

# Isopycnal mixing and backscatter parameterizations

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# Parameterization at coarse resolution

## Advective transport

(Gent and McWilliams, 1990)

$$\mathbf{u}_* = -\partial_z(\kappa_{GM}\mathbf{s}),$$

$$w_* = \nabla_h \cdot (\kappa_{GM}\mathbf{s})$$

$$\mathbf{s} = -\nabla_h \rho / \partial_z \rho$$

$$A_{GM} = \kappa_{GM} \begin{bmatrix} 0 & 0 & -s_x \\ 0 & 0 & -s_y \\ s_x & s_y & 0 \end{bmatrix}$$

- adiabatic slumping of isopycnals
  - eddy-induced transport velocity
  - GM coefficient (“thickness diffusivity”)
-

# Parameterization at coarse resolution

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- eddy-induced transport velocity
- GM coefficient (“thickness diffusivity”)

## Diffusive mixing

(Redi, 1982)

$$\nabla \cdot (S_{\text{R}} \nabla \bar{c})$$

where

$$S_{\text{R}} = \kappa_{\text{R}} \begin{bmatrix} 1 & 0 & s_x \\ 0 & 1 & s_y \\ s_x & s_y & |\mathbf{s}|^2 \end{bmatrix}$$

- eddy stirring enhances dissipation of tracer variance
- downgradient diffusion of tracers on isopycnals
- Redi coefficient (isopycnal tracer diffusivity)

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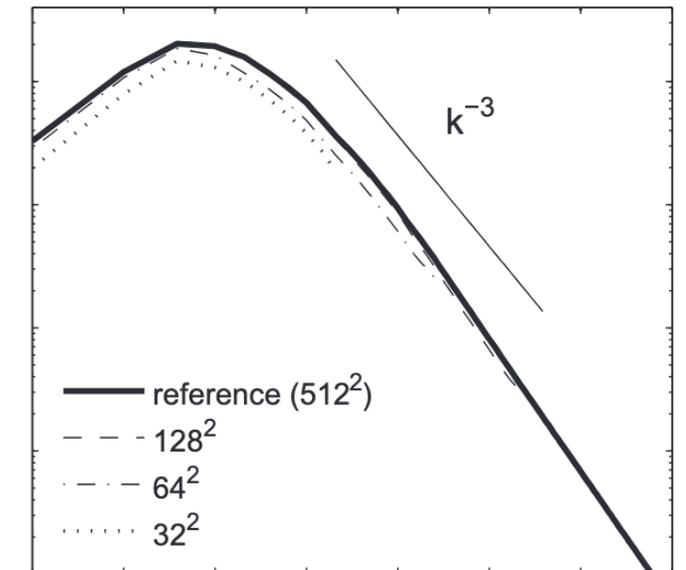
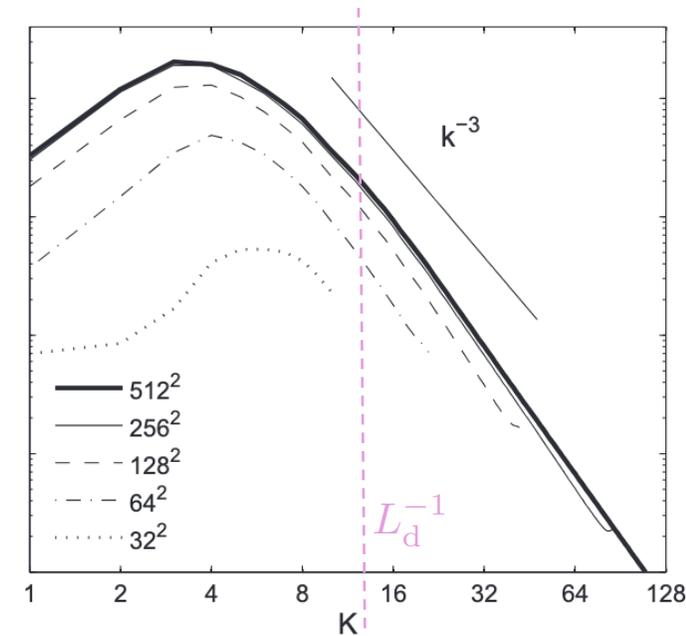
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$$\partial_t \bar{c} + \bar{\mathbf{v}} \cdot \nabla \bar{c} - \nabla \cdot (A_{\text{GM}} \nabla \bar{c}) = \nabla \cdot (S_{\text{R}} \nabla \bar{c}) + \text{source/sink}$$

# Parameterization at eddy-permitting resolution

- GM-Redi designed for non-eddy-permitting resolutions
- Mesoscale turbulence generated by baroclinic instability **partially resolved**
  - resolution near deformation radius
    - dissipation cuts off energy source for inverse cascade
- Energy backscatter to re-energize partially resolved eddies using negative viscosity



*Jansen and Held (2014)*

# Parameterization at eddy-permitting resolution

Backscatter in more realistic set-ups:

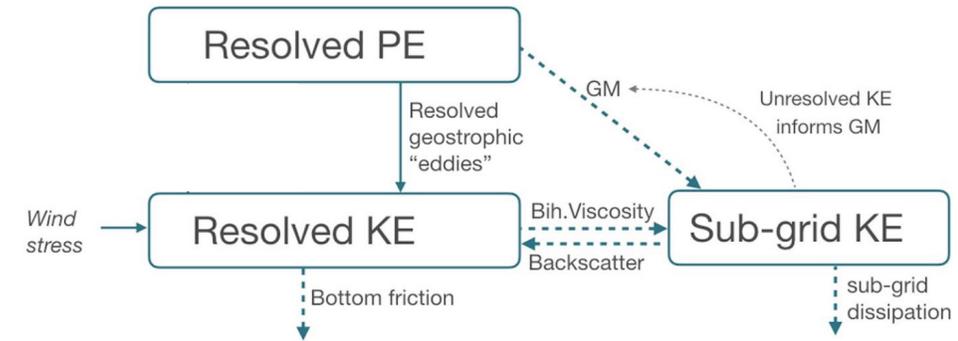
$$\partial_t \mathbf{u} + (f + \zeta) \hat{\mathbf{z}} \times \mathbf{u} + \nabla(K + M) = \frac{1}{\rho_0} \partial_z \boldsymbol{\tau} - \nabla \cdot [\nu_4 \nabla(\nabla^2 \mathbf{u})] + \nu_2 \nabla^2 \mathbf{u}$$

