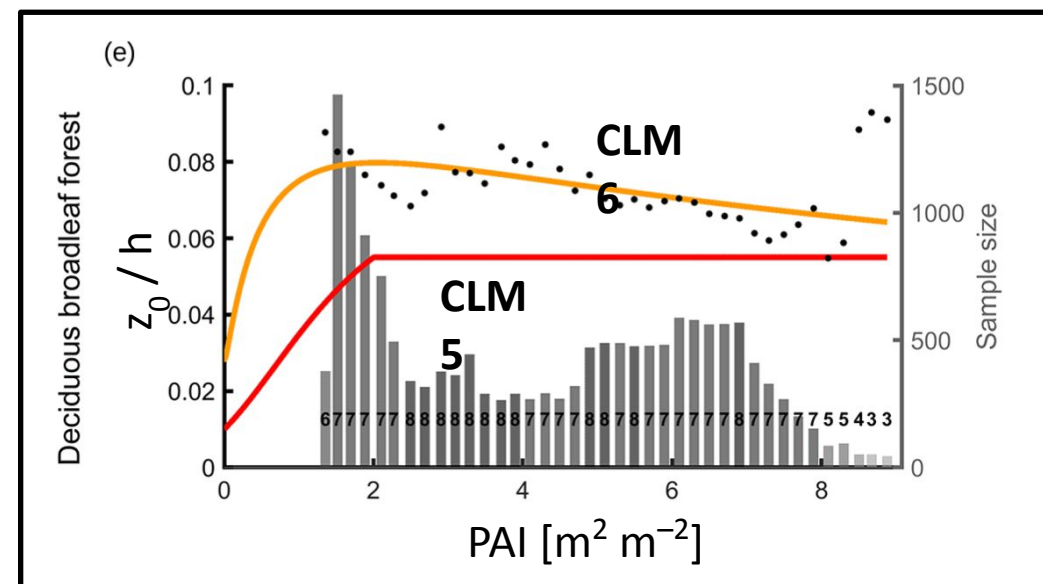
Roughness length (z_0) and displacement height (d)

Treated as canopy parameters: vary with h and PAI

CLM5: Zeng & Wang (2007)

CLM6: Meier et al. (2022) → Raupach (1994) fitted to FLUXNET data

FATES: passes z_0 and d to CLM: $f(h)$
& CLM6 adjusts for PAI using
Zeng & Wang (2007) as in CLM5



CLM5/6 canopy physics

Roughness length (z_0) and displacement height (d)

Treated as canopy parameters: vary with h and PAI

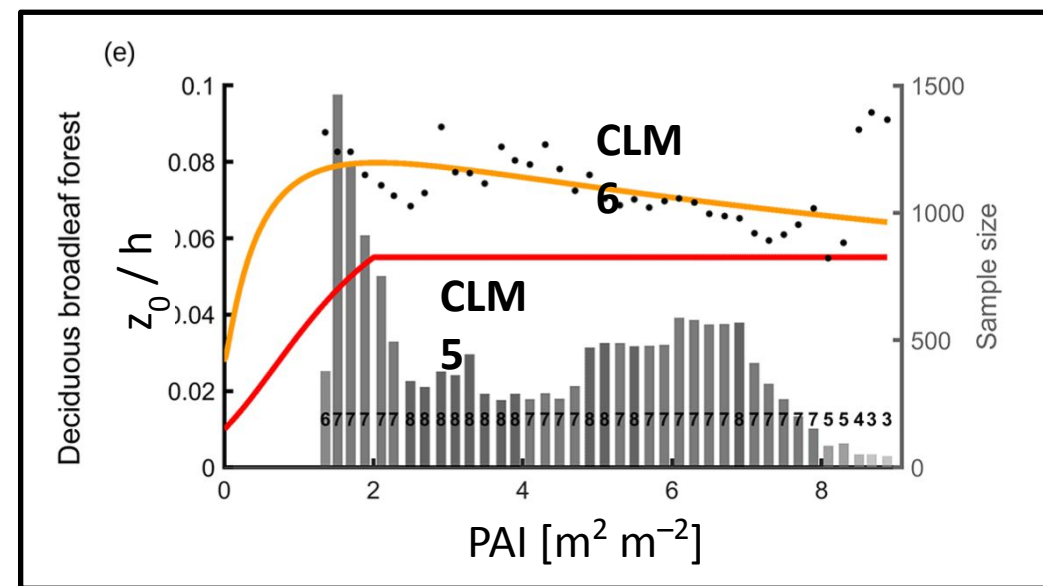
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Technical debt:

