

# Sources of the Arctic Atlantic Water Biases in CESM2

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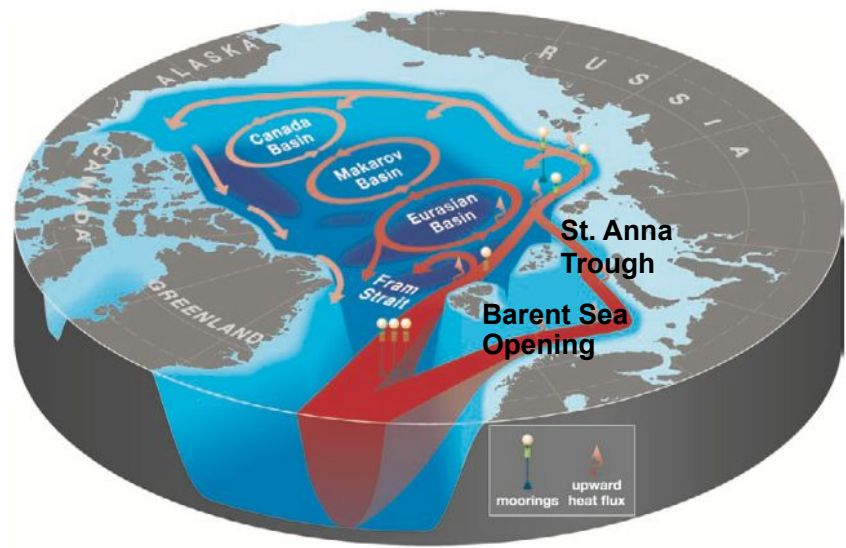
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**OMWG Winter Meeting 2025**

**February 27, 2025**

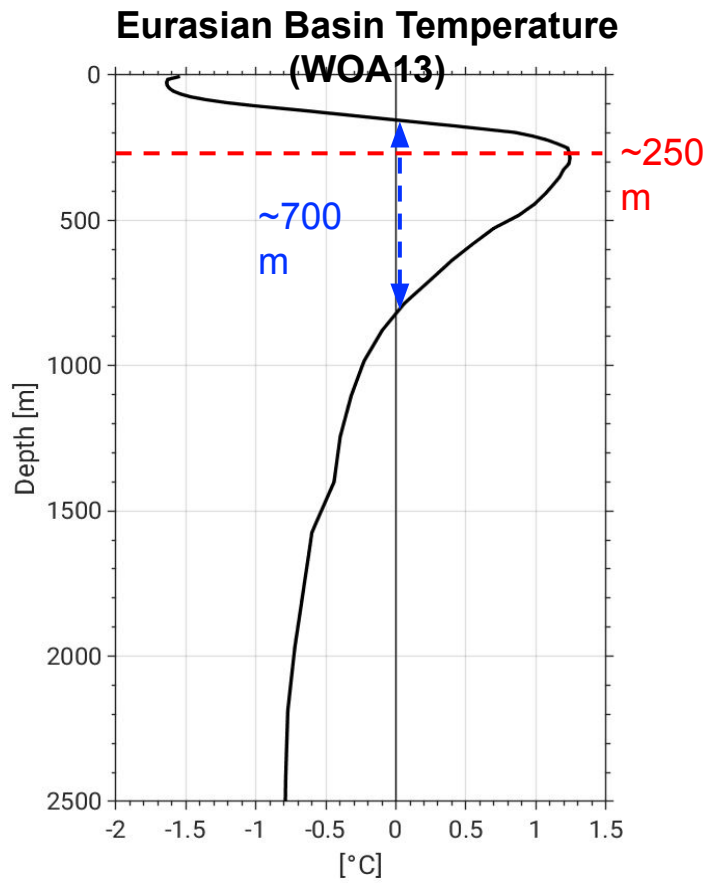


# Arctic Atlantic Water (AW)

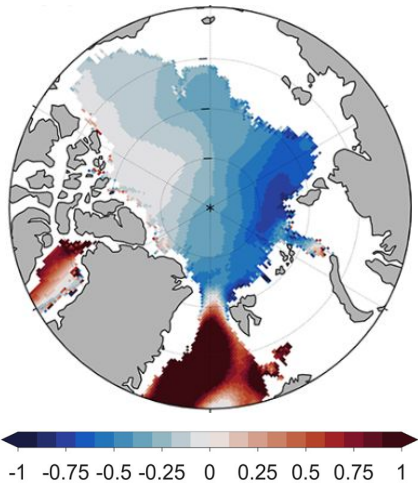
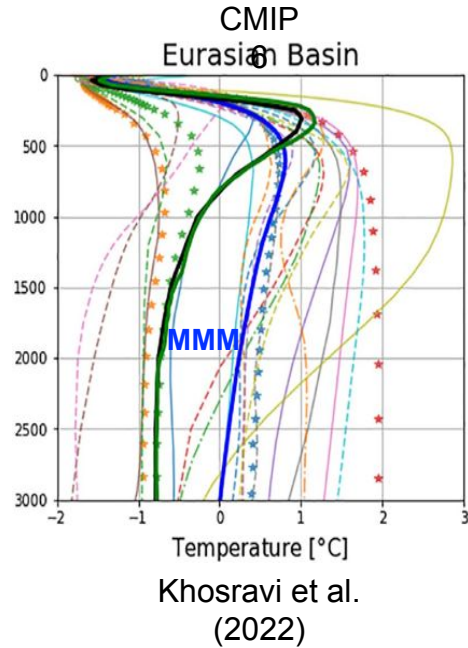
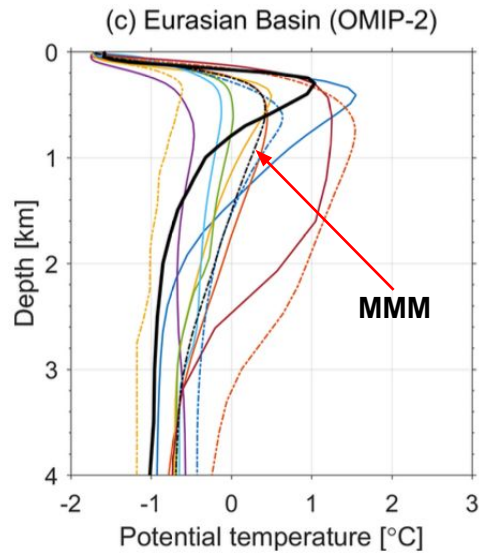


Polyakov et al. (2011)

- AW inflow to the Arctic Ocean**
- Fram Strait branch (~2 Sv): relatively warmer and saltier
  - Barents Sea branch (~2 Sv): relatively colder and fresher due to heat loss and mixing, entering the Arctic Ocean mainly through the St. Anna Trough
  - Flows cyclonically along the continental slope
  - AW in the interior Arctic Ocean:  $\theta > 0^\circ\text{C}$

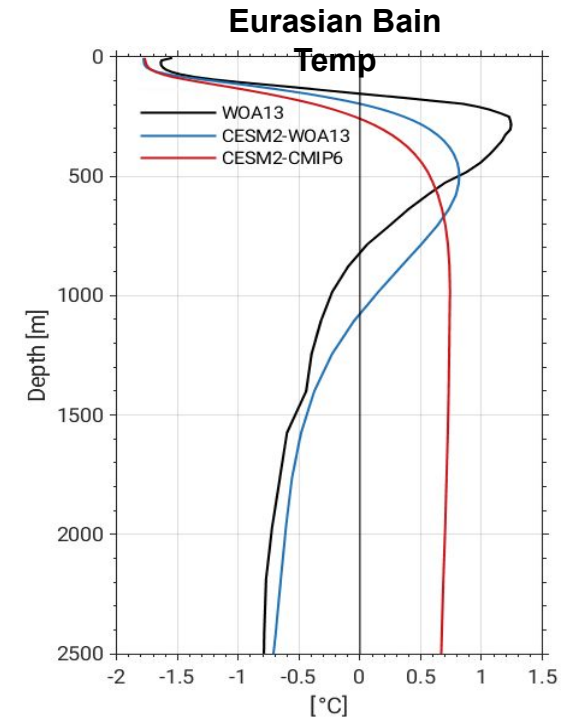


# Atlantic Water Biases

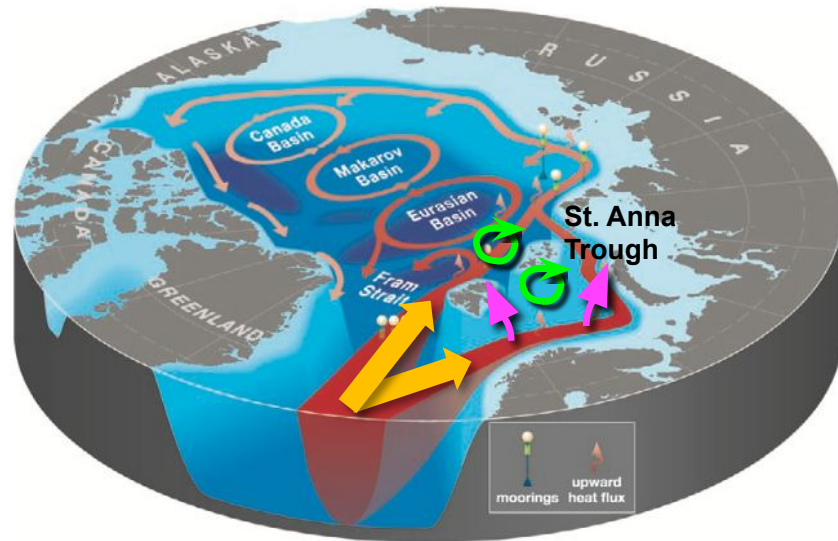


Shu et al. (2023)

- Too deep and thick
- Too cold around the observed AW core and too warm below
- Salinity: a fresh bias in the AW layer
- **Short-term** vs. long-term biases



# Potential Sources of the Biases



Polyakov et al.  
(2011)

- Biased AW inflow from the GIN Seas
- Incorrect water mass transformation, due to surface fluxes and/or mixing, over the shelf of Barents and Kara Seas and/or within the troughs and channels
- Inaccurate or misrepresented horizontal and vertical mixing processes along the slope and in the interior
- ▶ **No study systematically examining the bias sources using a single model**

# Model and Experiments

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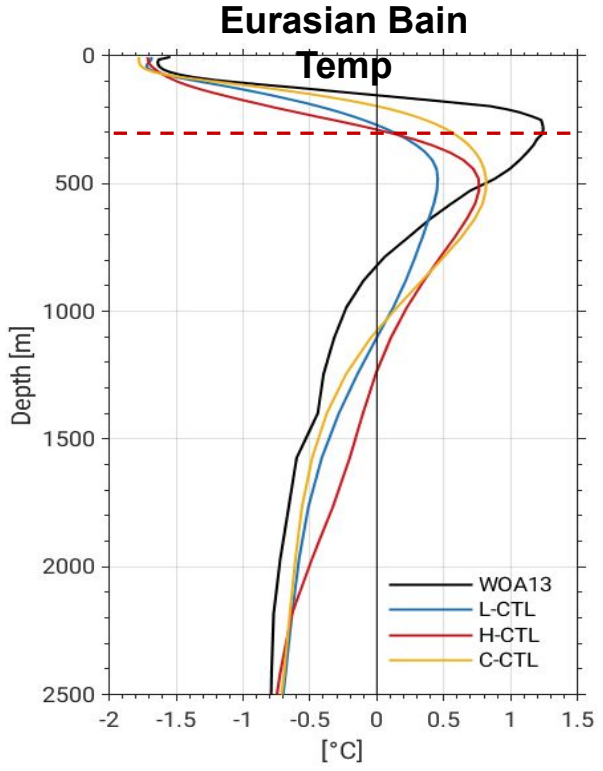
## Forced ocean–sea-ice simulations