↓ Dissipation
↑ Backscatter

- negative viscosity from prognostic subgrid KE budget

$$\frac{\partial e}{\partial t} = \dot{e} = \dot{e}_{\text{GM}} + \dot{e}_{\text{Smag}} - \dot{e}_{\text{BScat}} - \dot{e}_{\text{diss}} - \dot{e}_{\text{adv}}$$

$$\nu_2 = c\sqrt{2e}L_{\text{mix}}, \quad e = e(x, y, t)$$



# Parameterization at eddy-permitting resolution

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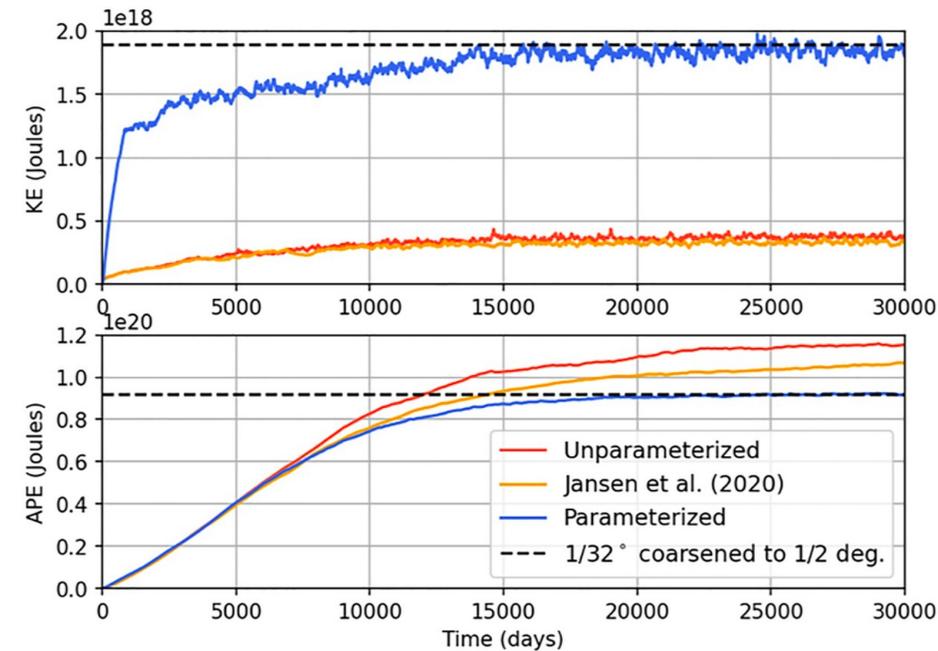
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- negative viscosity from prognostic subgrid KE budget

Backscatter **with no GM** can get both the kinetic energy *and* the stratification to match a hi-res truth  
(Yankovsky et al. 2024)

$$\frac{\partial e}{\partial t} = \dot{e} = \dot{e}_{\text{GM}} + \dot{e}_{\text{Smag}} - \dot{e}_{\text{BScat}} - \dot{e}_{\text{diss}} - \dot{e}_{\text{adv}}$$

$\nu_2 = c\sqrt{2e}L_{\text{mix}}, \quad e = e(x, y, t)$



Yankovsky et al. (2024)

# Hypothesis & goal

Hypothesis:

That backscatter induces sufficient  
isopycnal mixing



Goal:

Measure mixing in model  
simulations

# Experimental set-up

- Adiabatic primitive equations on sphere
- MOM6, NeverWorld2 configuration (Marques et al. 2022)

→ 15 isopycnal layers

→ constant wind stress at surface

→ meridional ridge and continental slopes

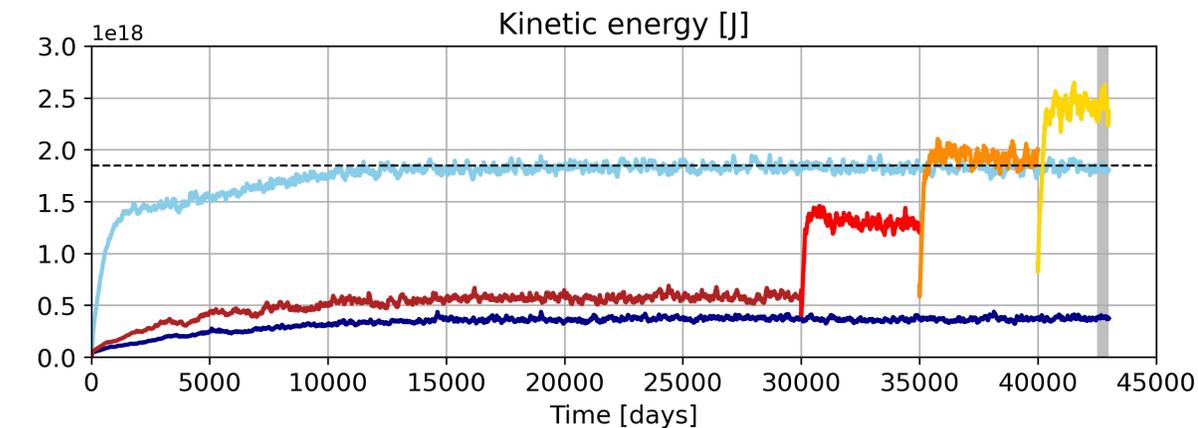
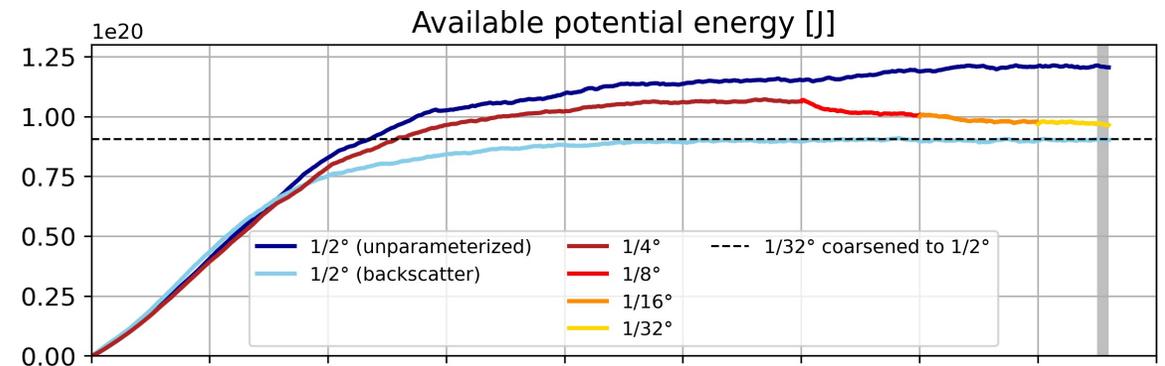
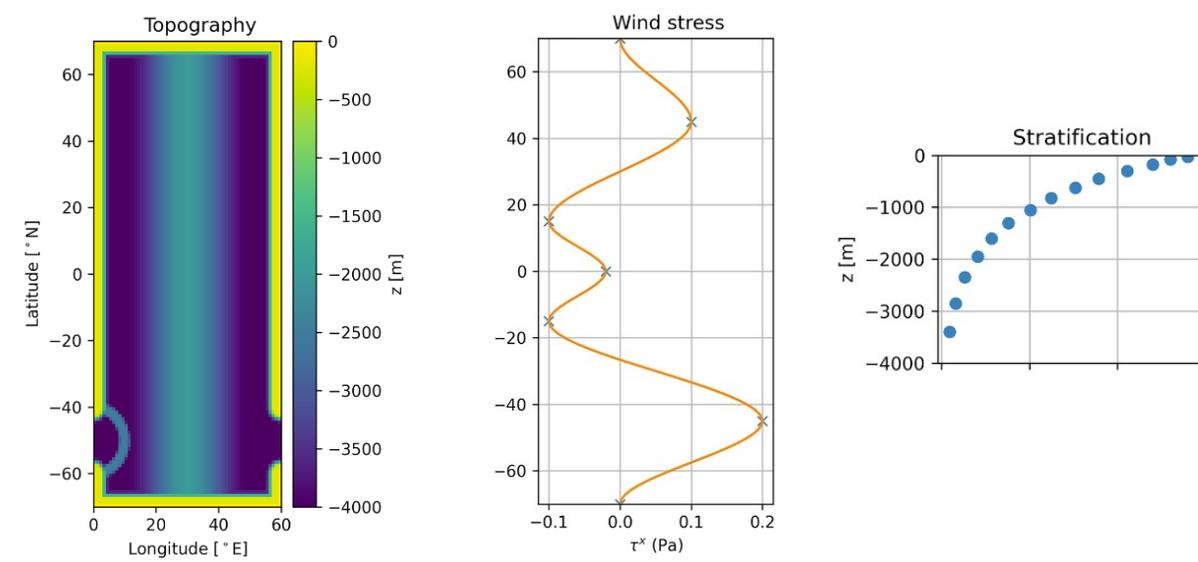
→ periodic southern channel