5 subroutines in 4 modules
6 variables in two data types



CLM5/6 canopy physics

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CLM5/6 and FATES treat z_0 and d as canopy structural parameters not as turbulent flow parameters

Leads to scientific debt in how z_0 and d are used throughout CLM, e.g.

$$\text{forc_hgt} = z_{\text{atm}} + z_0 + d$$

Canopy physics: a complex system of equations

Fluxes depend on canopy states but canopy states depend on fluxes:

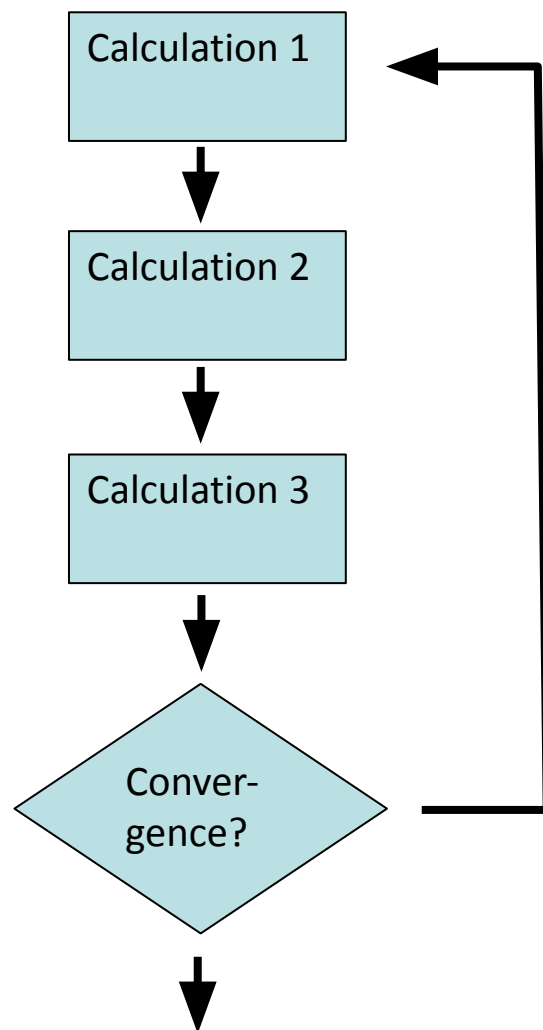
- $g_s = f(T_{\text{leaf}})$ and $T_{\text{leaf}} = f(g_s)$
- $L = f(H, \lambda E)$ and $H = f(L)$, $\lambda E = f(L)$

The physiological state of a plant community substantially influences the microclimate within it; in turn, the microclimate influences the physiological state, so that neither is independent of the other.

Finnigan & Raupach (1987) in *Stomatal Function*
(Zeiger, Farquhar & Cowan, eds.; Stanford Univ. Press)

CLM5/6 canopy physics: a deep dive

CLM5/6 use a multistep iterative solution for fluxes/temperature



Up to 40 iterations (standard)
and 3 iterations (fast)

- Initial values for canopy air temperature and specific humidity are obtained from

$$T_s = \frac{T_g + \theta_{atm}}{2} \quad (2.5.136)$$

$$q_s = \frac{q_g + q_{atm}}{2}. \quad (2.5.137)$$

- An initial guess for the wind speed V_a is obtained from (2.5.24) assuming an initial convective velocity $U_c = 0 \text{ m s}^{-1}$ for stable conditions ($\theta_{v,atm} - \theta_{v,s} \geq 0$ as evaluated from (2.5.50)) and $U_c = 0.5$ for unstable conditions ($\theta_{v,atm} - \theta_{v,s} < 0$).