- Both low- (LR) and high-resolution (HR) configurations of CESM2
- **Sensitivity experiments (L-Exp\_Name)**
  - 1) Interior T&S restoring experiments
  - 2) Mixing parameter experiments
  - 3) Combinations of (1) and (2)
- Selected experiments repeated using HR (H-Exp\_Name)
- 2003-2004 Repeat Year Forcing (RYF), starting from WOA13 ICs
  - *Similar biases appear whether IAF or different RYFs are used*

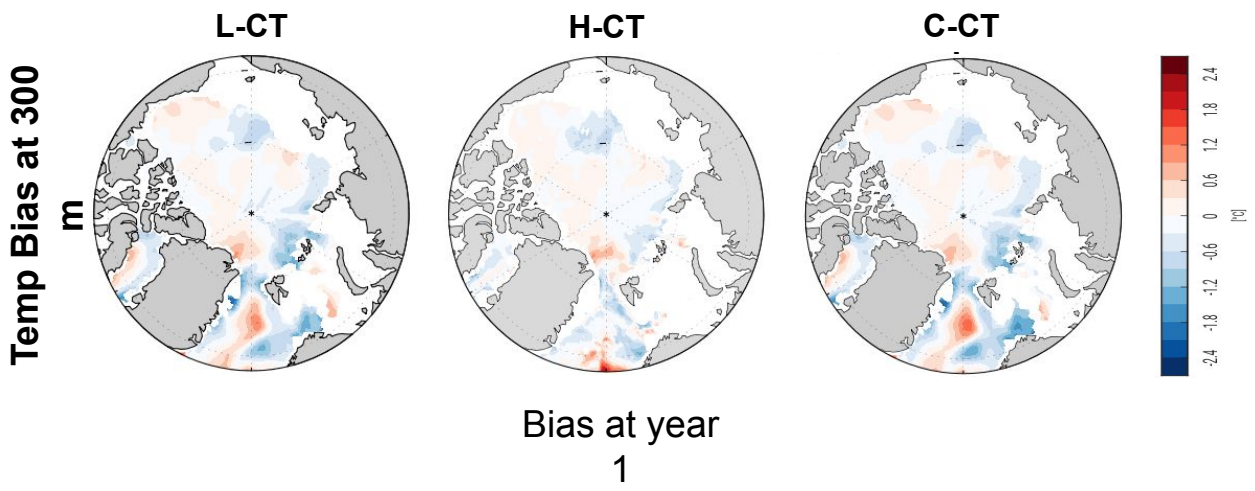
## Fully Coupled simulations

- Standard, LR configuration of CESM2
  - New pi-control, starting from WOA13 ICs
  - Selected experiments repeated (C-Exp\_Name)
- 
- All simulations run for 30 years
  - Last 5-year averages used to examine biases, focusing on the temperature and circulation in the Eurasian Basin (EB)

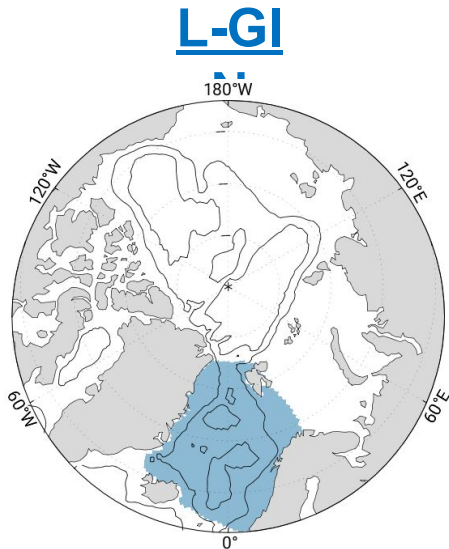
# AW Biases in the Control Simulations



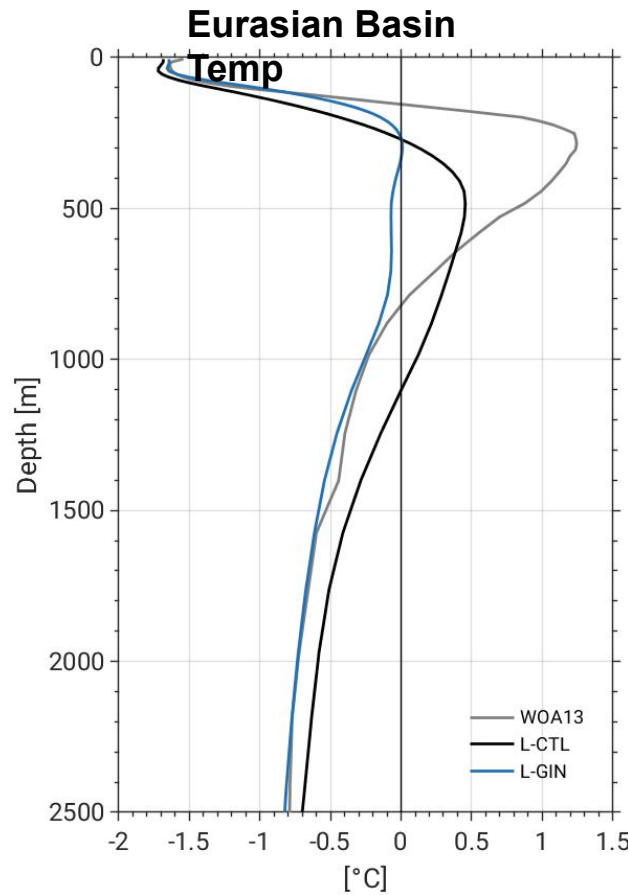
- AW core found at deeper depths
  - ~450-500 m vs. 250 m in WOA13
  - Deepest in H-CTL
- AW layer thicker
  - ~800-900 m vs. 700 m in WOA13
  - Thickest in H-CTL
- Colder at the observed core depth and warmer below
- Max temperature is coldest in L-CTL (<0.5°C vs. 1.2 °C in WOA13)
- Biases initially developing quickly along the slope



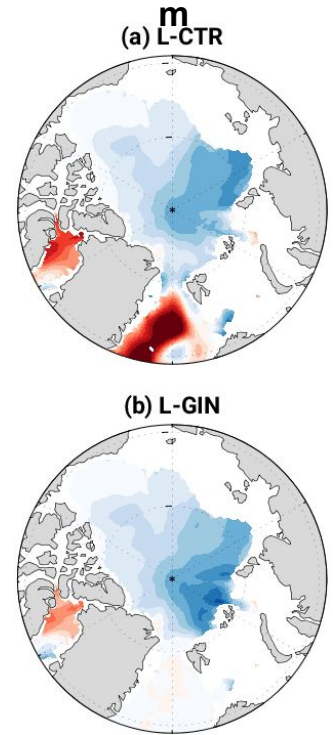
# Restoring Experiments (L-GIN)



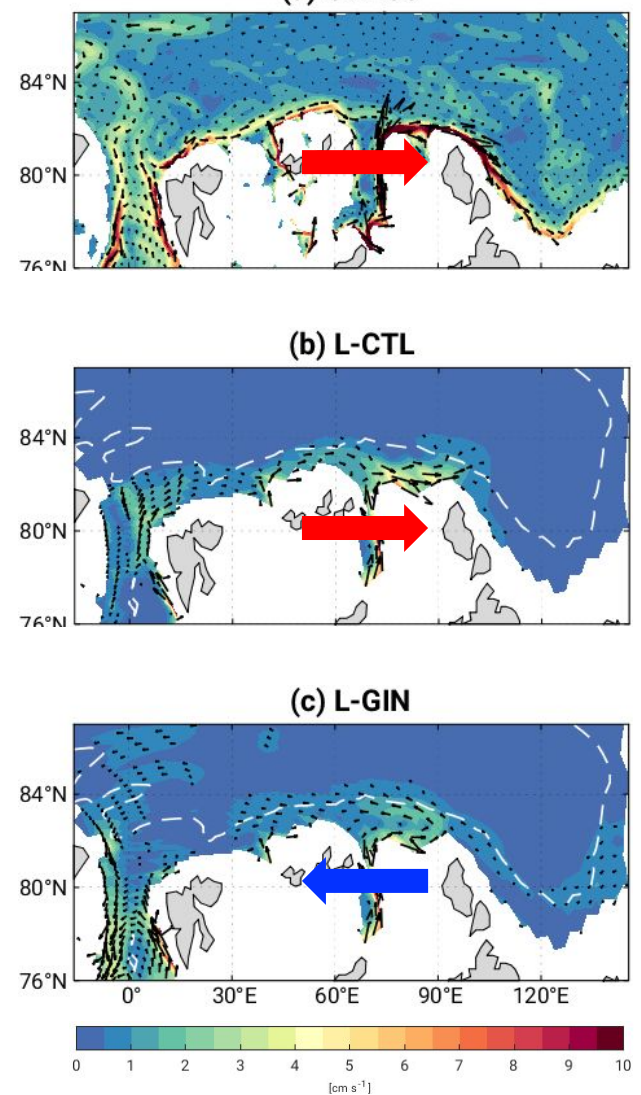
10 m < z < bottom



## Temp Bias at 300 m



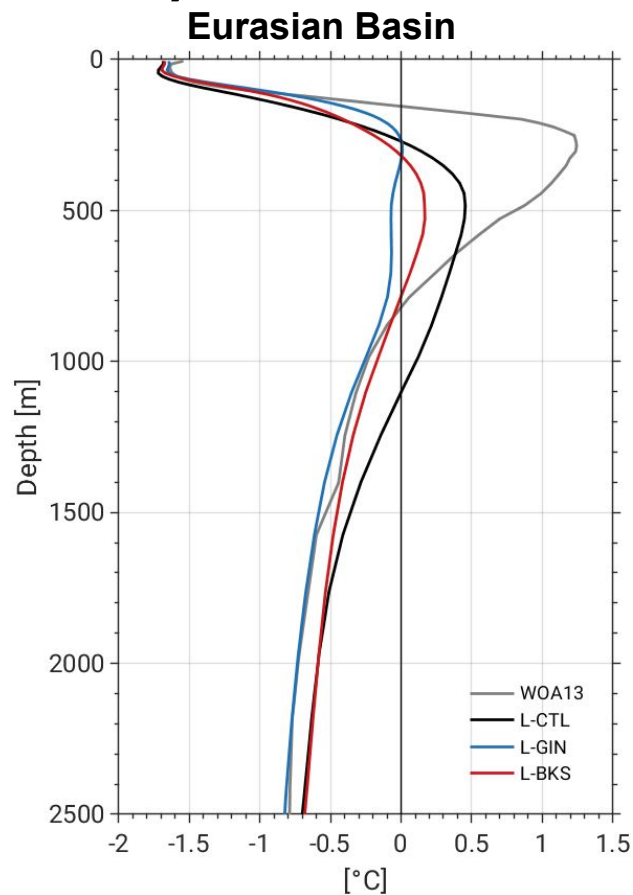
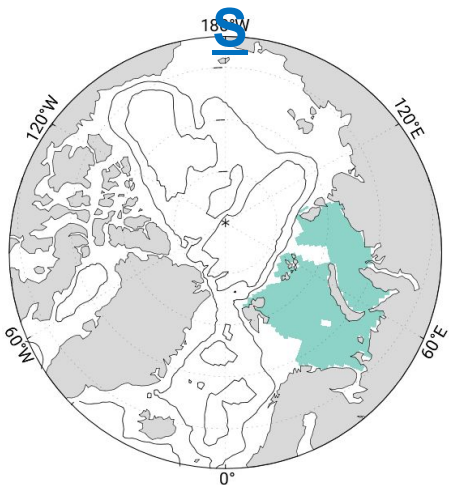
## U<sub>o</sub> at 300 m