→ horizontal resolution:

- $1/2^\circ$  (~50 km), “eddy-permitting”
- $1/32^\circ$  (~3 km), “eddy-resolving”

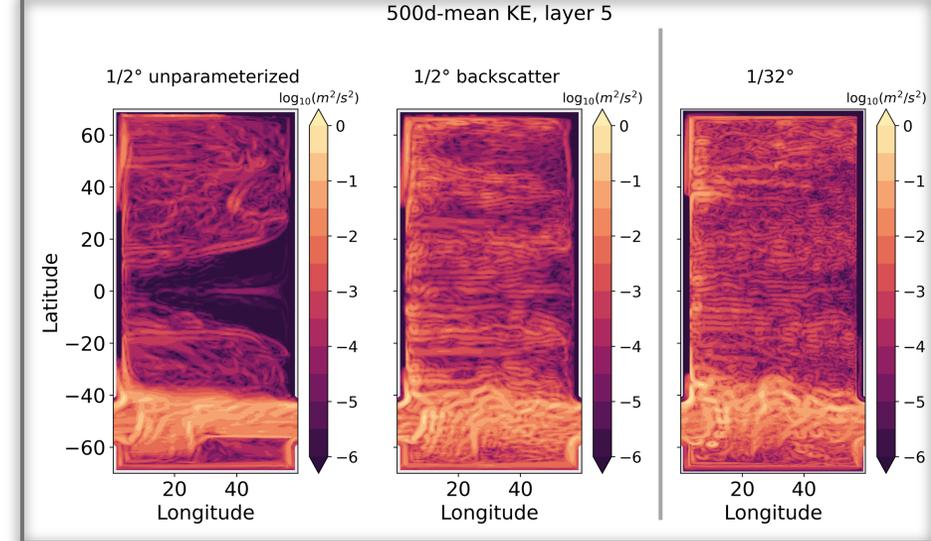
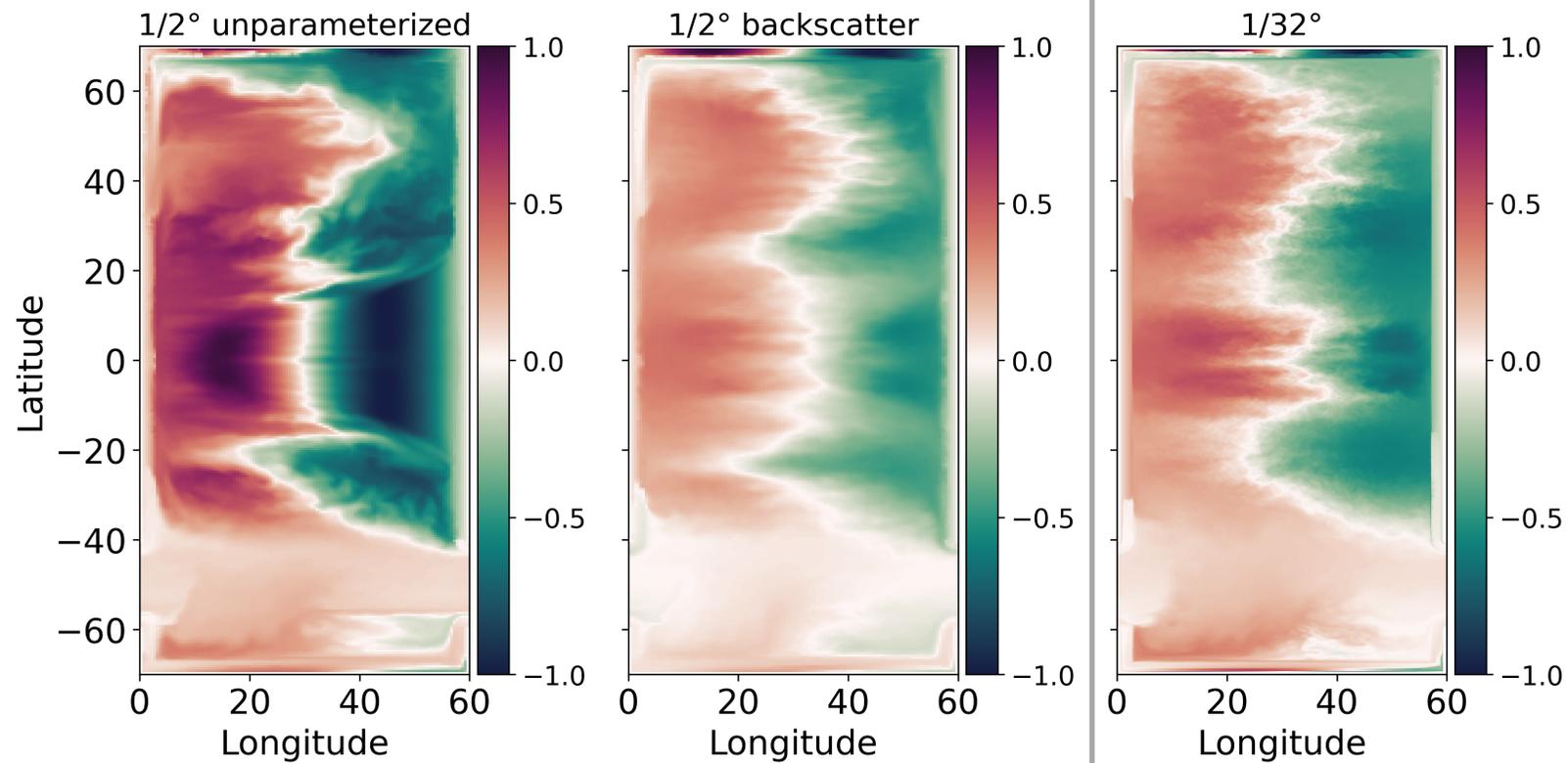
→ backscatter with **no GM** (Yankovsky et al. 2024)

→ passive tracers advected online (forced-dissipative equilibrium)



# Tracer distributions

500d-mean  $\sin(x)$  tracer distribution, layer 5



# Measuring eddy tracer fluxes

$$\hat{c} \equiv \frac{\overline{hc}}{\overline{h}} \quad c'' \equiv c - \hat{c}$$

$$\begin{bmatrix} \widehat{u''c''} \\ \widehat{v''c''} \end{bmatrix} = - \begin{bmatrix} K^{11} & K^{12} \\ K^{21} & K^{22} \end{bmatrix} \begin{bmatrix} \partial_x \hat{c} \\ \partial_y \hat{c} \end{bmatrix}$$