- An initial guess for the Monin-Obukhov length L is obtained from the bulk Richardson number using equation and (2.5.46) and (2.5.48).

- Iteration proceeds on the following system of equations:

- Friction velocity u_* ((2.5.32), (2.5.33), (2.5.34), (2.5.35))

- Ratio $\frac{\theta_*}{\theta_{atm} - \theta_s}$ ((2.5.37), (2.5.38), (2.5.39), (2.5.40))

- Ratio $\frac{q_*}{q_{atm} - q_s}$ ((2.5.41), (2.5.42), (2.5.43), (2.5.44))

- Aerodynamic resistances r_{am} , r_{ah} , and r_{aw} ((2.5.55), (2.5.56), (2.5.57))

- Magnitude of the wind velocity incident on the leaves U_{av} ((2.5.117))

- Leaf boundary layer resistance r_b ((2.5.136))

- Aerodynamic resistances r'_{ah} and r'_{aw} ((2.5.116))

- Sunlit and shaded stomatal resistances r_s^{sun} and r_s^{sha} (Chapter 2.9)

- Sensible heat conductances c_a^h , c_g^h , and c_v^h ((2.5.94), (2.5.95), (2.5.96))

- Latent heat conductances c_a^w , c_v^w , and c_g^w ((2.5.108), (2.5.109), (2.5.110))

- Sensible heat flux from vegetation H_v ((2.5.97))

- Latent heat flux from vegetation λE_v ((2.5.101))

- If the latent heat flux has changed sign from the latent heat flux computed at the previous iteration ($\lambda E_v^{n+1} \times \lambda E_v^n < 0$), the latent heat flux is constrained to be 10% of the computed value. The difference between the constrained and computed value ($\Delta_1 = 0.1\lambda E_v^{n+1} - \lambda E_v^n$) is added to the sensible heat flux later.

- Change in vegetation temperature ΔT_v ((2.5.129)) and update the vegetation temperature as $T_v^{n+1} = T_v^n + \Delta T_v$. T_v is constrained to change by no more than 1°K in one iteration. If this limit is exceeded, the energy error is

$$(2.5.138)$$

$$\Delta_2 = \vec{S}_v - \vec{L}_v - \frac{\partial \vec{L}_v}{\partial T} \Delta T_v - H_v - \frac{\partial H_v}{\partial T} \Delta T_v - \lambda E_v - \frac{\partial \lambda E_v}{\partial T} \Delta T_v$$

- Water vapor flux E_v ((2.5.133))

- Transpiration E_v^t ((2.5.134) if $\beta_t > 0$, otherwise $E_v^t = 0$)

- The water vapor flux E_v is constrained to be less than or equal to the sum of transpiration E_v^t and the water available from wetted leaves and stems $W_{can}/\Delta t$. The energy error due to this constraint is

$$\Delta_3 = \max\left(0, E_v - E_v^t - \frac{W_{can}}{\Delta t}\right). \quad (2.5.139)$$

The error $\lambda\Delta_3$ is added to the sensible heat flux later.

- Sensible heat flux H_v ((2.5.135)). The three energy error terms, Δ_1 , Δ_2 , and $\lambda\Delta_3$ are also added to the sensible heat flux.
- The saturated vapor pressure e_i (Chapter 2.9), saturated specific humidity $q_{sat}^{T_v}$ and its derivative $\frac{dq_{sat}^{T_v}}{dT_v}$ at the leaf surface (section 2.5.5), are re-evaluated based on the new T_v .
- Canopy air temperature T_s ((2.5.93))