- Inflow from the SAT is cold-biased in both L-CTL and L-GIN
- Too weak circulation in the eastern EB

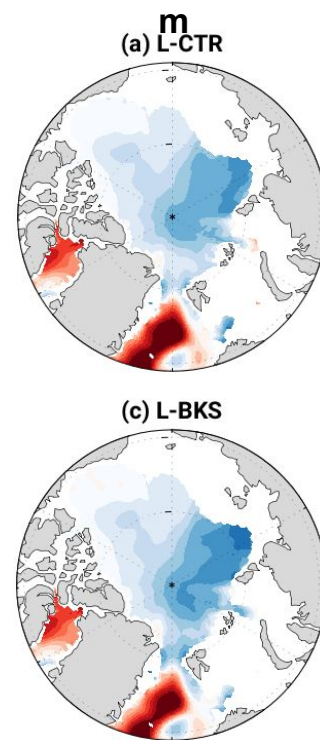
# Restoring Experiments (L-BKS)

**L-BK**

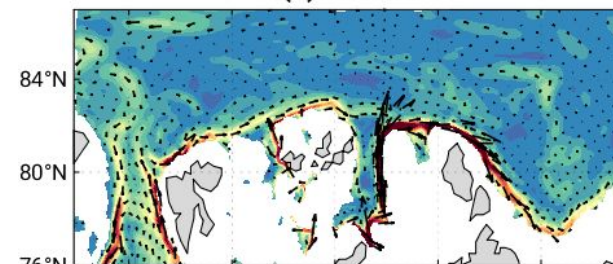


10 m < z < bottom or 500 m

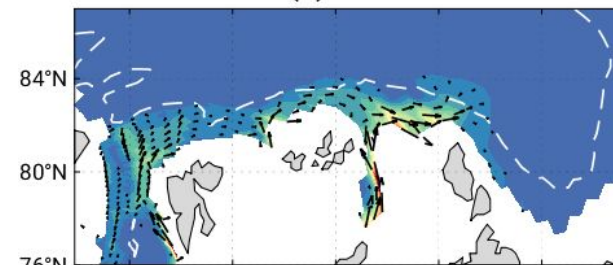
Temp Bias at 300 m



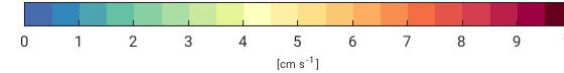
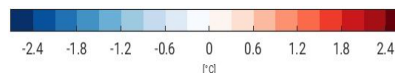
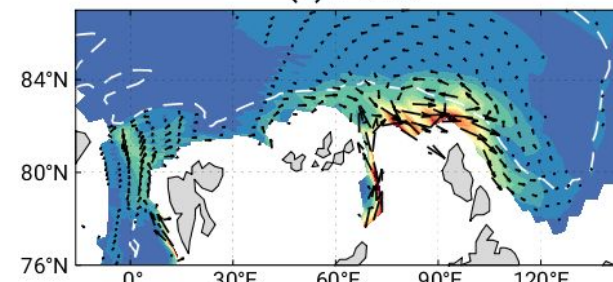
(a) ORAS5



(b) L-CTL



(d) L-BKS



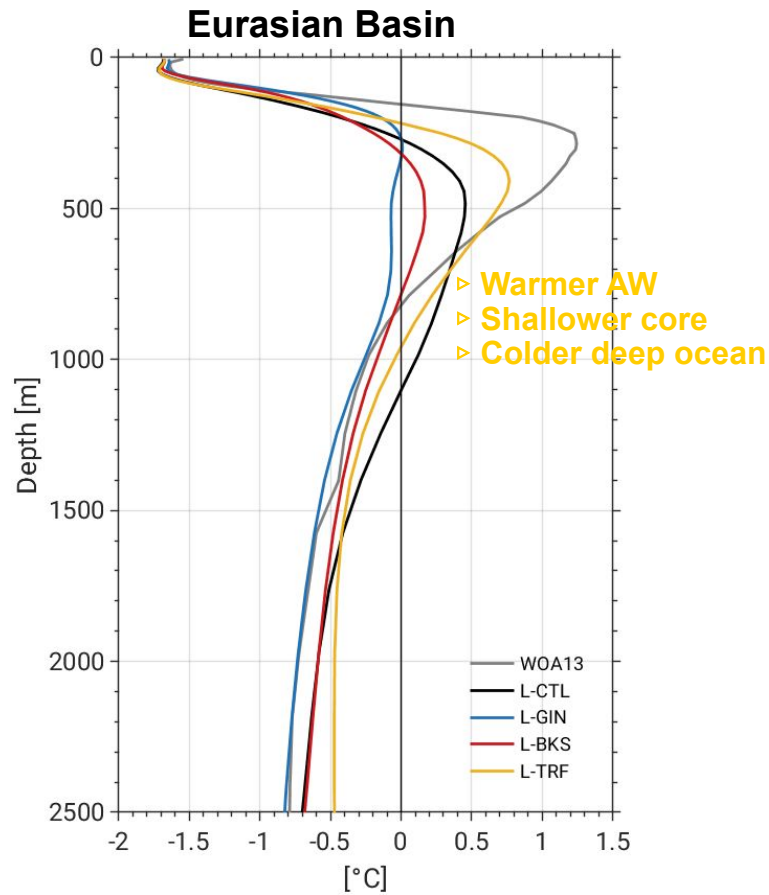
- Inflow from the SAT is still cold-biased in L-BKS
- Getting right over the shelf is not enough



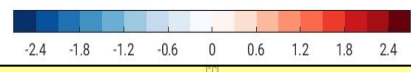
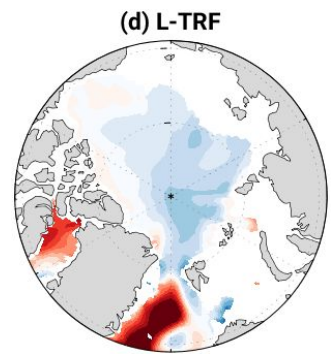
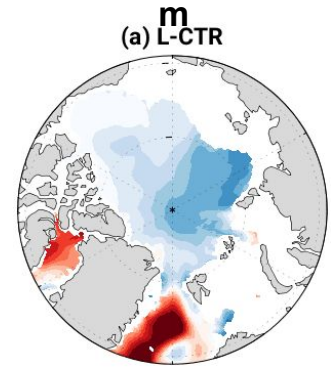
# Restoring Experiments (L-TRF)



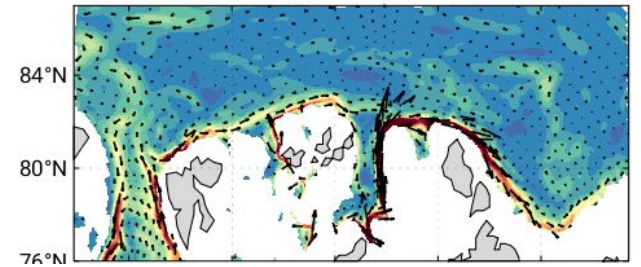
250 m < z < bottom



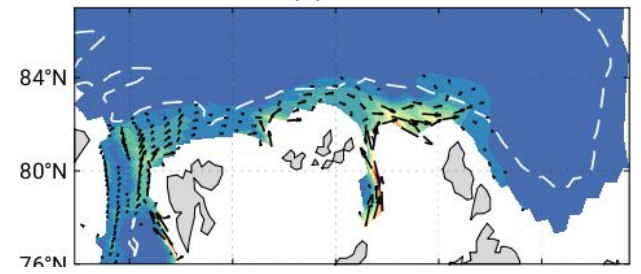
Temp Bias at 300 m



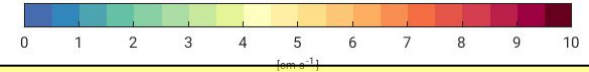
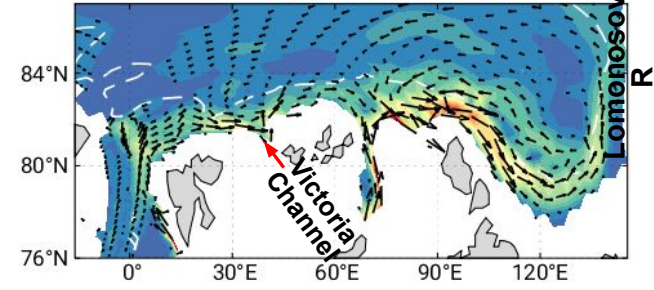
(a) ORAS5



(b) L-CTL

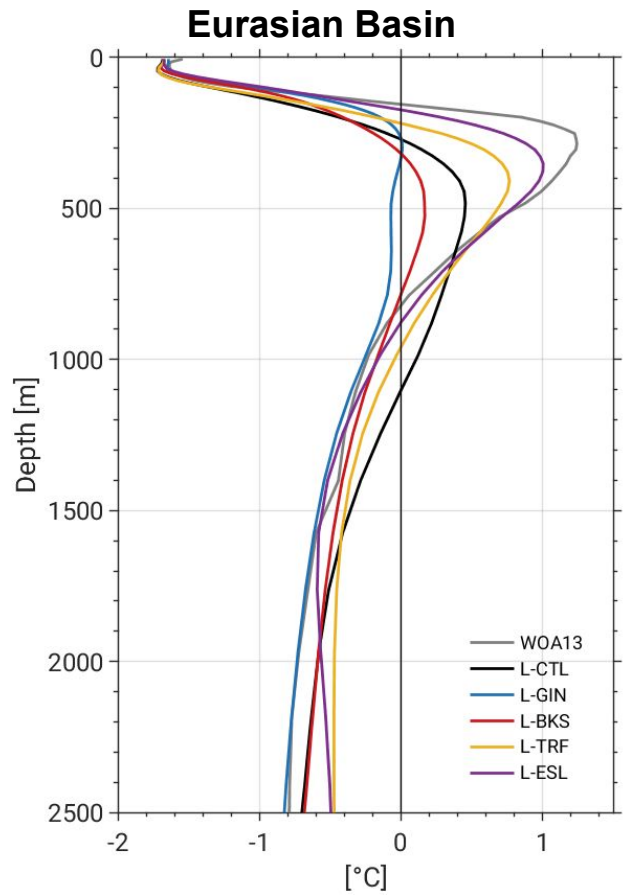


(e) L-TRF

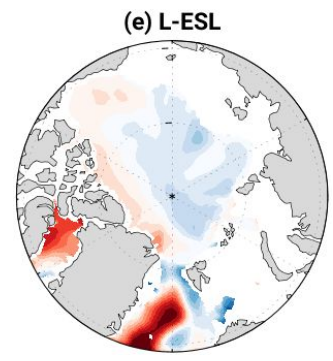
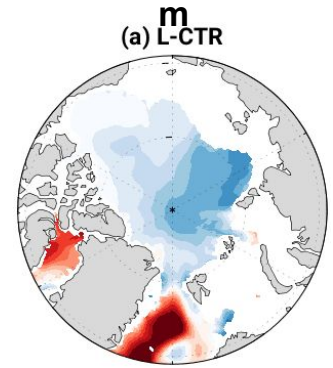