# Measuring eddy tracer fluxes

Method of Multiple Tracers (MMT)

$$\begin{bmatrix} \widehat{u''c''} \\ \widehat{v''c''} \end{bmatrix} = - \begin{bmatrix} K^{11} & K^{12} \\ K^{21} & K^{22} \end{bmatrix} \begin{bmatrix} \partial_x \widehat{c} \\ \partial_y \widehat{c} \end{bmatrix}$$

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## Method of Multiple Tracers (MMT)

- advect  $n$  tracers,  $c_j(x, y, z, t)$ ,  $j = 1, \dots, n$
- linearly restore to target fields:  $\partial_t c_j + \dots = -(c_j - c_j^*)/\tau_j$
- define 2 x 2 mixing tensor (at each  $x, y, z$ )

$$-K \nabla \widehat{c}_j \equiv \widehat{\mathbf{u}''c''_j}, \quad K(x, y, z) \in \mathbb{R}^{2 \times 2}$$

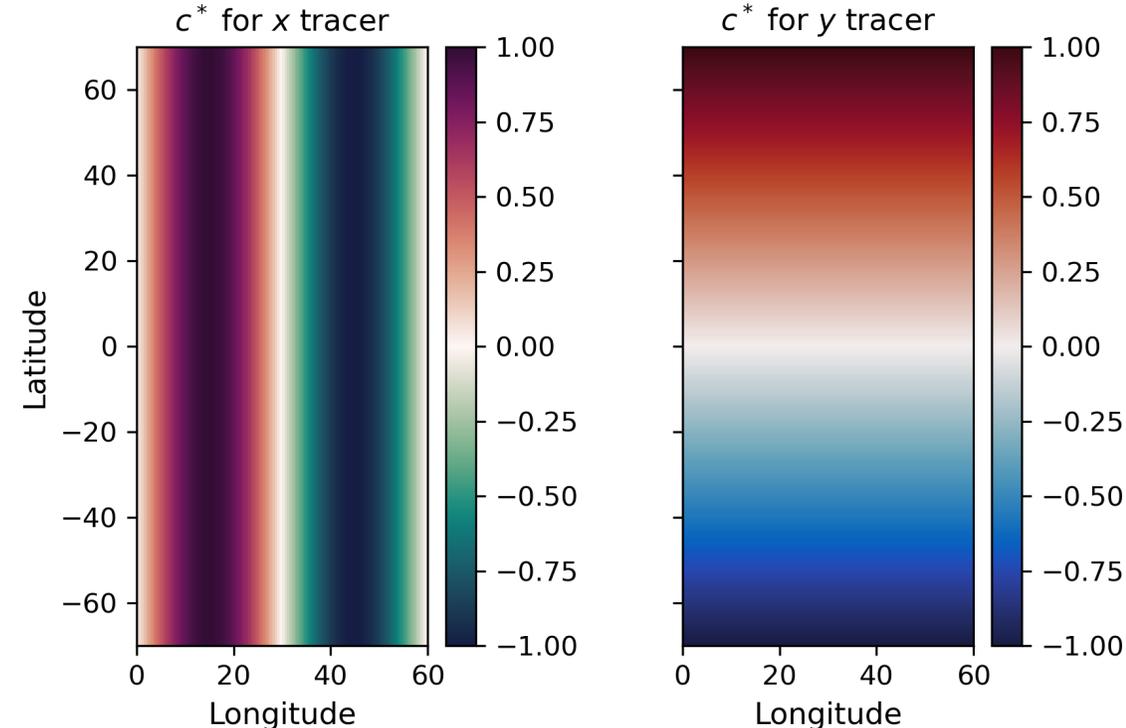
- write as over-determined matrix equation

$$-G^T K^T = F^T, \quad G, F \in \mathbb{R}^{2 \times n}, \quad n > 2$$

- optimal solution from least squares

$$K_{\text{lsq}}^T = -(GG^T)^{-1}GF^T$$

$$K \simeq K_{\text{lsq}}$$



# Tracer mixing

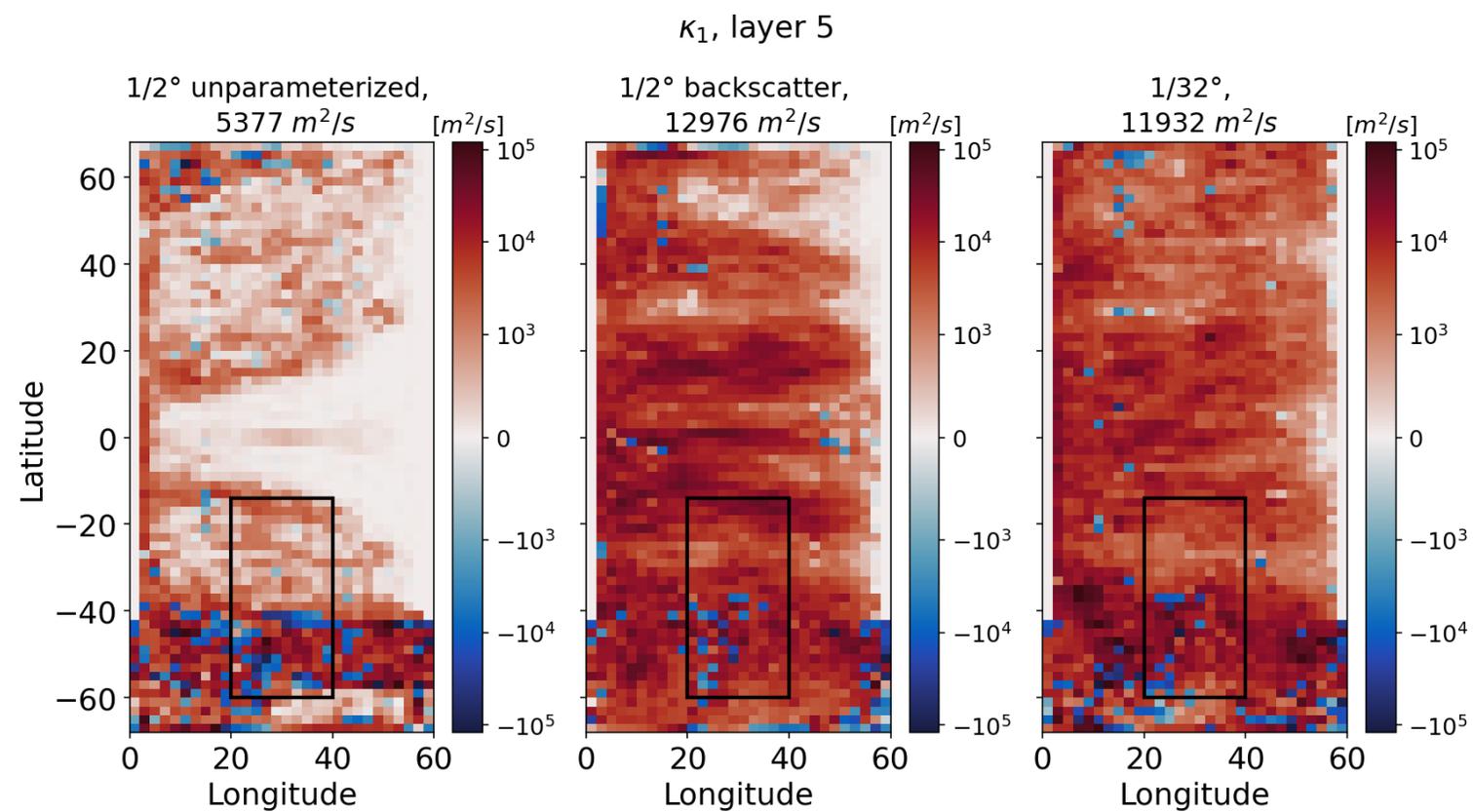
$$A = \frac{1}{2}(K - K^T)$$

$$S = \frac{1}{2}(K + K^T)$$

$$\implies S = U \begin{bmatrix} \kappa_1 & 0 \\ 0 & \kappa_2 \end{bmatrix} U^T$$