37 step solution

- Canopy air specific humidity q_s ((2.5.107))
- Temperature difference $\theta_{atm} - \theta_s$
- Specific humidity difference $q_{atm} - q_s$
- Potential temperature scale $\theta_* = \frac{\theta_*}{\theta_{atm} - \theta_s} (\theta_{atm} - \theta_s)$ where $\frac{\theta_*}{\theta_{atm} - \theta_s}$ was calculated earlier in the iteration
- Humidity scale $q_* = \frac{q_*}{q_{atm} - q_s} (q_{atm} - q_s)$ where $\frac{q_*}{q_{atm} - q_s}$ was calculated earlier in the iteration
- Virtual potential temperature scale θ_{v*} ((2.5.17))
- Wind speed including the convective velocity, V_a ((2.5.24))
- Monin-Obukhov length L ((2.5.49))
- The iteration is stopped after two or more steps if $\tilde{\Delta}T_v < 0.01$ and $|\lambda E_v^{n+1} - \lambda E_v^n| < 0.1$ where $\tilde{\Delta}T_v = \max(|T_v^{n+1} - T_v^n|, |T_v^n - T_v^{n-1}|)$, or after forty iterations have been carried out.
- Momentum fluxes τ_x , τ_y ((2.5.5), (2.5.6))
- Sensible heat flux from ground H_g ((2.5.89))
- Water vapor flux from ground E_g ((2.5.102))
- 2-m height air temperature T_{2m} , specific humidity q_{2m} , relative humidity RH_{2m} ((2.5.58), (2.5.59), (2.5.60))

The end!

https://escomp.github.io/ctsm-docs/versions/release-clm5.0/html/tech_note/index.html

CLM5/6 canopy physics: a deep dive

However, if ζ changes sign more than four times during the temperature iteration, $\zeta = -0.01$. This helps prevent “flip-flopping” between stable and unstable conditions.

If the latent heat flux has changed sign from the latent heat flux computed at the previous iteration ($\lambda E_v^{n+1} \times \lambda E_v^n < 0$), the latent heat flux is constrained to be 10% of the computed value. The difference is added to the sensible heat flux later.

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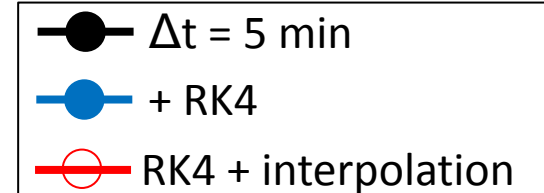
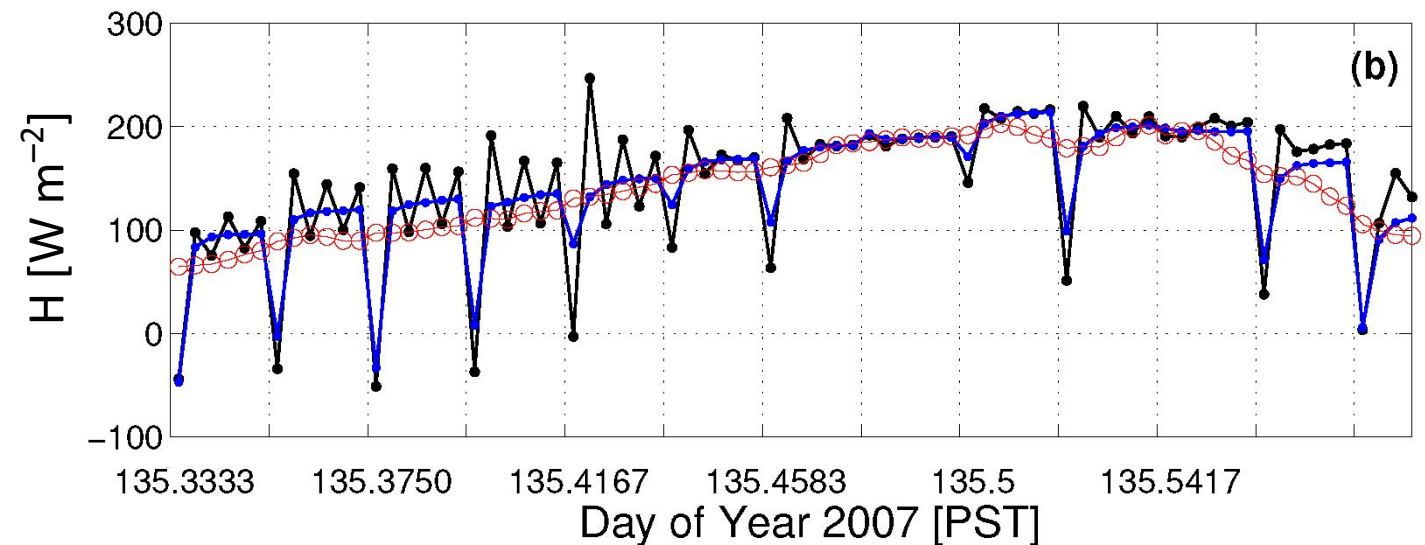