- Stronger cyclonic circulation in the eastern EB and along the Lomonosov Ridge with a stronger recirculation interior
- Suggesting the importance of the inflow from the troughs in simulating the the properties and circulation of the AW

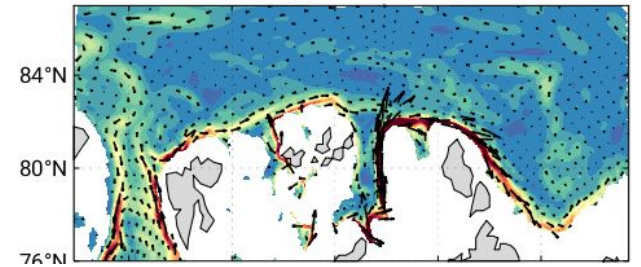
# Restoring Experiments (L-ESL)



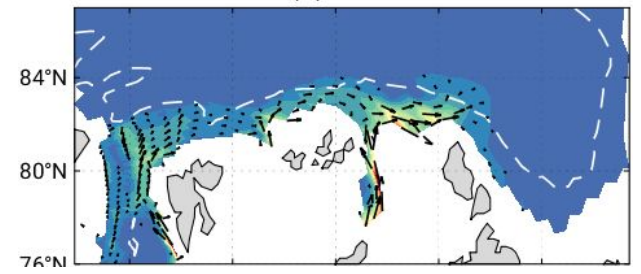
Temp Bias at 300 m



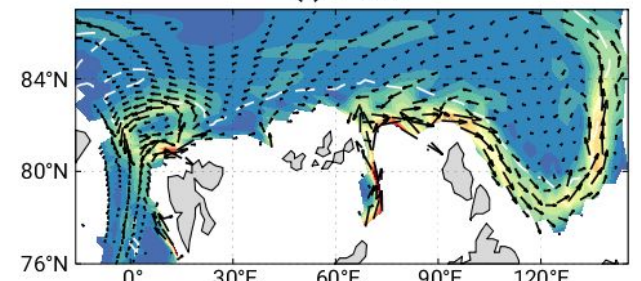
(a) ORAS5



(b) L-CTL



(f) L-ESL



# Mixing Experiments (L-WVM)

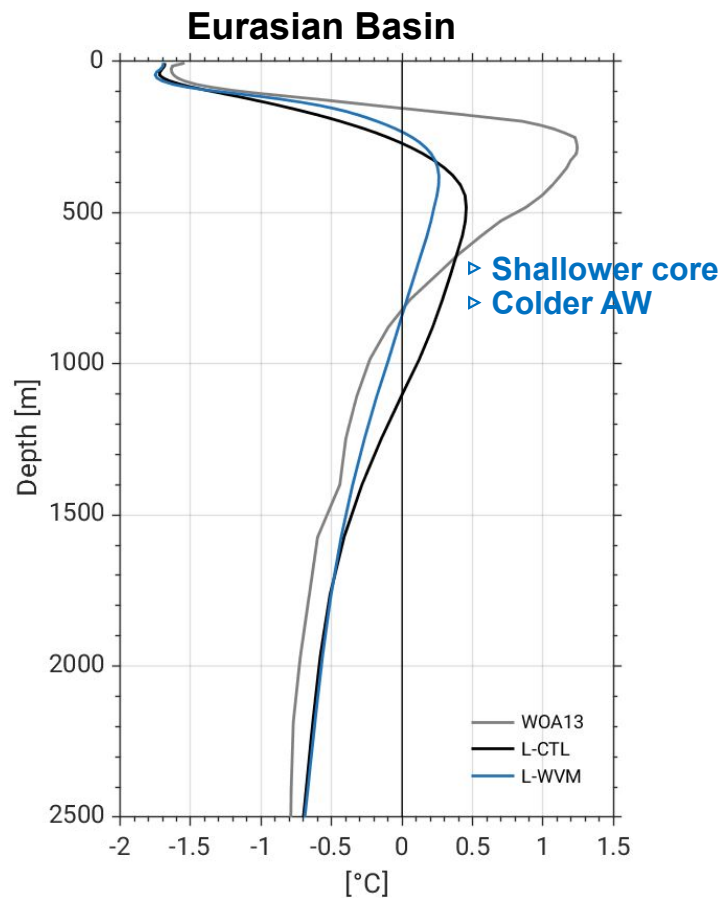
**L-WV**

**M**

$$\kappa_{v\_bg} = 1.4 \times 10^{-5} \text{ s}^{-1}$$

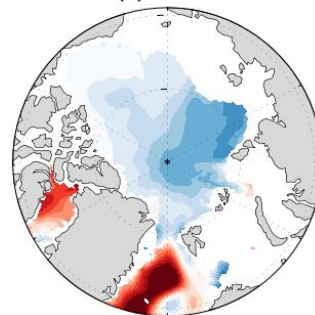
$$\rightarrow 0.1 \times 10^{-5} \text{ m}^2 \text{ s}^{-1}$$

in the Arctic

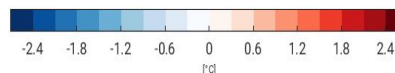
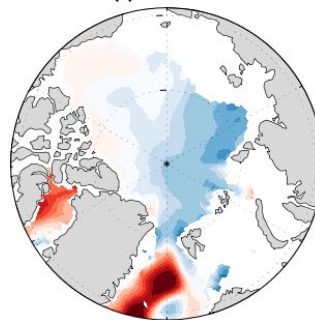


**Temp Bias at 300 m**

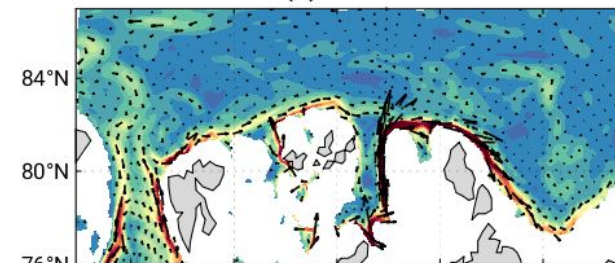
(a) L-CTR



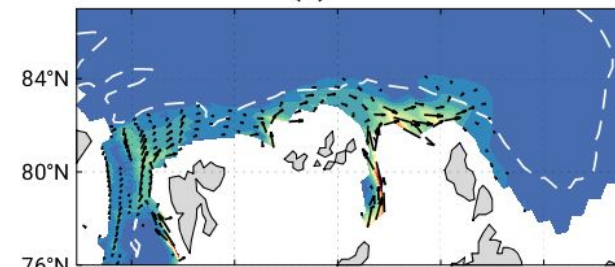
(f) L-WVM



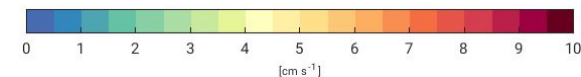
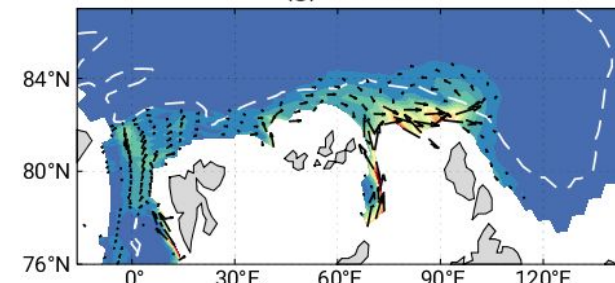
(a) ORAS5



(b) L-CTL



(g) L-WVM



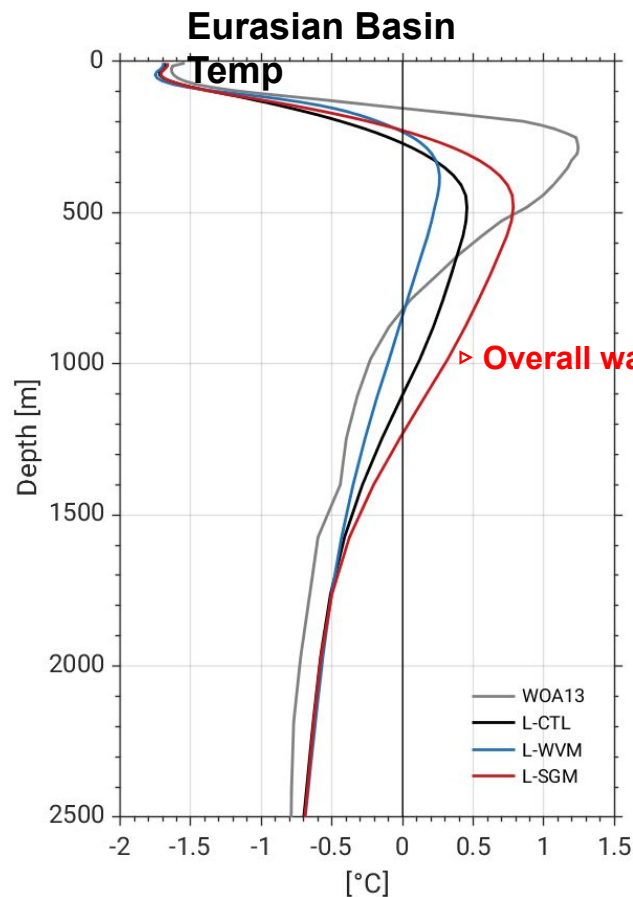
# Mixing Experiments (L-SGM)

## L-SGM

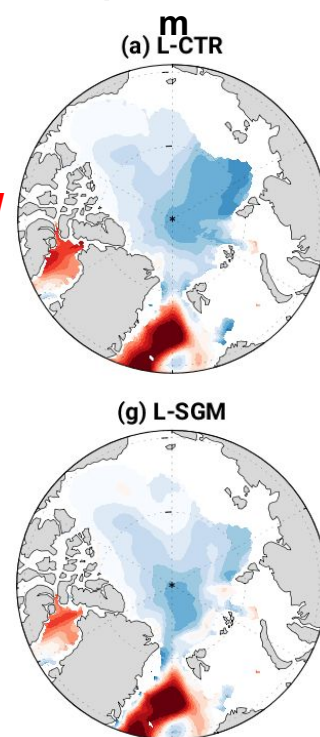
$\kappa_{GM\_isop,thic}$  scaled by area/max(area), suggested by Hunk et al. (2008)