- Only symmetric part contributes to irreversible mixing (cf. Redi)

→ eigenvalues set magnitude



# Tracer mixing

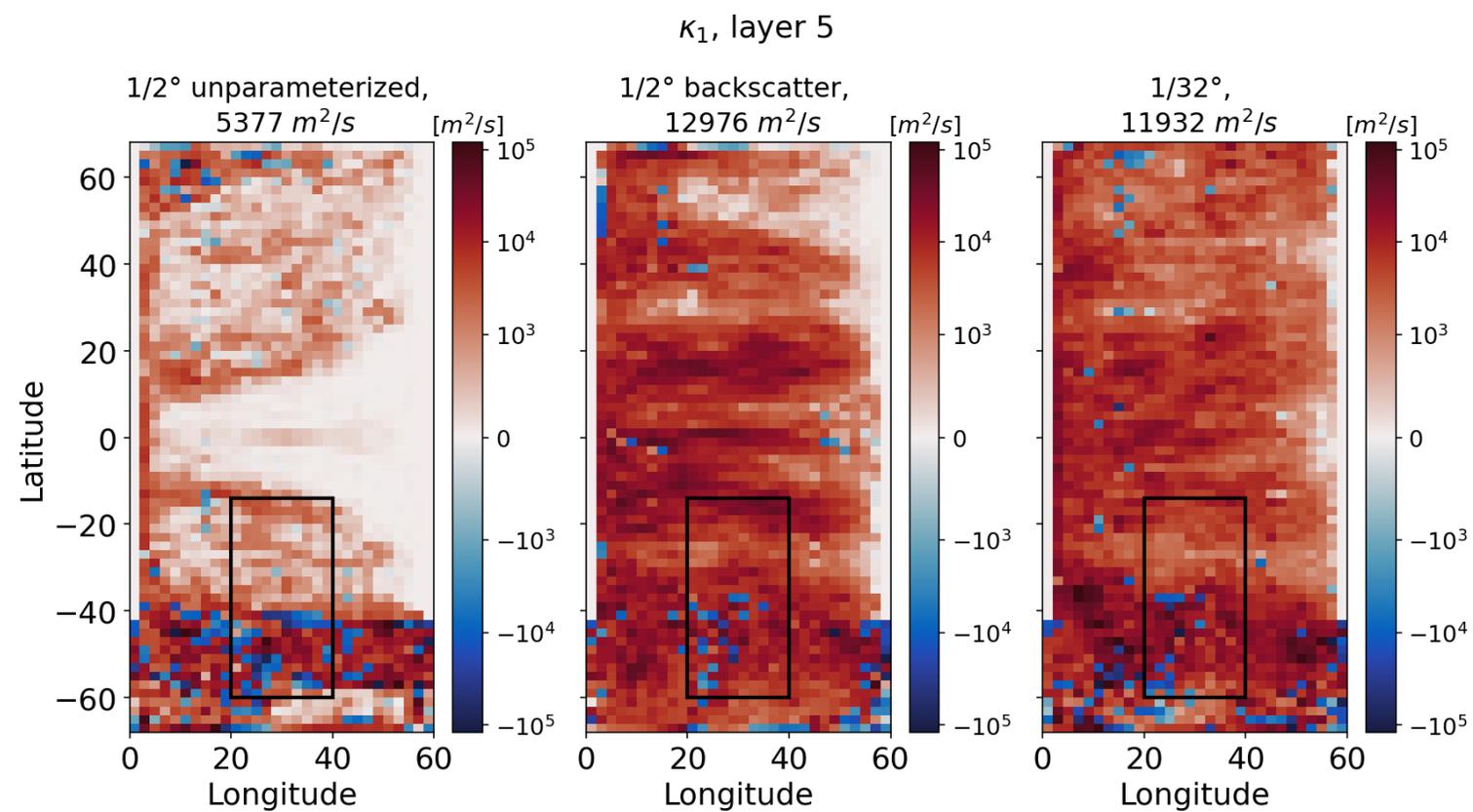
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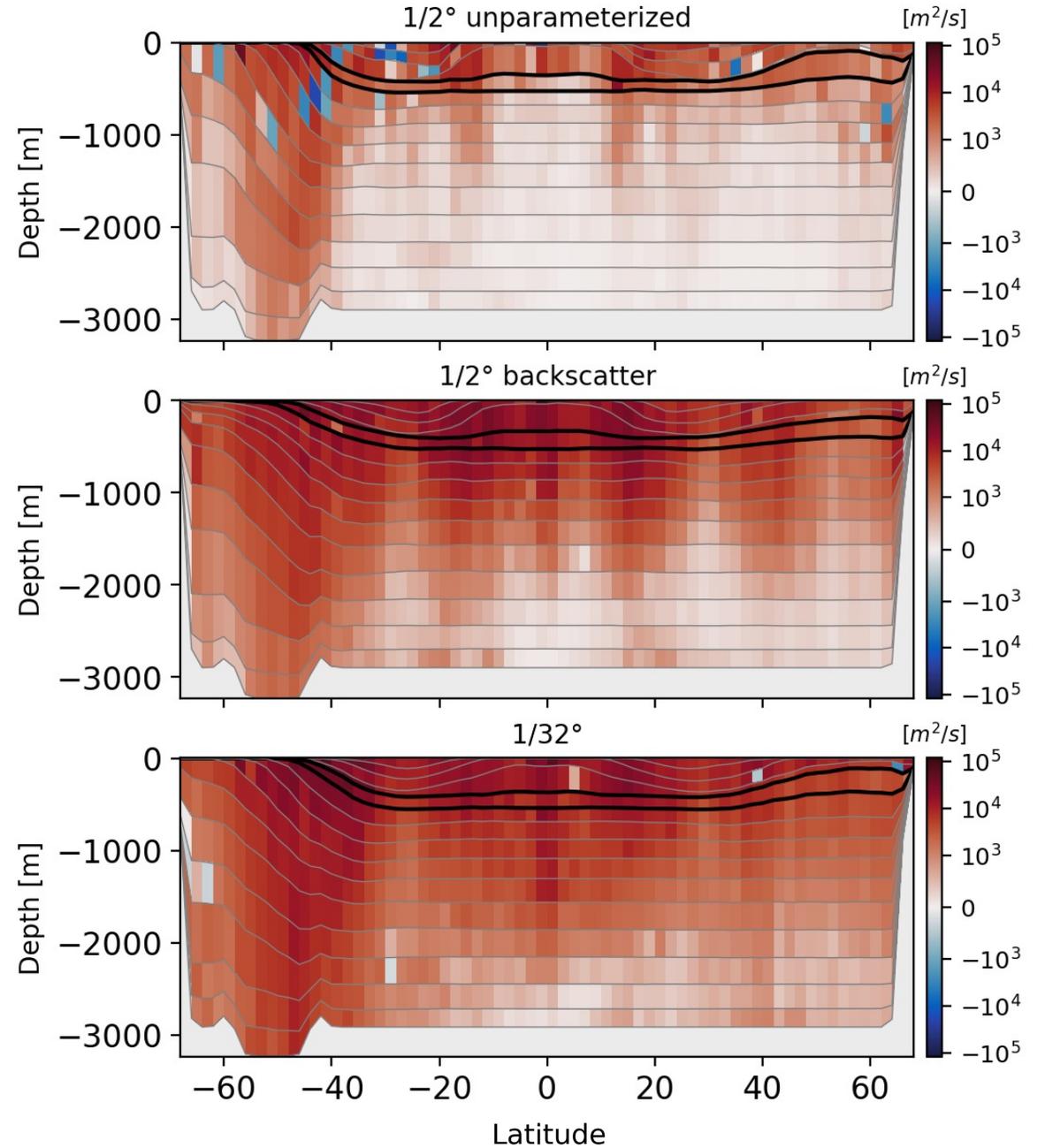