Numerical methods

CLM-ml v2:

- 5-min timestep instead of 30-min
- 4th-order Runge-Kutta (RK4)
- Interpolate 30-min forcing to 5-min

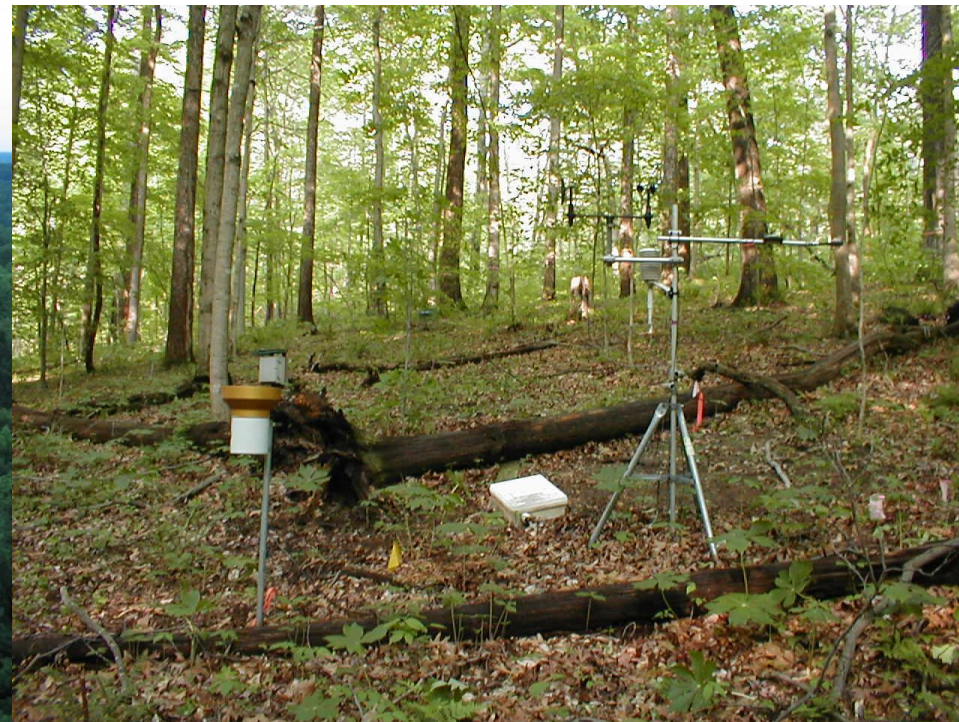
6-hour period from morning to mid-afternoon



Advanced numerical methods can eliminate need for ad-hoc “fixes”

Two ways to model plant canopies

Photographs of Morgan Monroe State Forest tower site illustrate two different representations of a plant canopy: as a “big leaf” (below) or with vertical structure (right)



Big-leaf canopy

“incorrect but useful”