→ smaller

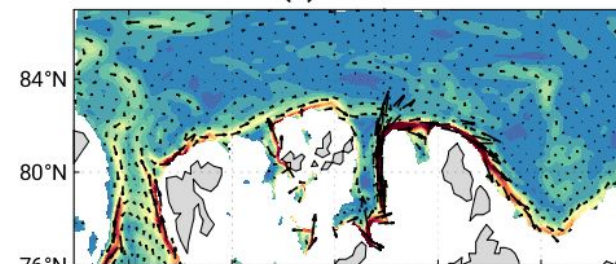
$\kappa_{GM\_isop,thic}$  in the Arctic



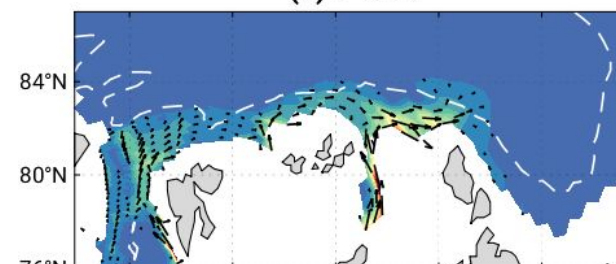
## Temp Bias at 300m



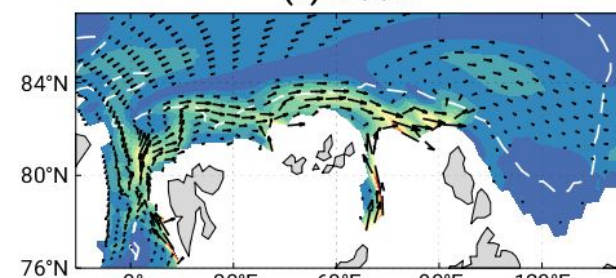
## (a) ORAS5



## (b) L-CTL



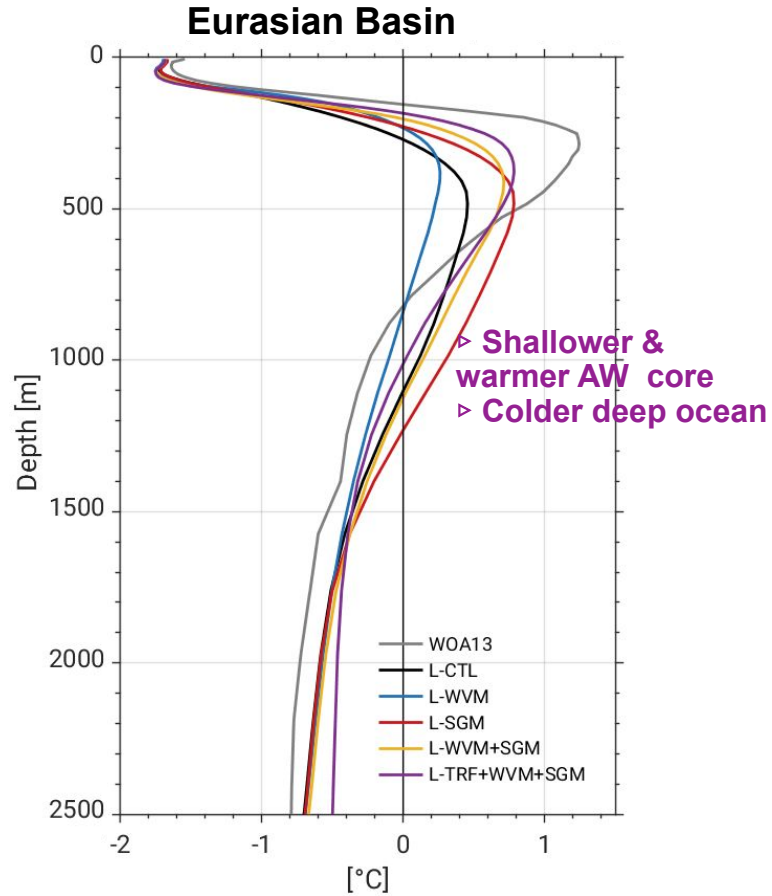
## (h) L-SGM



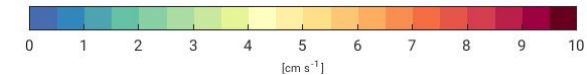
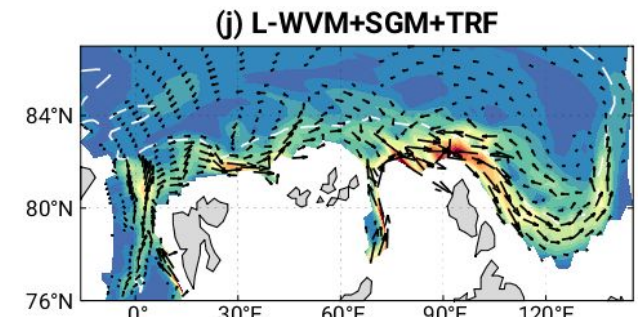
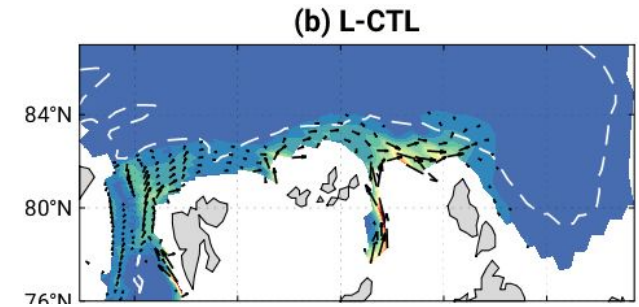
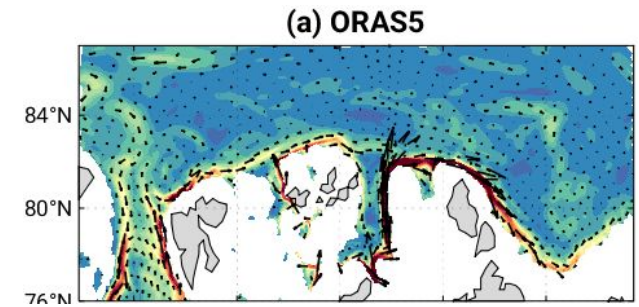
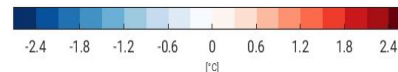
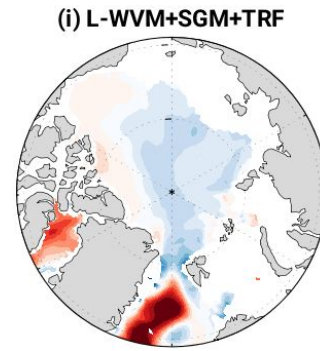
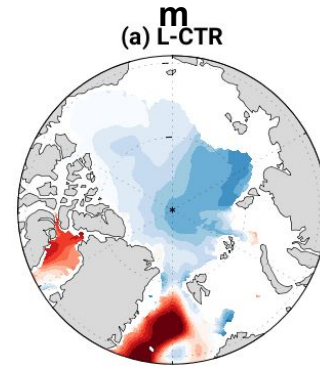
- Stronger cyclonic circulation in the western EB; still weak in eastern EB
- More heat (warm-biased) entering through FS; cold-biased inflow from the SAT

# Combined Experiments (L-WVM+SGM+TRF)

L-WVM+SGM+TR  
F



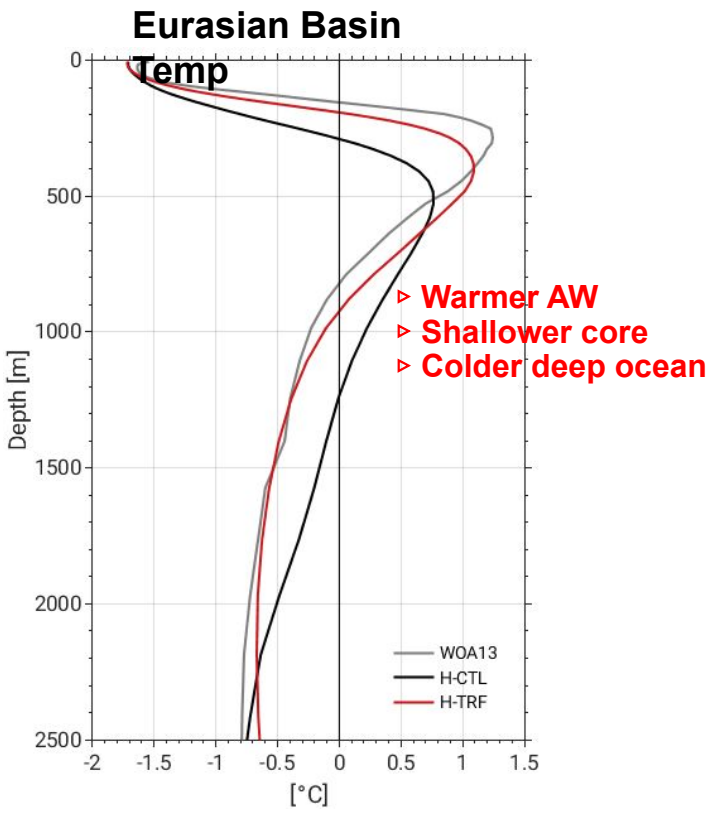
Temp Bias at 300m



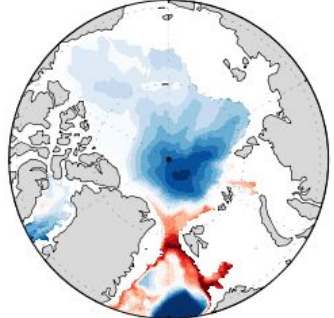
- Least biased AW core (other than L-ESL)
- Deep ocean still too warm

# HR Experiments (H-TRF)

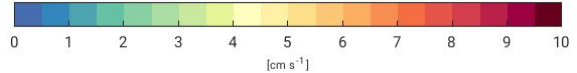
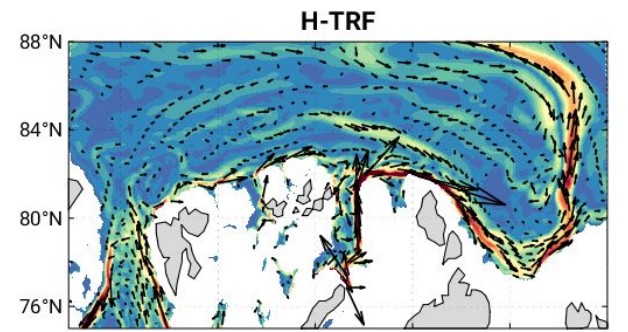
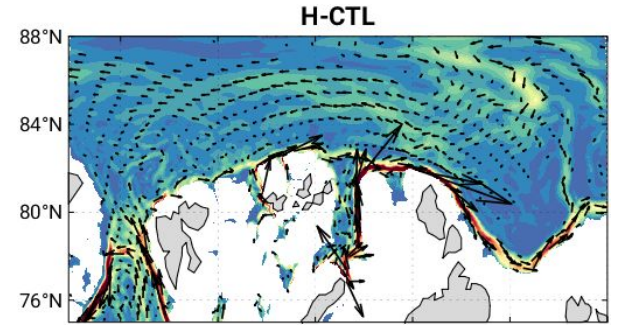
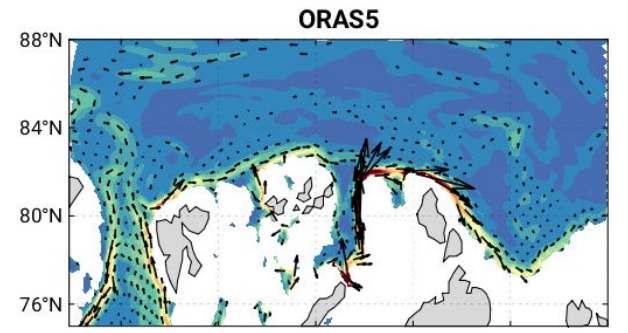
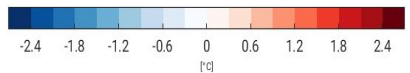
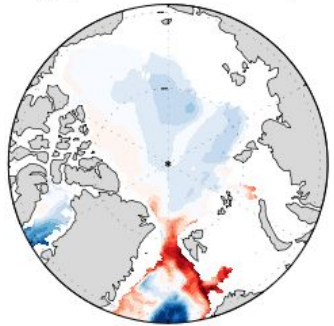
## H-TR F