Global mean of eigenvalues (of positive values)

- 1/2° unparameterized:  
2781 m<sup>2</sup>s<sup>-1</sup>
- 1/2° backscatter:  
5341 m<sup>2</sup>s<sup>-1</sup>
- 1/32°:  
5795 m<sup>2</sup>s<sup>-1</sup>

# Tracer mixing

$\kappa_1$ , zonal average



- similar zonal structure of mixing in backscatter and hi-res, but

→ weaker at depth in Northern part of domain

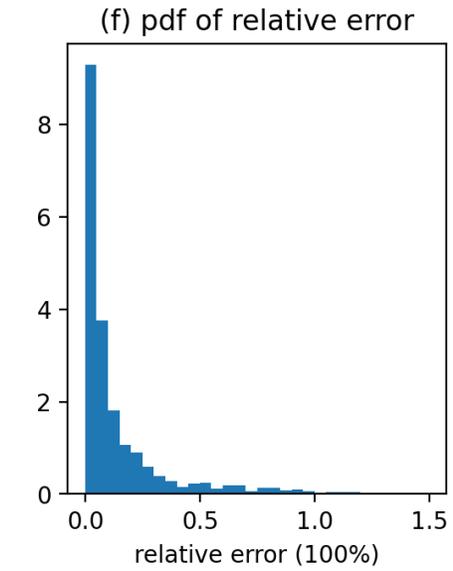
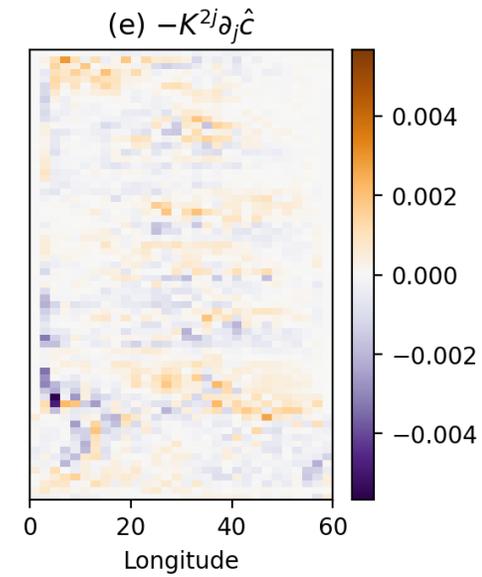
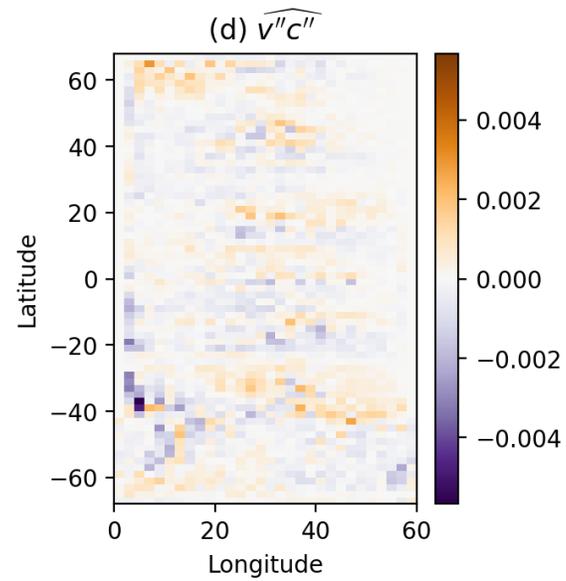
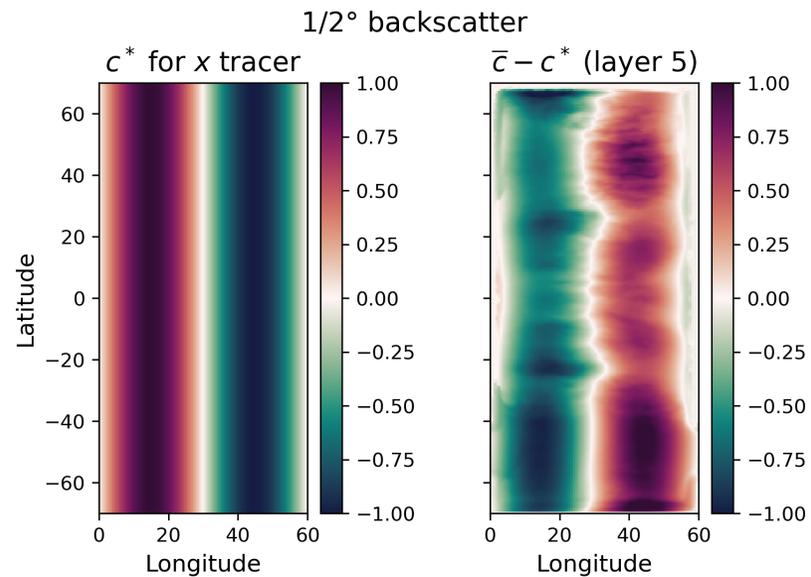
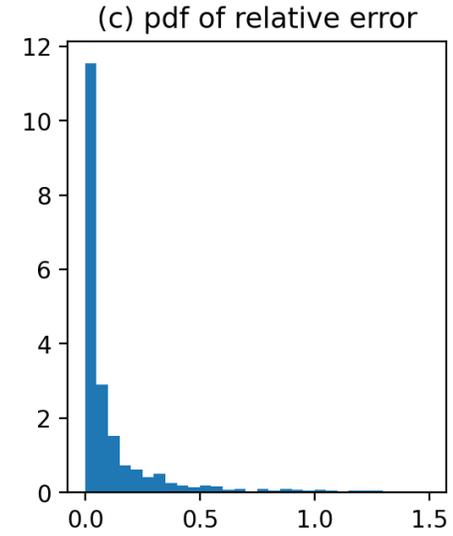
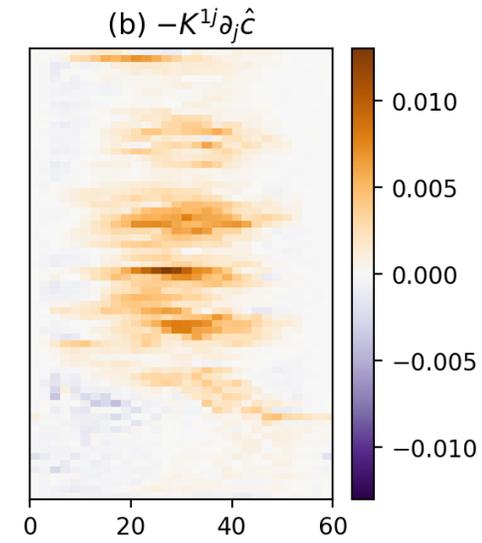
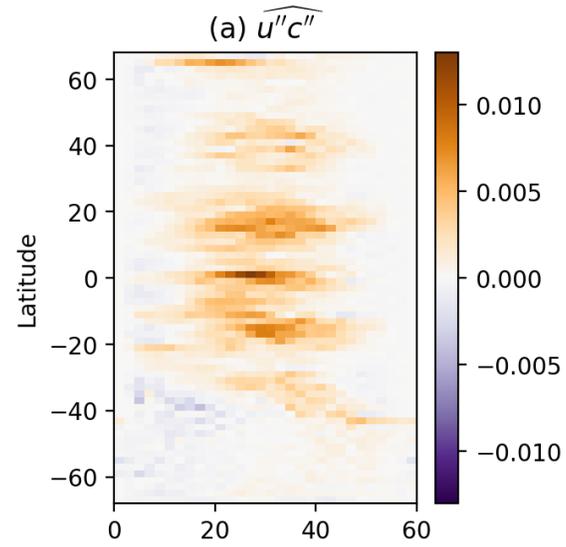
→ slightly lower magnitude in S.O.