Multilayer canopy

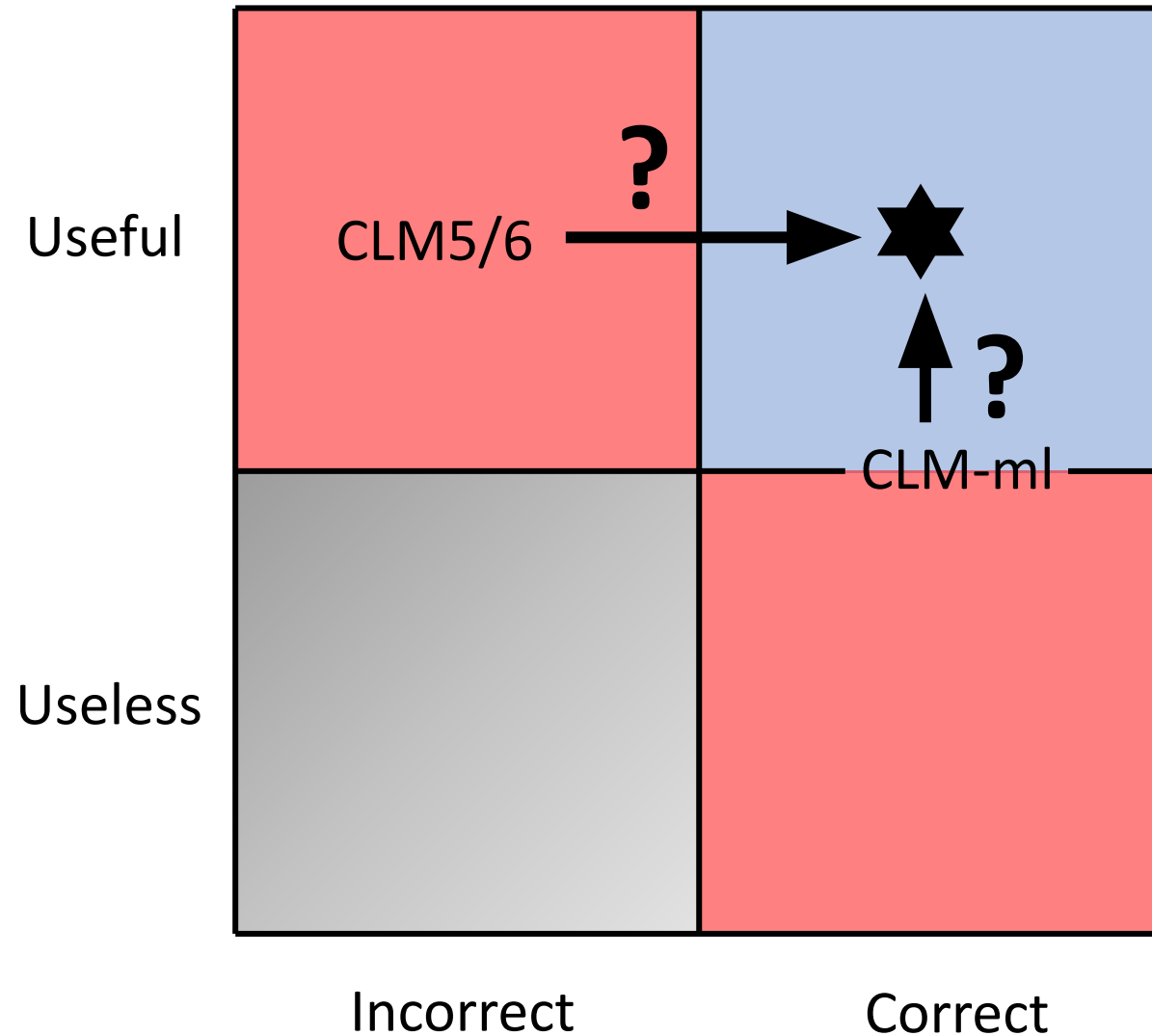
“correct but useless”