Temp Bias at 300 m (b) H-CTL (305 m)

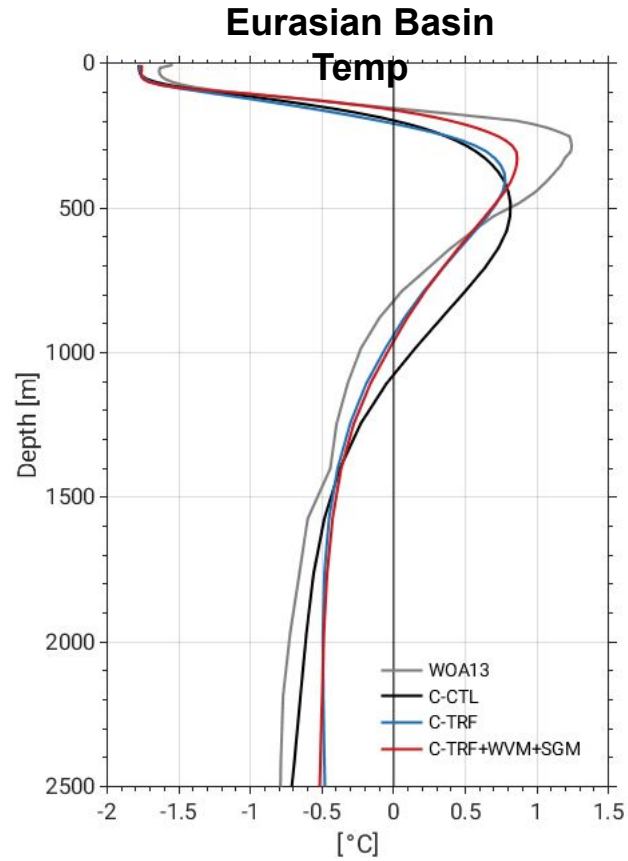


(c) H-TRF (305 m)



- Similar improvement as in L-TRF – shallower, thinner, warmer AW with a stronger cyclonic circulation in the eastern EB

# Coupled Experiments



## 1) Important role of the inflow from the troughs and channels in setting up the hydrographic properties and circulation of the AW

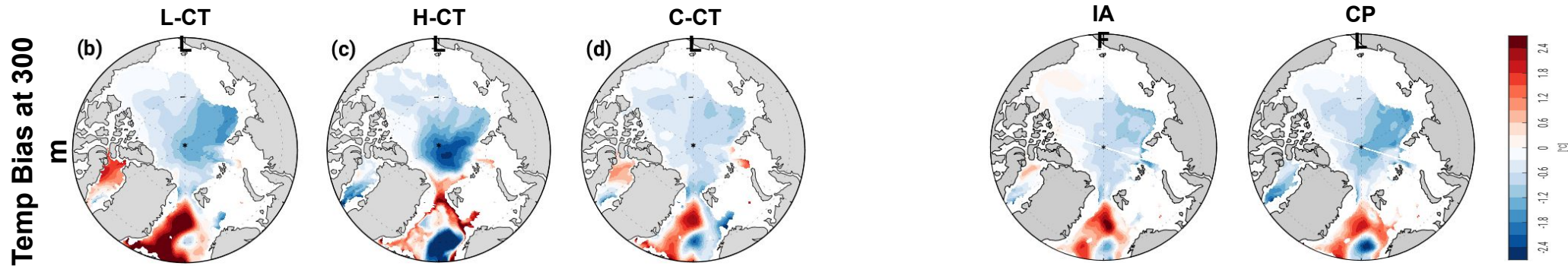
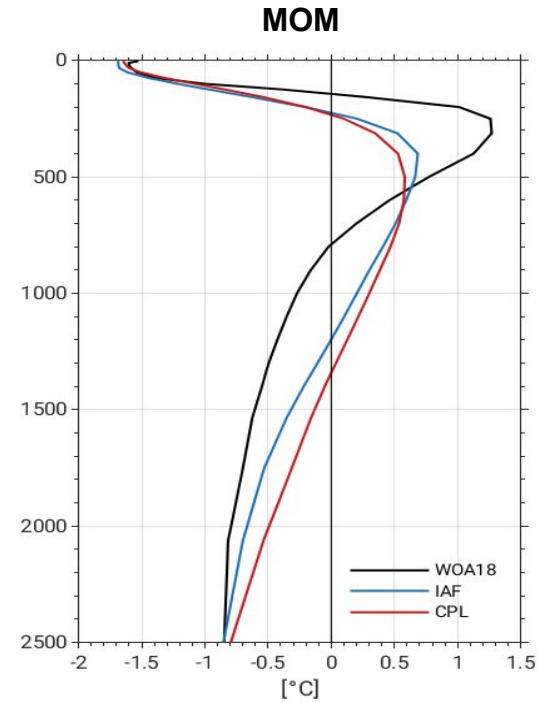
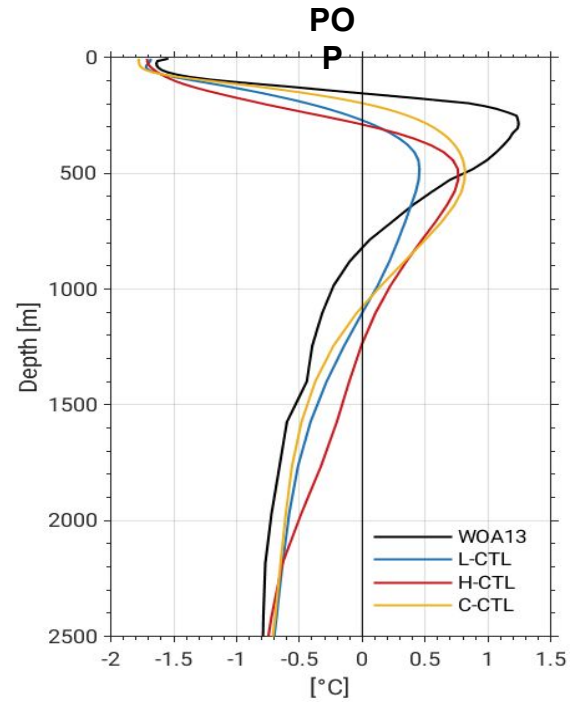
- The AW biases are substantially reduced (warmer, shallower, and thinner)
- The cyclonic circulation becomes substantially stronger in the eastern EB, which feeds recirculation in the interior
- **Calls for the development of an overflow parametrization (even in HR)**

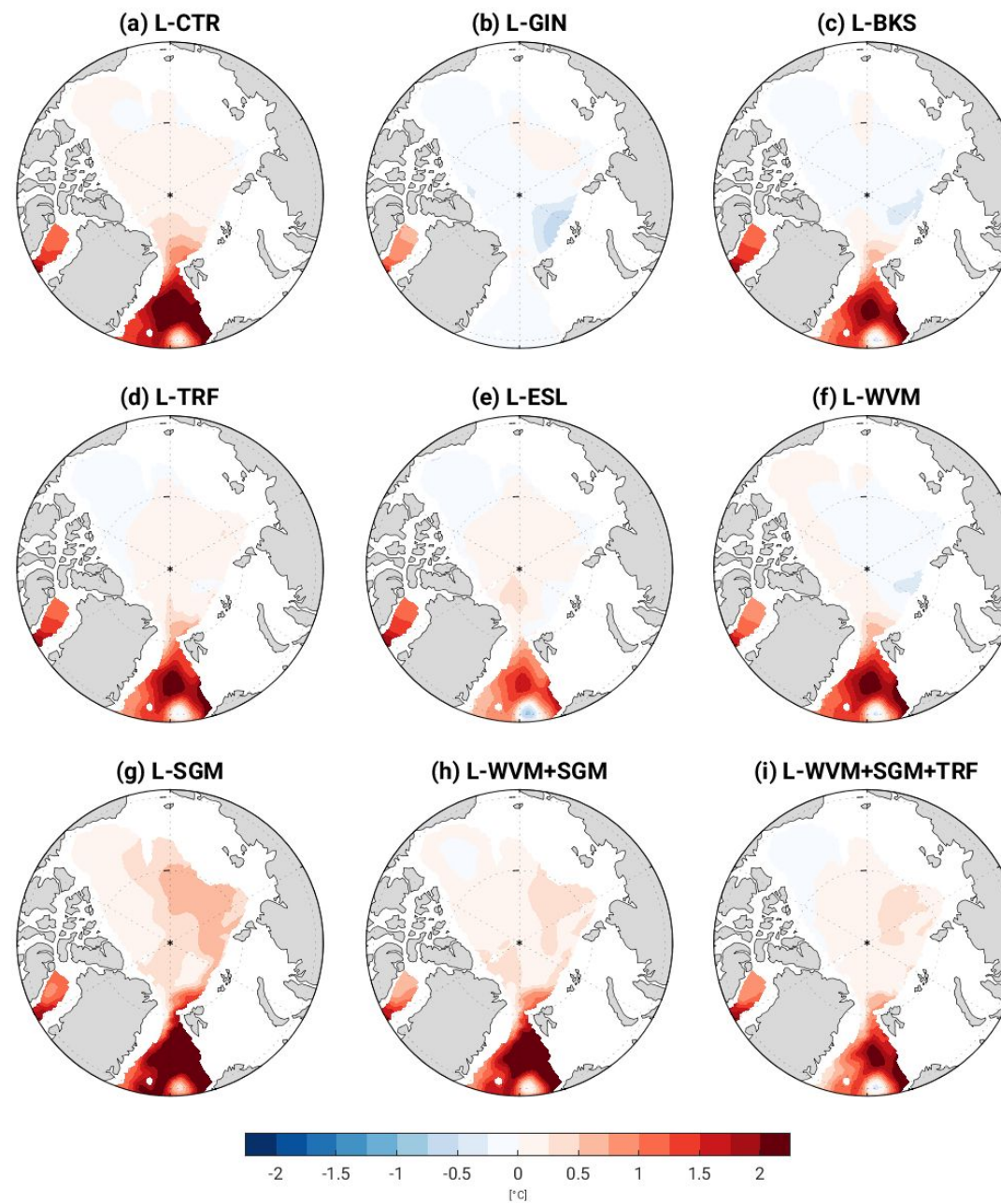
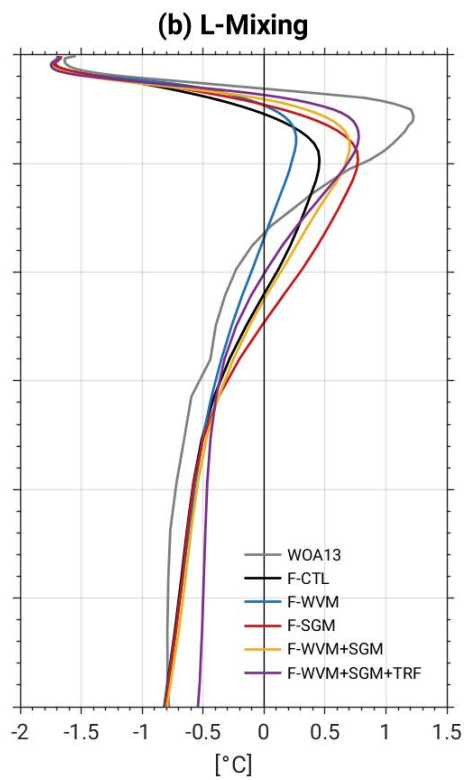
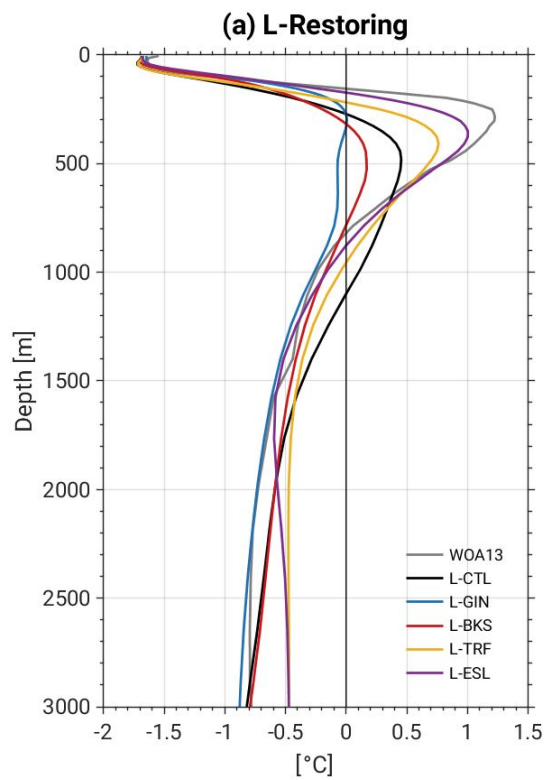
## 2) Tuning of mixing parameters (more relevant to LR)