# Flux-gradient reconstruction

Reconstruct tracer fluxes using full diffusivity tensor

→ using  $\sin(x)$  tracer, *included* in the MMT inversion

→ verify fidelity of estimated diffusivities

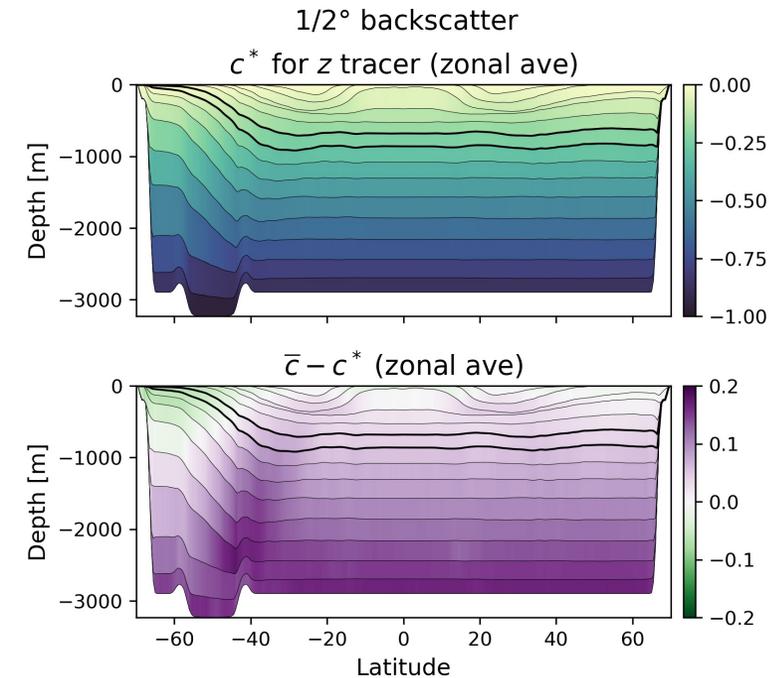
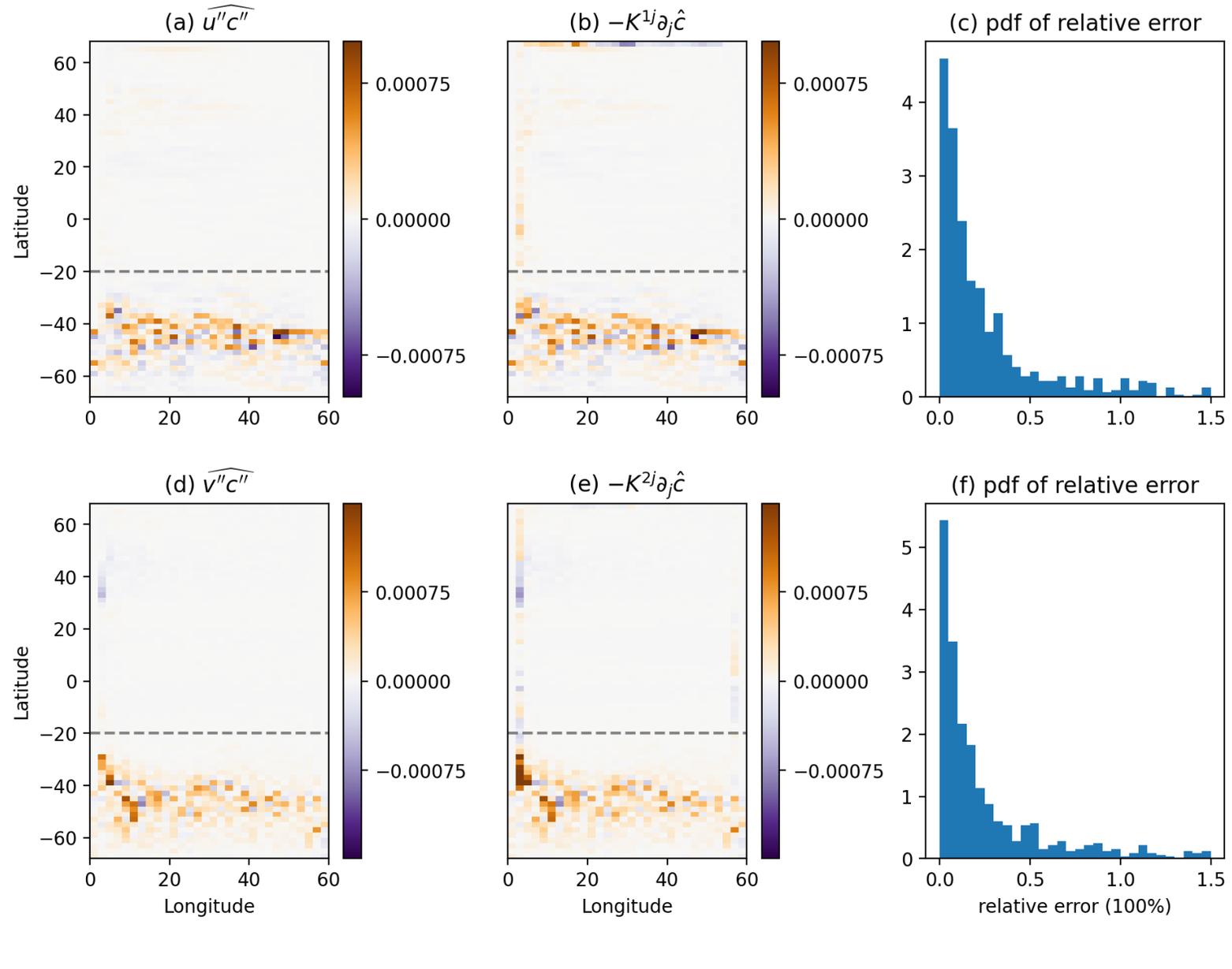