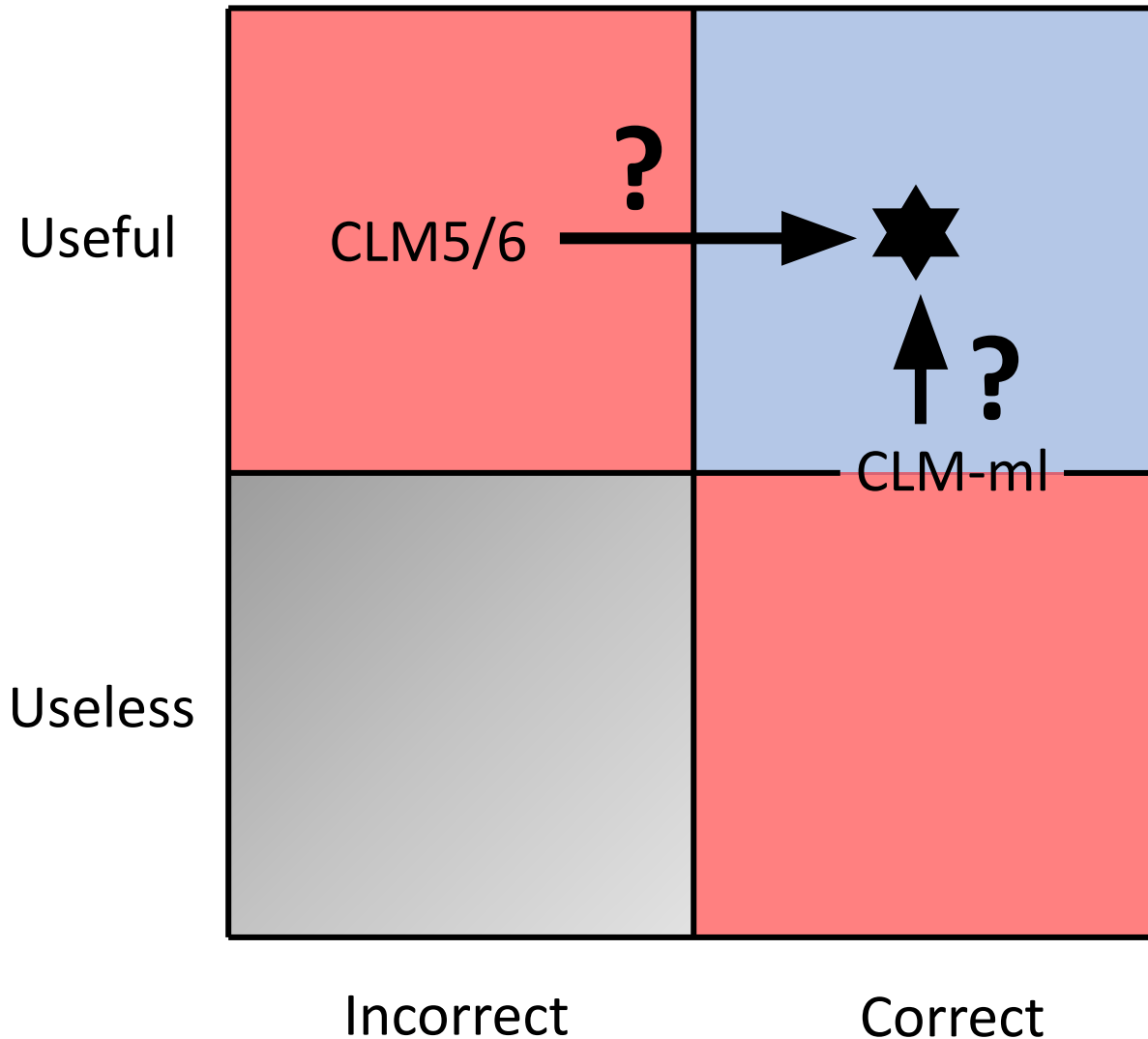
Raupach & Finnigan (1988) *Aust. J. Plant Physiol.*, 15, 705-716