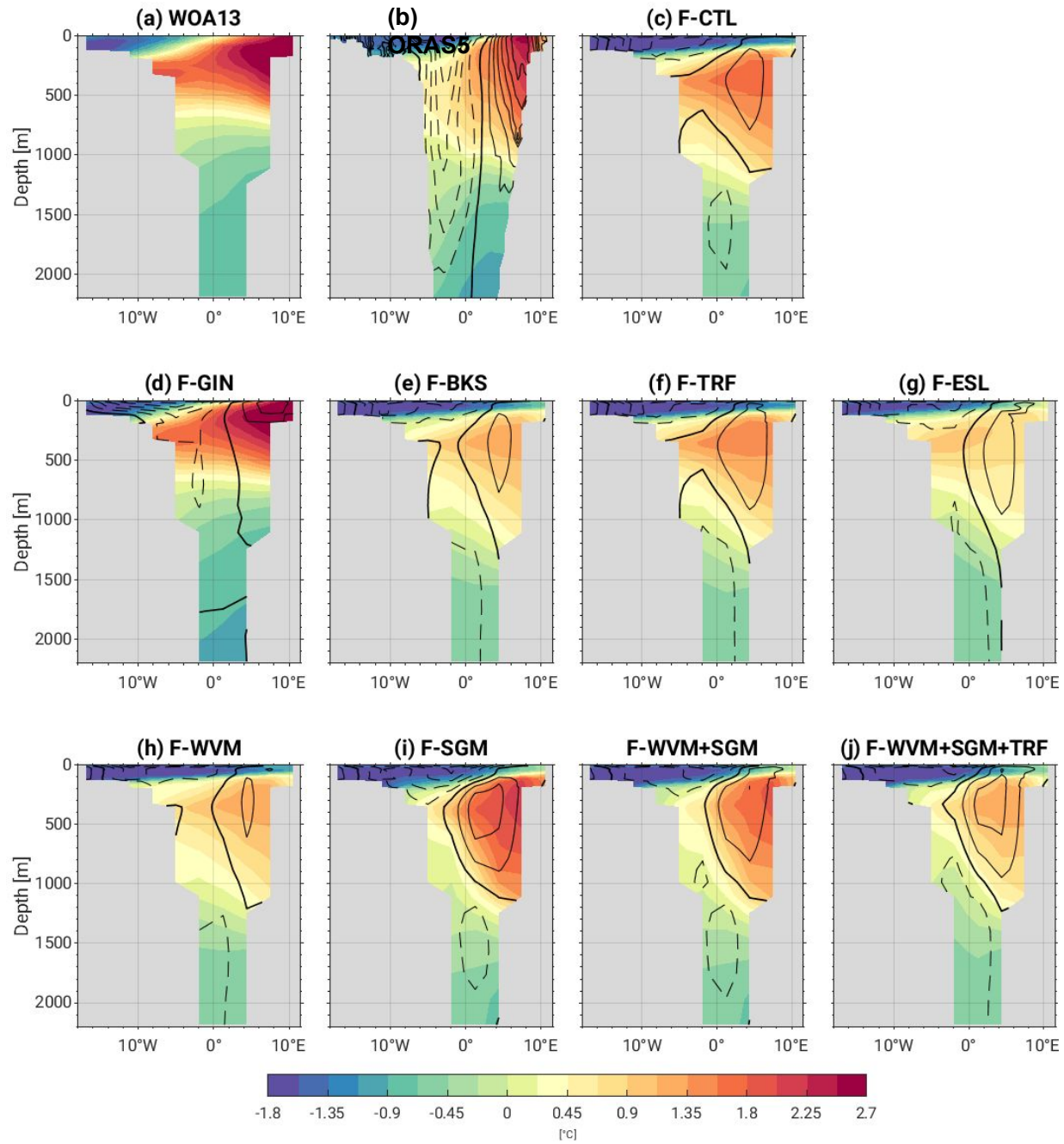
- Scaling GM diffusivity coefficients to grid area (i.e., weaker mixing in the Arctic) substantially strengthens the circulation in the western EB
- Flow pattern becomes less organized
- Biases in the GIN Seas become more matter (too warm in the deep ocean)

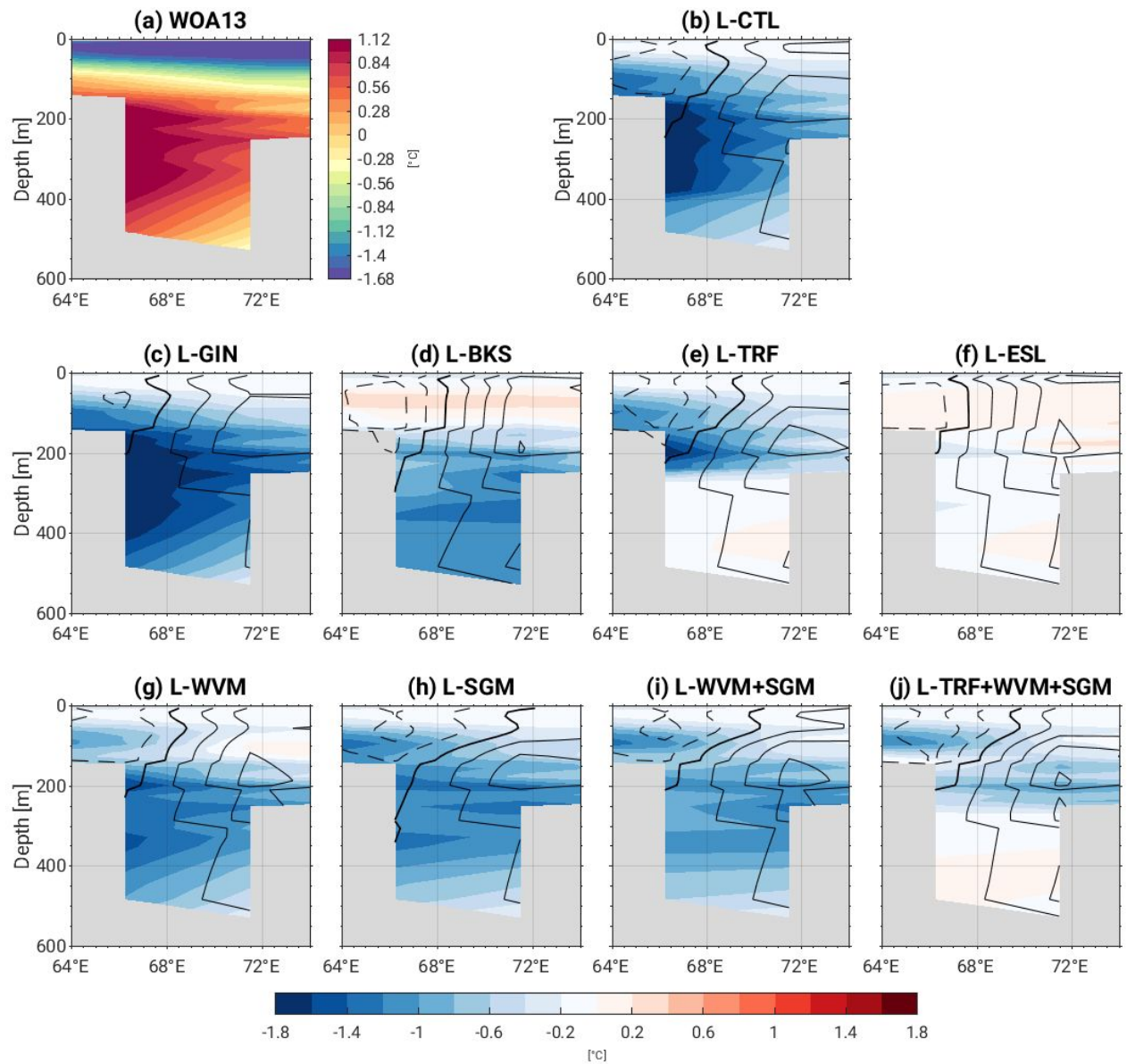


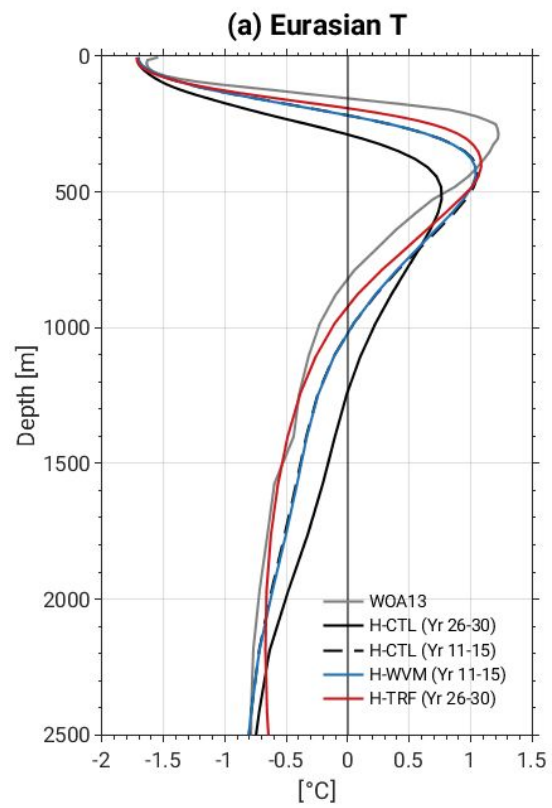
# Is MOM better than POP?