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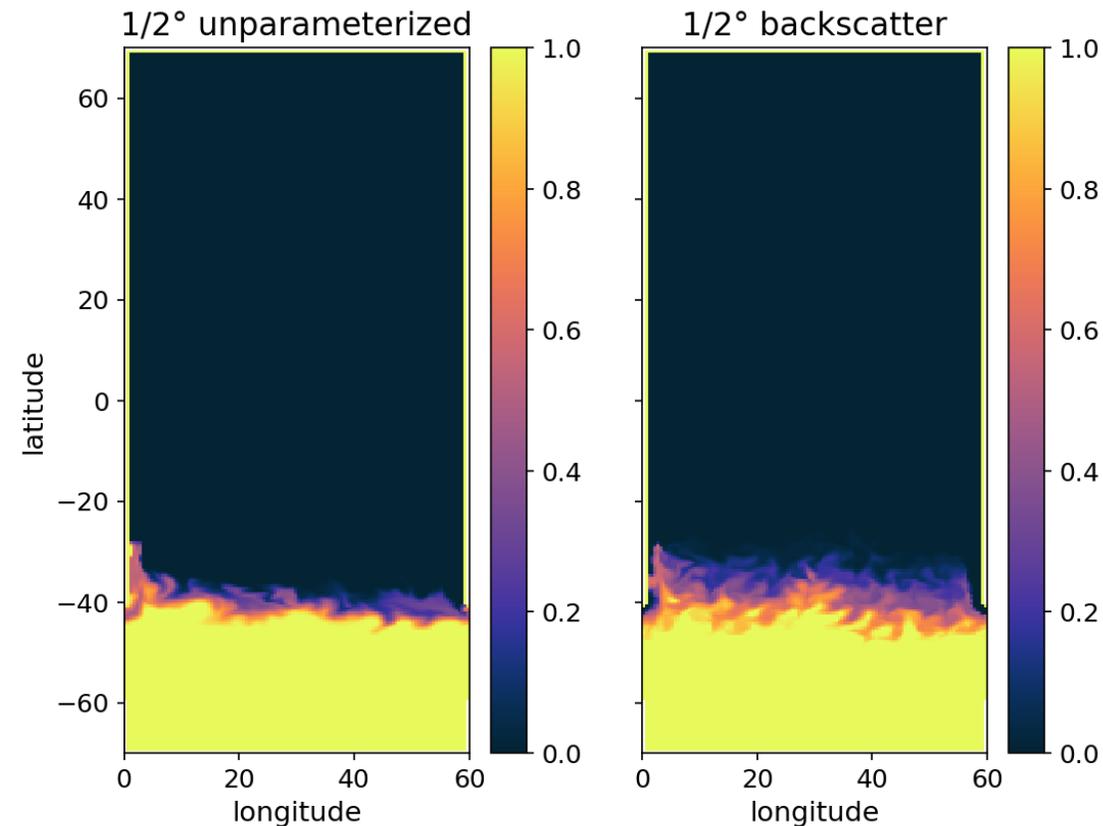
→ using  $z$  tracer, *not* included in the MMT inversion

→ fluxes/gradients confined to where isopycnals shoal

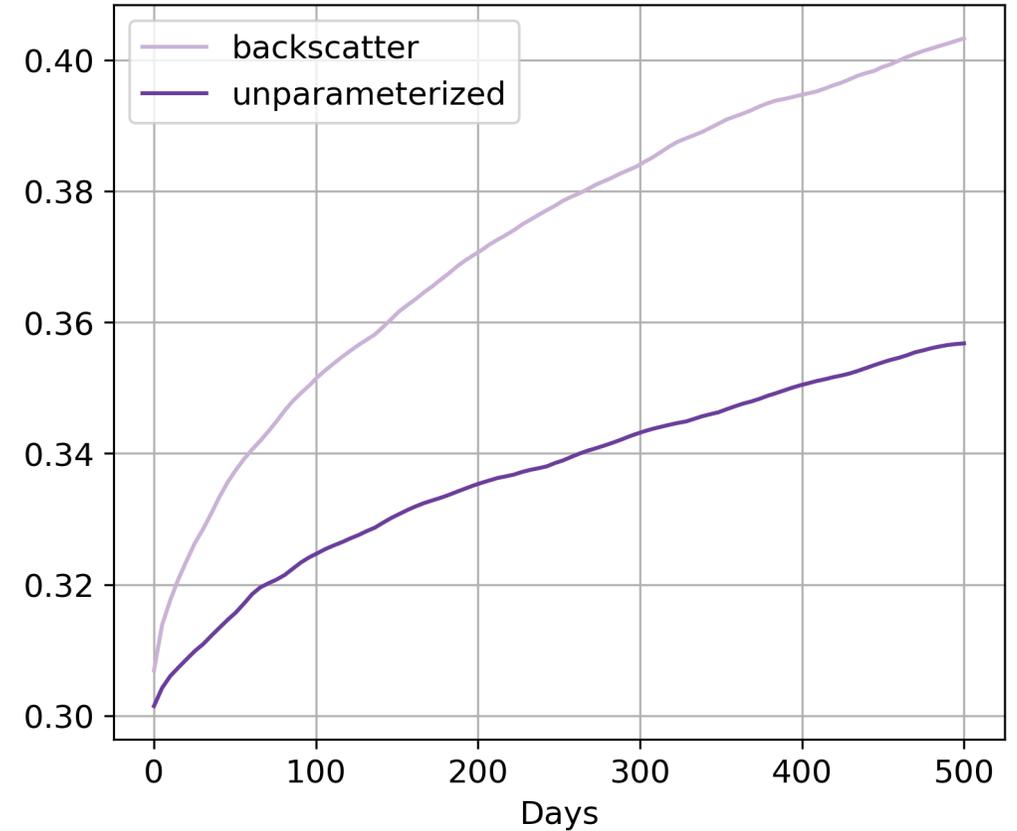


# Surface dye tracer

- Surface dye tracer to assess uptake
  - Set to 1 if layer center < 100 m depth
  - Otherwise passively advected by flow



area-averaged dye concentration on first S.O.-only outcropping isopycnal



- ~13% more uptake in backscatter simulation

# Summary and next steps

## Summary

- Backscatter parameterization with **no GM or Redi** largely induces sufficient isopycnal mixing in  $1/2^\circ$  simulation to match hi-res truth
- Multiple tracers used to estimate spatial maps of isopycnal diffusivities
- More uptake of surface tracer in backscatter than unparameterized simulation

## Next steps

- Currently examining  $1/4^\circ$  (unparameterized and backscatter)
- Other metrics to assess where backscatter is getting things right/wrong