Attaining the correct and useful model

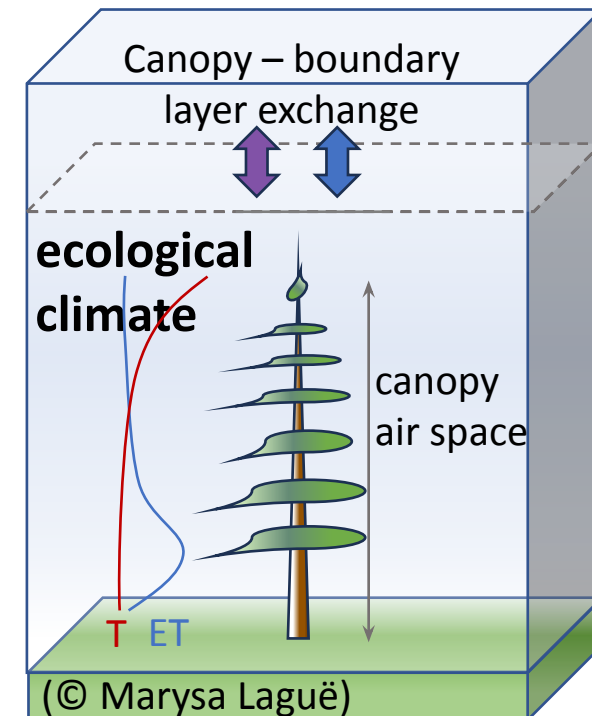


Attaining the correct and useful model

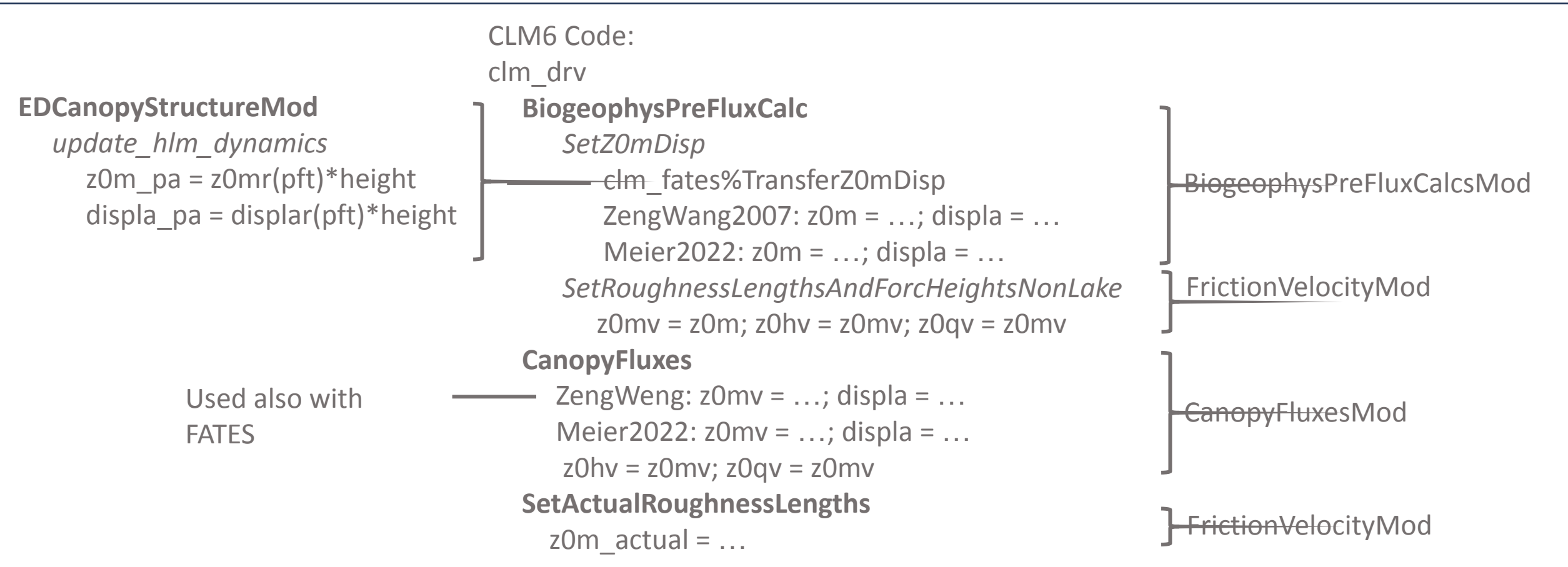


New science

- Directly observable (A_n , g_s , ψ_l , T_l , $\delta^{13}C$)
- Canopy chemistry (e.g., BVOCs, dry deposition)
- Forest microclimates (overstory, understory)
- Orphan air space



Technical debt



5 subroutines over 4 modules

6 variables over two data types

- canopystate_type%z0m_patch
- canopystate_type%displa_patch
- frictionvel_type%z0mv_patch
- frictionvel_type%z0hv_patch
- frictionvel_type%z0qv_patch
- frictionvel_type%z0m_actual_patch