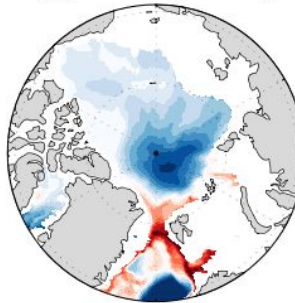




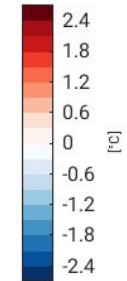
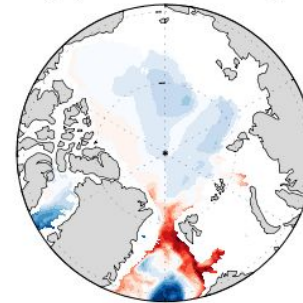




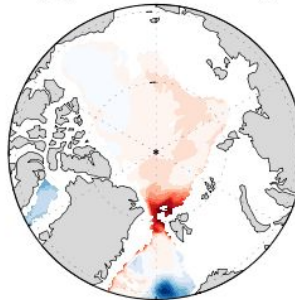
**(b) H-CTL (305 m)**



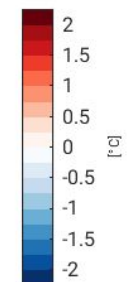
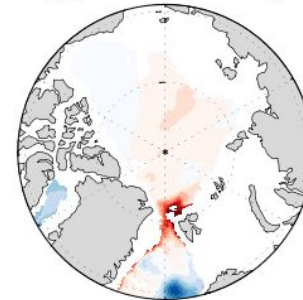
**(c) H-TRF (305 m)**

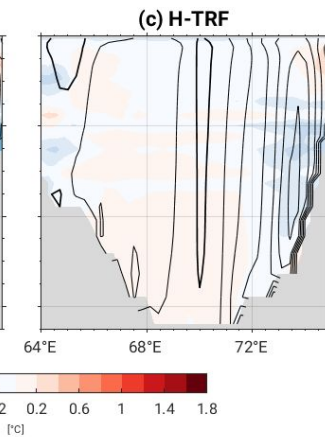
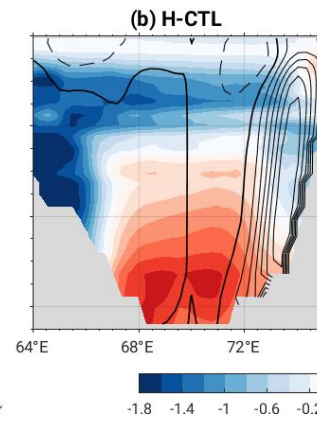
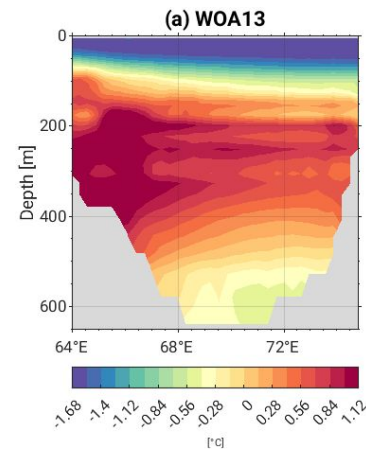
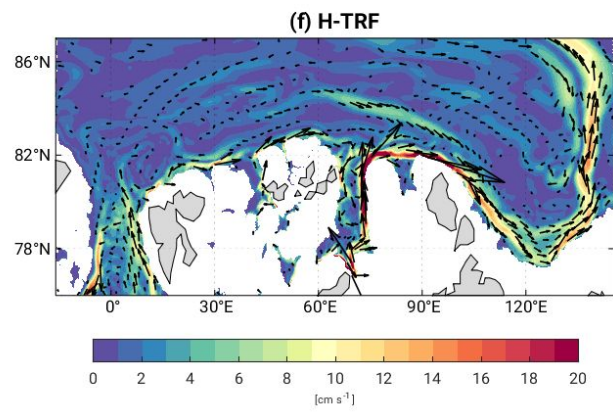
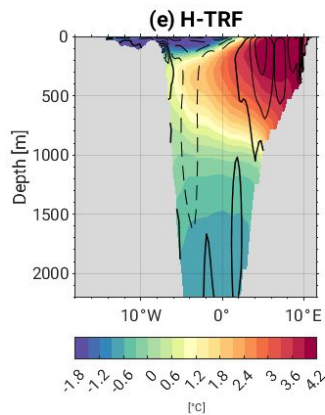
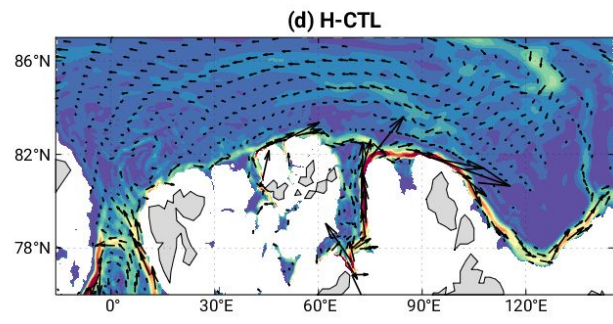
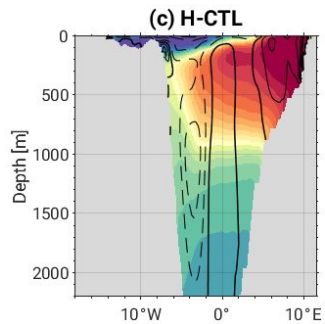
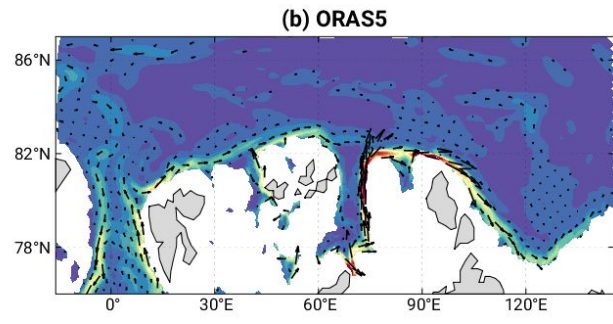
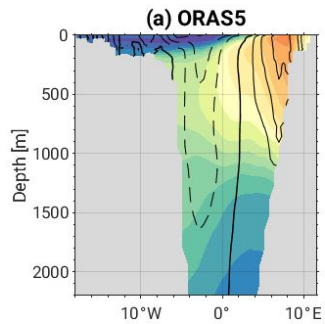


**(d) H-CTL (787 m)**



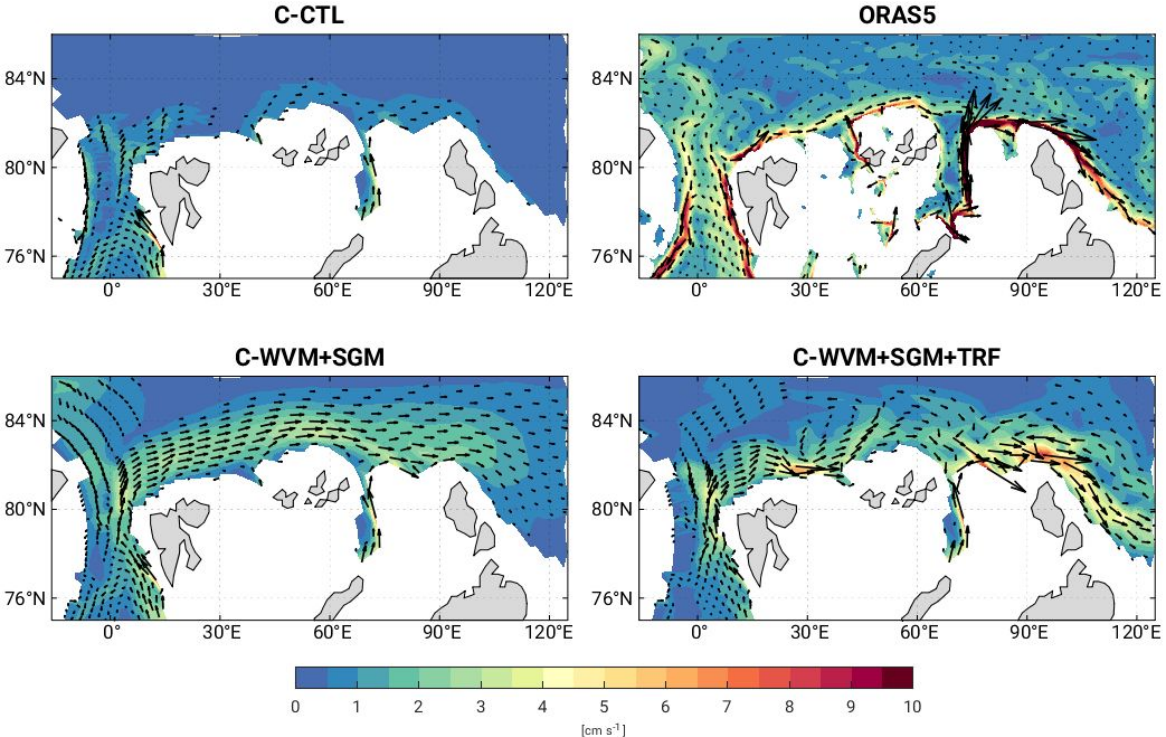
**(e) H-TRF (787 m)**





**Fig. 11. Hovmoller diagram of T bias in EB and spatial map of T bias at 300 m from coupled experiments**

- Blah



## **Tentative Conclusions:**

- Both lateral mixing and waters entering the Arctic Ocean through channels (esp., St. Anna Trough) appear to be significant sources of the AWL temperature layer bias
- TRF is most effective in reducing both cold and warm biases at the observed AW core and bottom, respectively.
- TRF also enhance the cyclonic circulation substantially
- SGM also reduce the cold bias at the AW core and enhance the cyclonic circulation substantially
- However, SGM makes the warm bias worse at depths
- Need to optimize lateral mixing
- Need to improve water properties that exit the troughs (overflow parameterization?)



